

2ND TRAINING COURSE ON WMO SDS-WAS

(SATELLITE AND GROUND OBSERVATION AND MODELLING OF ATMOSPHERIC DUST)

21-25 November 2011, Antalya

Ground observation of mineral dust

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Index

- Aerosols and dust background
- In-situ dust monitoring (PM)
- In-situ dust estimations (Visibility)
- Ground based remote sensing
- Summary



GOBIERNO
DE ESPAÑA

MINISTERIO
DE MEDIO AMBIENTE
Y MEDIO RURAL Y MARINO

AEmet
Agencia Estatal de Meteorología



Chad-2004 by Jahi-Chikwendiu
World Press Prize 2004

Aerosols: suspended fine solid or liquid particles in a gas

- Size range of particles in the atmosphere, 0.001 to 100 μm ($1 \mu\text{m} = 10^{-6} \text{ m}$)

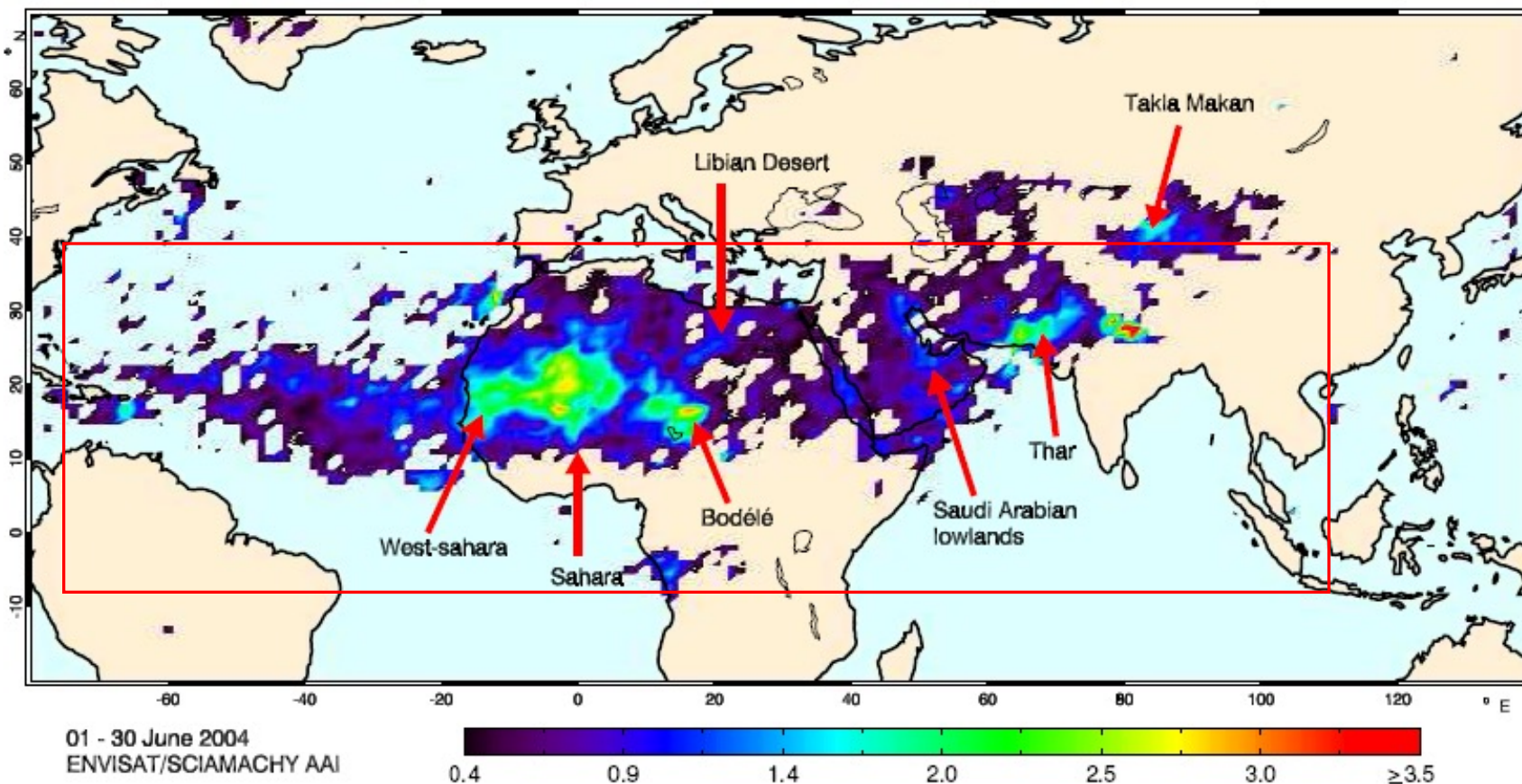
Mineral dust: Coarse aerosols

- One of the main natural sources of atmospheric aerosol particles and has been observed in the most remote regions in the world (Prospero, 1999).

Mineral dust is one of the most important tropospheric aerosols on the global scale

The global distribution is very heterogeneous

The "Global Dust Belt"



Sahara, Sahel, Arabian Peninsula, Thar desert (Middle East), Aral Sea (Central Asia), Taklamakan desert (China), Gobi Desert (China/Mongolia), Lake Eyre Basin (Australia)

(de Graaf, 2006)

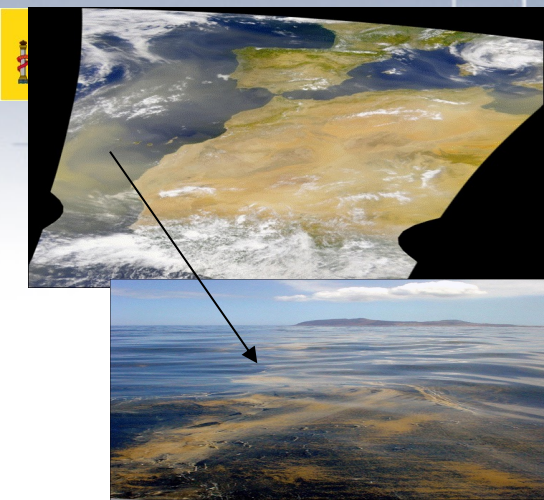
Background



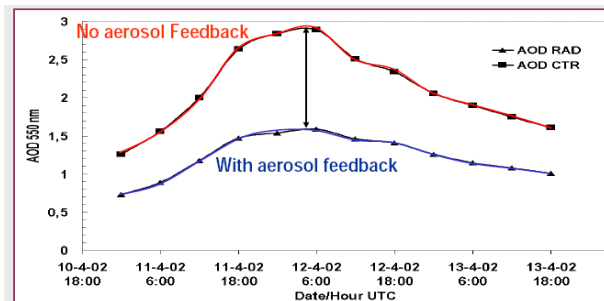
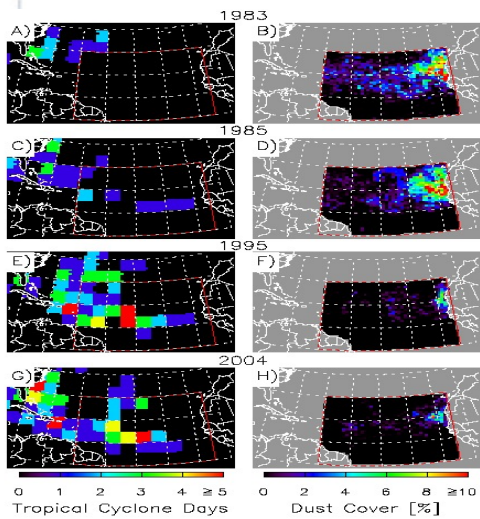
Human Health (Asthma, infections, Meningitis in Africa, Valley Fever in the America's)



Agriculture (negative & positive impacts)



Marine productivity (negative & positive impacts)



Improved Weather and Seasonal Climate prediction

Industry (Semi-conductor, etc.)

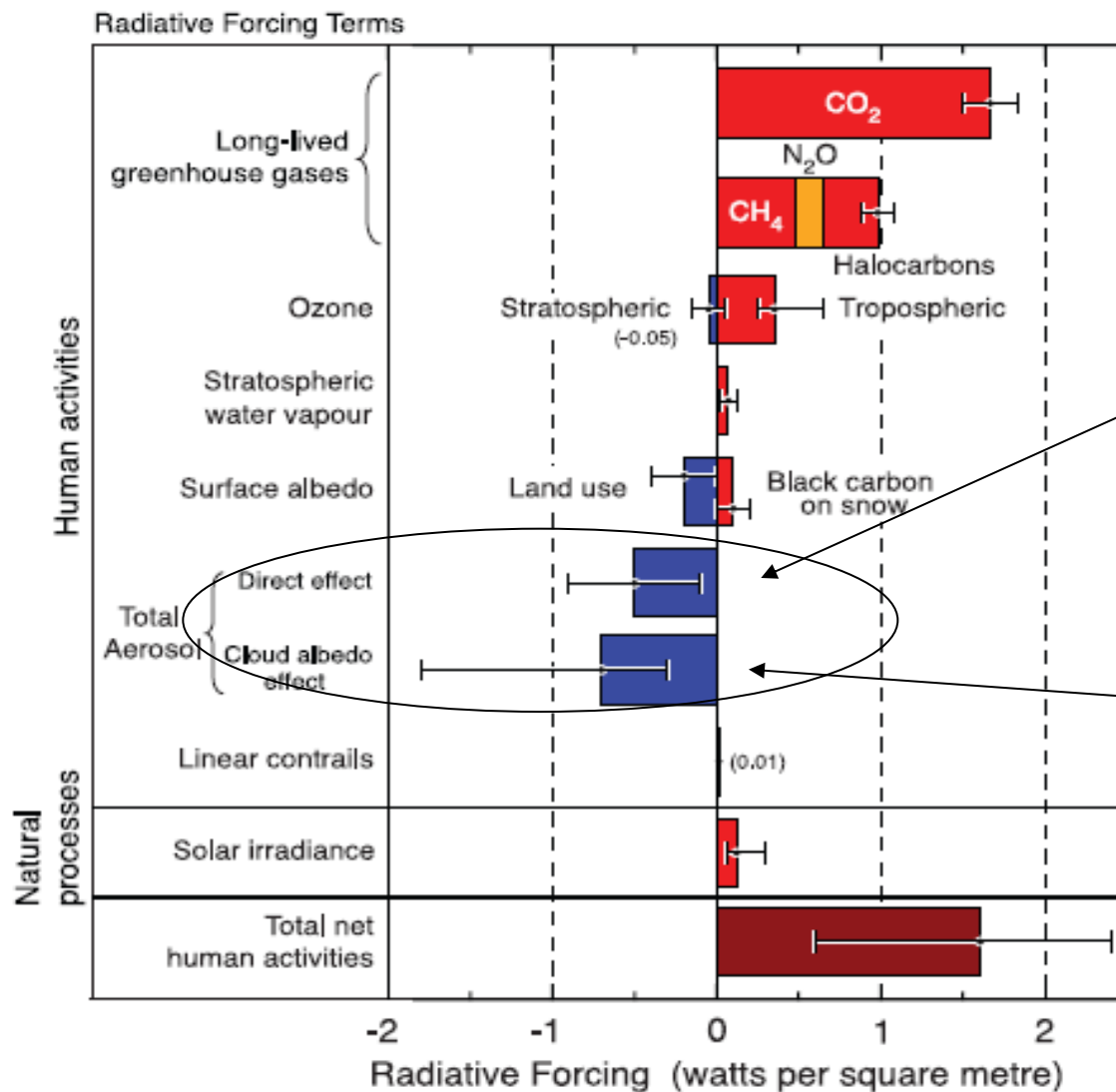
Energy (Thermal solar energy)



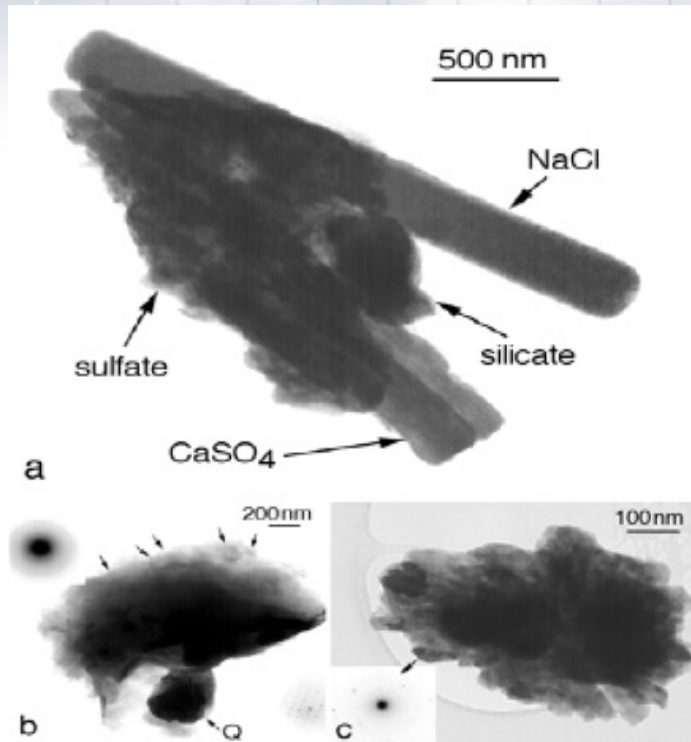
Aviation (air disasters)
Ground Transportation



Radiative forcing of climate between 1750 and 2005



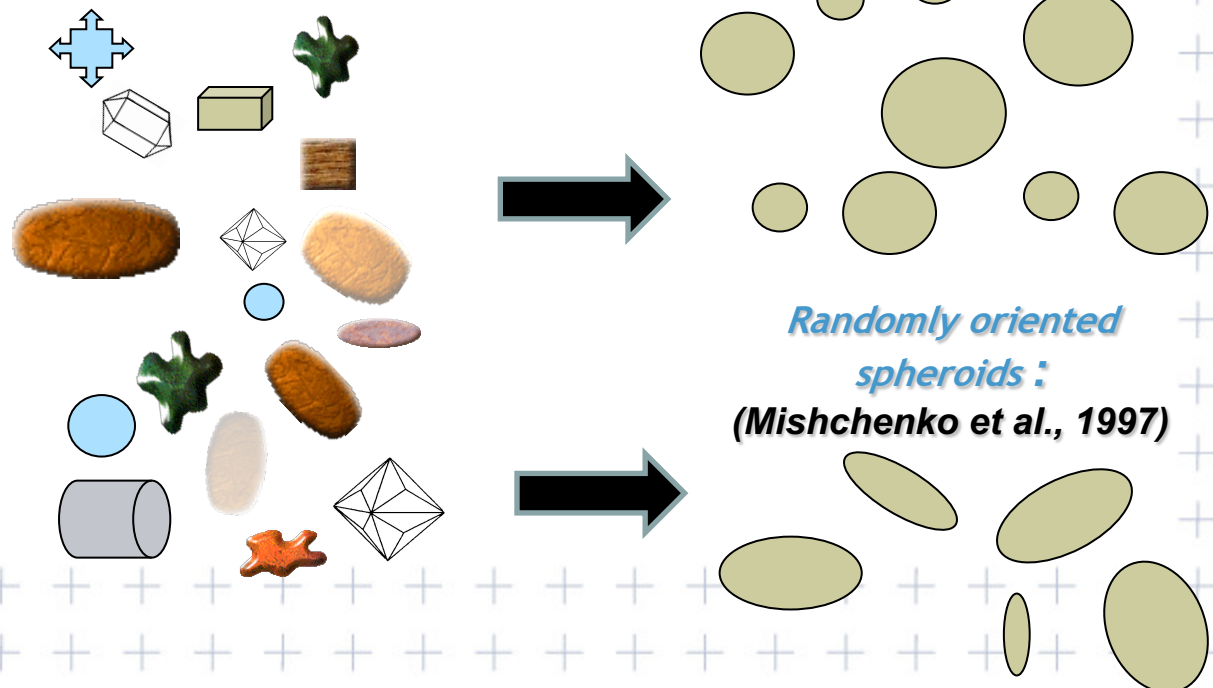
- The total direct aerosol RF as derived from models and observations is estimated to be $-0.5 [\pm 0.4] \text{ W m}^{-2}$, with a *medium-low* level of scientific understanding.
- The RF due to the cloud albedo effect (also referred to as first indirect), in the context of liquid water clouds, is estimated to be $-0.7 [-1.1, +0.4] \text{ W m}^{-2}$, with a *low* level of scientific understanding.



Busek and Posfai, 1996

Dust Particle Images

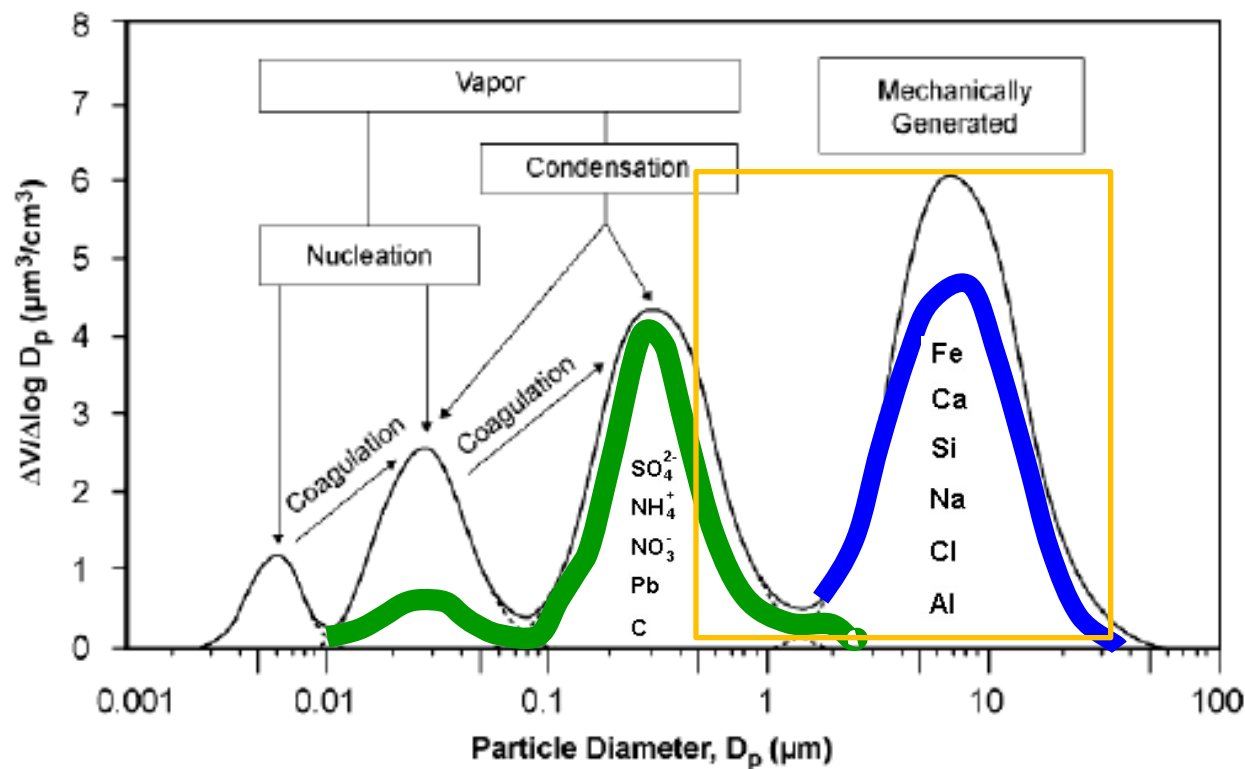
aerodynamic diameter, measured in microns or micrometers (μm), a unit equal to one millionth of a meter.



PM₁₀ (diameter <10 microm)

PM_{2.5}

PM_{2.5-10}



ultrafine
<0.1 μm

accumulation
0.1 - 1 μm

Coarse
1 - 10 μm

Mineral dust :

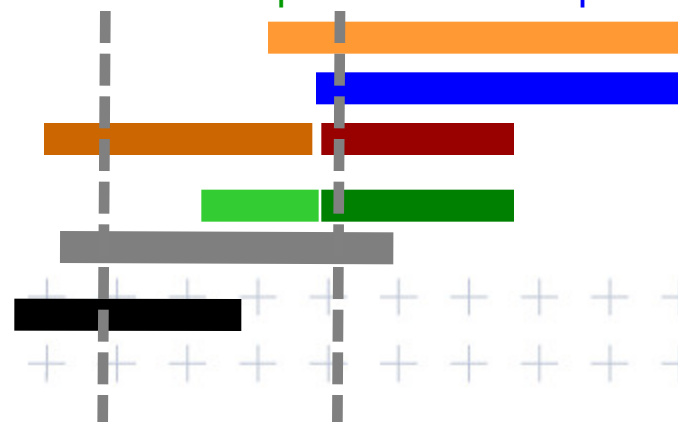
Marine salt:

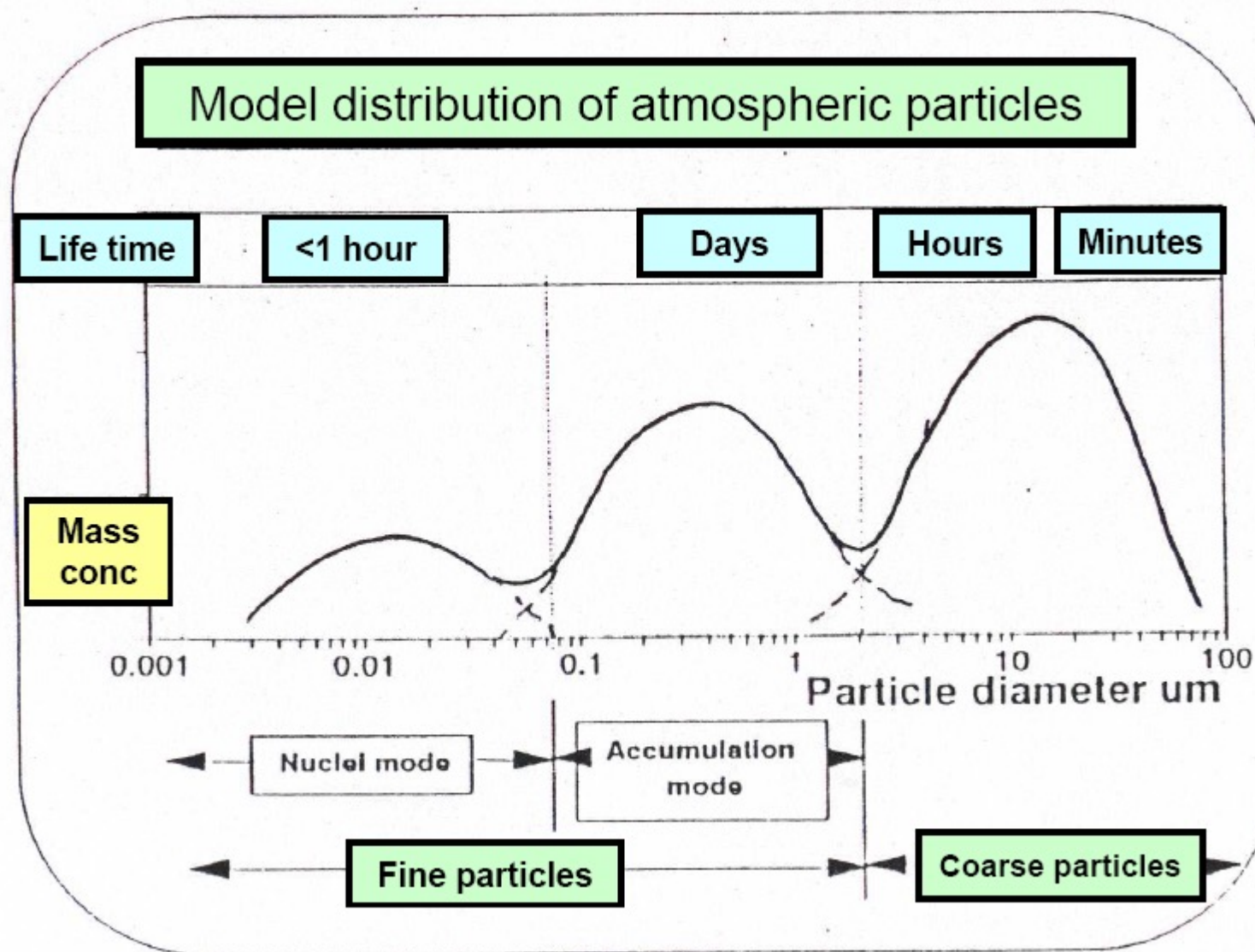
Sulfate:

Nitrate:

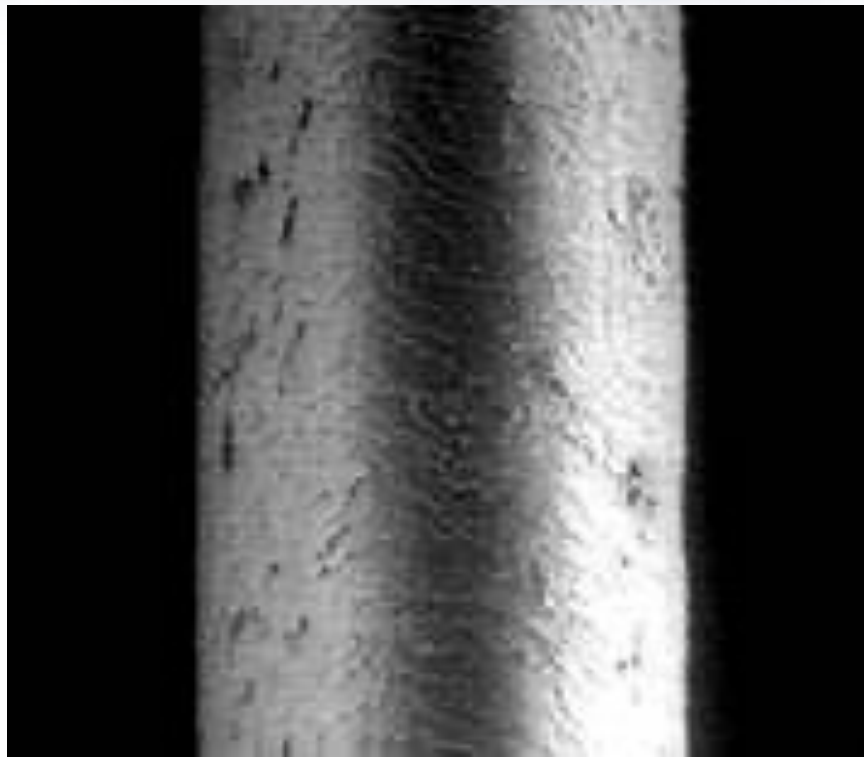
Organic aerosol:

black carbon:

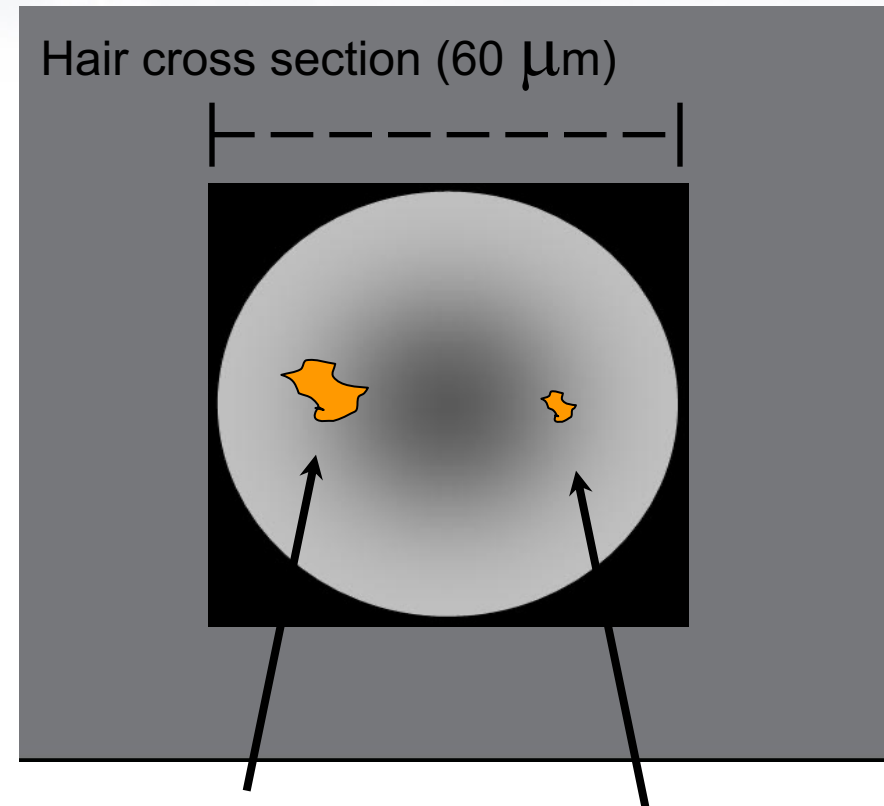




Particles in the Atmosphere: atmospheric residence time Model

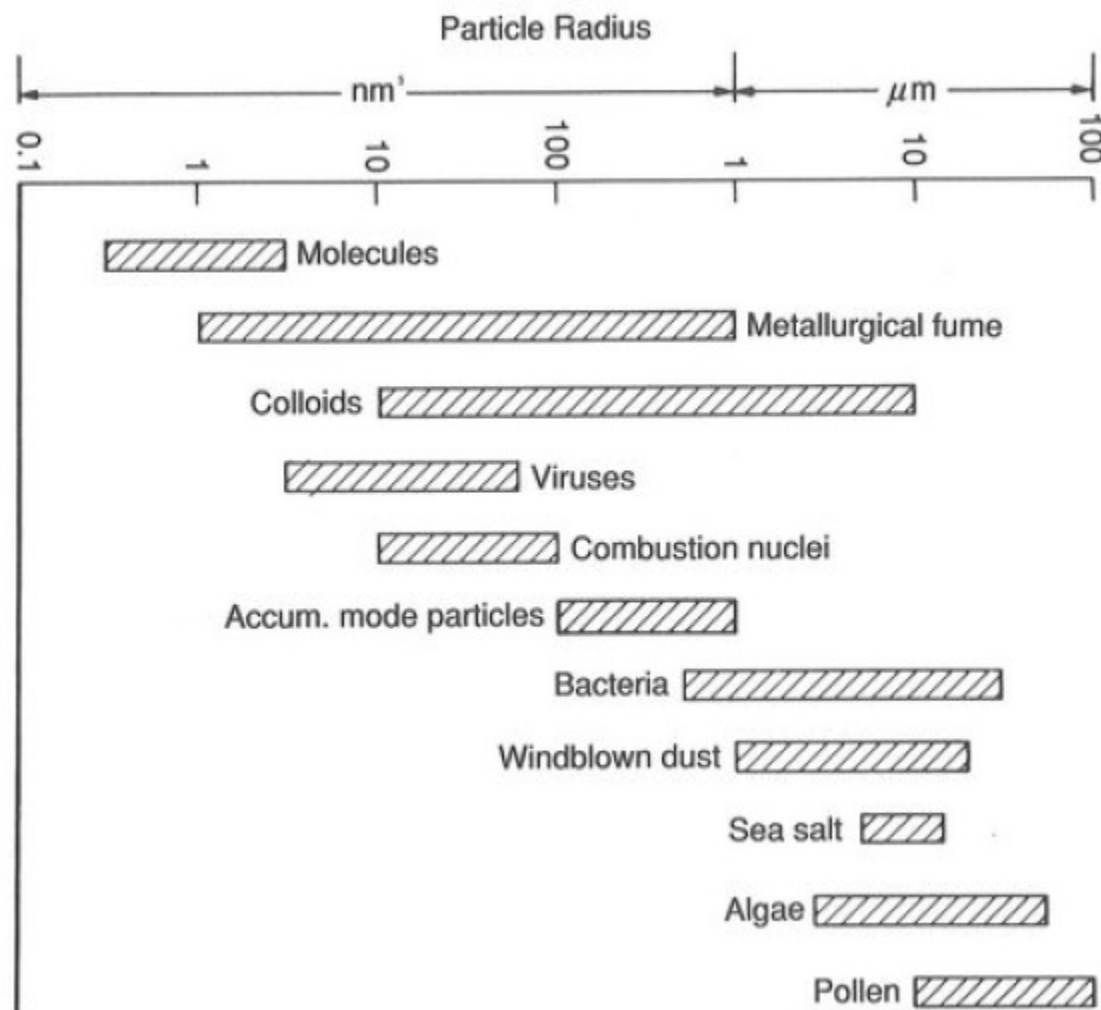


Human Hair
(60 μm diameter)



PM₁₀
(10 μm)

PM_{2.5}
(2.5 μm)



Size of different atmospheric aerosols, from (Graedel and Crutzen, 1994)

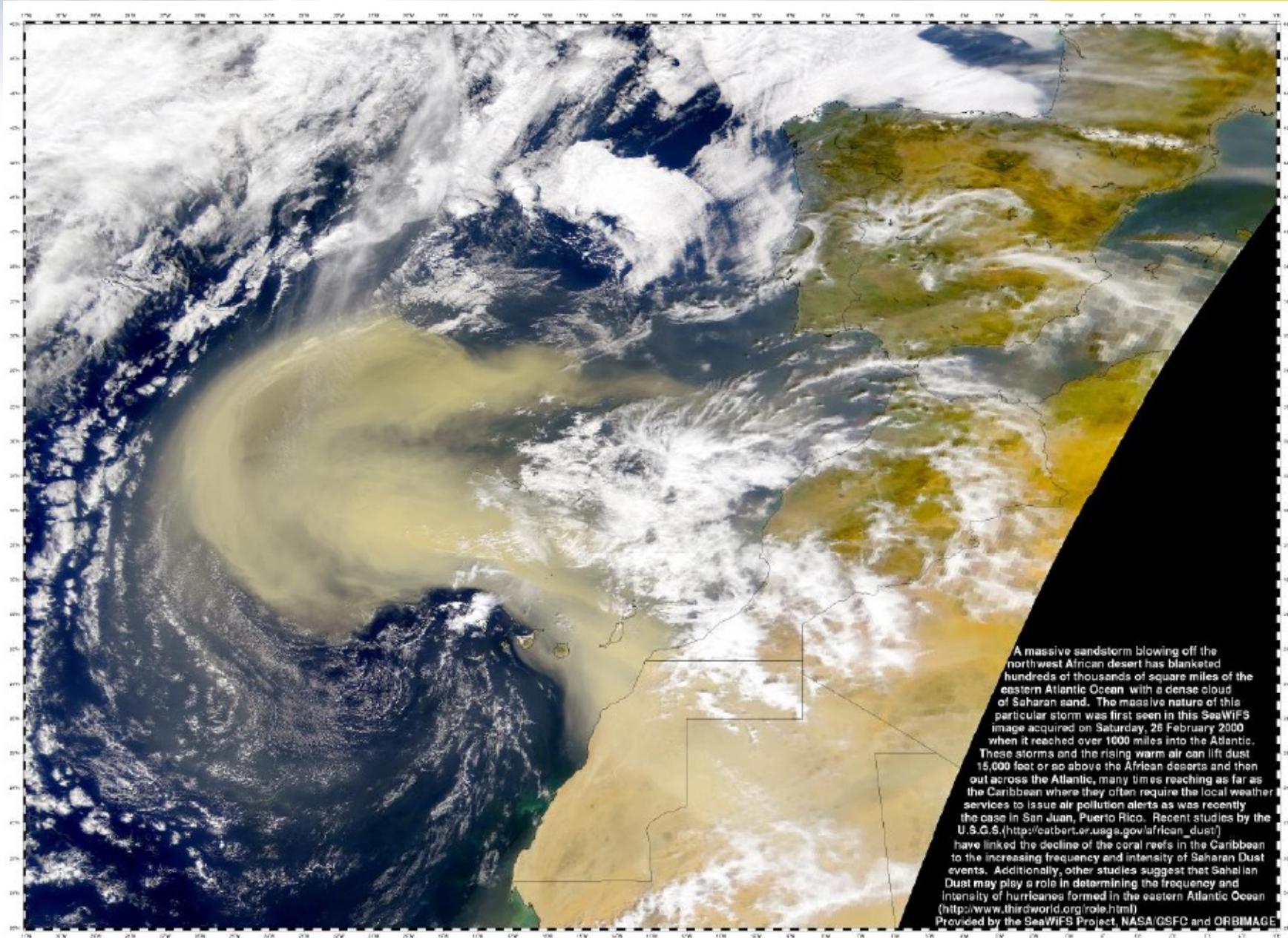
Conventionally **dust** refers to soil particles $<0.6 \text{ mm}$ ($600 \text{ }\mu\text{m}$)

but...

only those particles below 0.1 mm ($100 \text{ }\mu\text{m}$) can be transported by suspension

Dust particles move in one of three modes of transport depending on particle size, shape and density of the particle:

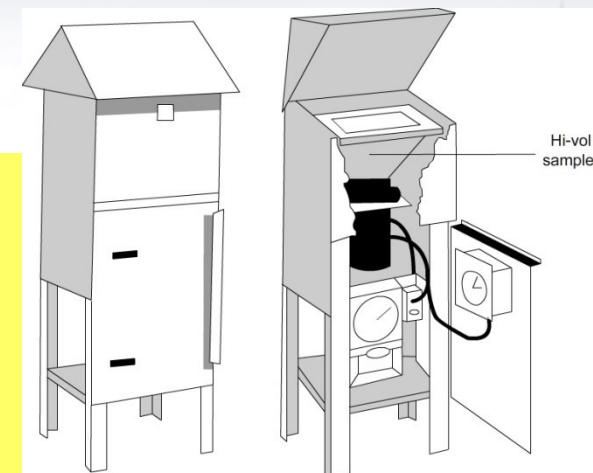
- The **suspension** mode includes dust particles $< 0.1 \text{ mm}$ ($100 \mu\text{m}$) in diameter. The fine particles may be transported at high altitudes ($6\text{-}8 \text{ km}$) and over distances of thousand kilometers.
- The **saltating** particles [i.e. 0.01 mm ($10 \mu\text{m}$) $<$ diameter $< 0.5 \text{ mm}$ ($500 \mu\text{m}$)] leave the surface, but are too large to be suspended.
- The remaining particles ($>0.5 \text{ mm}$) ($500 \mu\text{m}$) are transported in the **creep** mode. These roll along by the wind impacting particles upon the land surface, favoring the movement of other particles.



PM₁₀ and PM_{2.5} measurements in air quality networks

1. Reference method: gravimetric method

$$PM = \frac{(W2 - W1)}{\text{Volume}} \quad \mu\text{g}/\text{m}^3$$

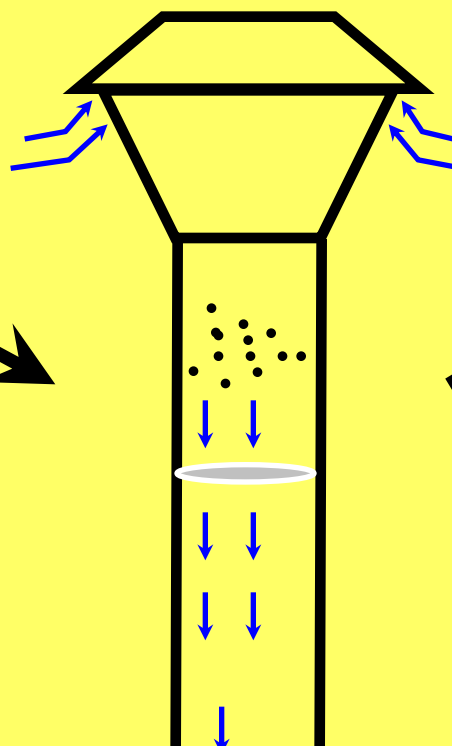


Blank filter

Conditioning

RH (50±5%) y T(20±1°C) 24-h

- Filter weight (W1)



Pump



Sampled filter

Conditioning

RH (50±5%) y T(20±1°C) 24-h

- Filter weight (W2)



Common Gravimetric Ambient Aerosol Sampling Techniques

- High volume methods: TSP, PM_{10} , $PM_{2.5}$, Air Toxics Sampler (Ultrafine particles -UFP-)
- Low volume methods: (PM_{10} , $PM_{2.5}$, PM_{Coarse})

Micro-Balance room



- Filters conditioning 48-h, $HR=50\pm5\%$ and $T=20\pm1^\circ\text{C}$
- balance, LVS resolution ≥ 5 digits (0.00001g)
- balance, HVS resolution ≥ 6 digits (0.000001g)

This sample filter is equilibrated at some set of thermodynamic conditions for a period of time before and after sampling. Through the use of a laboratory gravimetric balance, the difference in pre- and post-sample weights yields the PM mass collected. Knowing the volume of air passed through the filter allows the determination of the PM mass concentration.



PM₁₀ and PM_{2.5} measurements in air quality networks

1. Reference method: gravimetric method

LVS: 2.3 m³/h

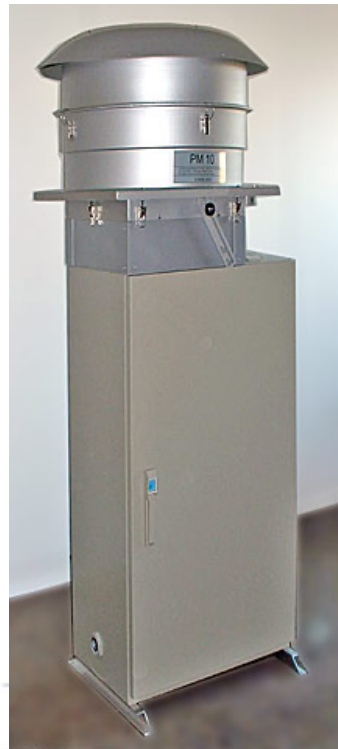
HVS: 68 m³/h

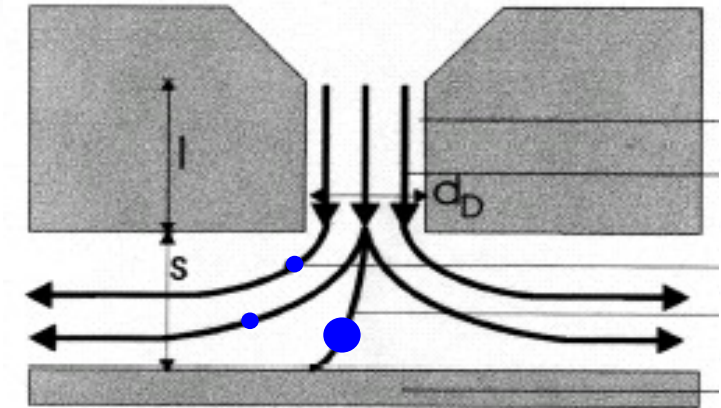
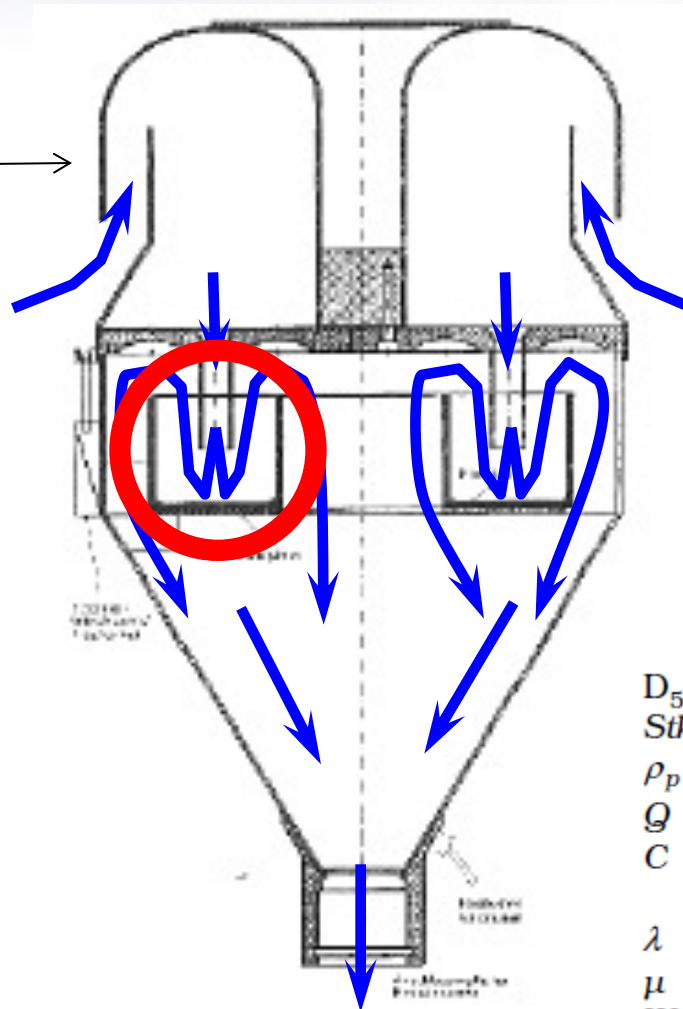
WRAC: 1966 m³/h

Low Volume Sampler

High Volume Sampler

Wide Range Aerosol Classifier





$$D_{50} = \sqrt{\frac{9\pi Stk \mu W^3}{4\rho_p CQ}}$$

- D_{50} = particle cut-point diameter (cm)
- Stk = Stokes number = 0.23
- ρ_p = particle density (g/cm^3)
- Q = volumetric flow rate (cm^3/s)
- C = Cunningham slip correction
 $= 1 + 2.492 \lambda/D_{50} + 0.84 \lambda/D_{50} \exp(-0.435 D_{50}/\lambda)$
- λ = gas mean free path
- μ = gas viscosity ($\text{dyne}\cdot\text{s}/\text{cm}^2$)
- W = nozzle diameter (cm)

The Stokes number is a dimensionless parameter that characterizes impaction.

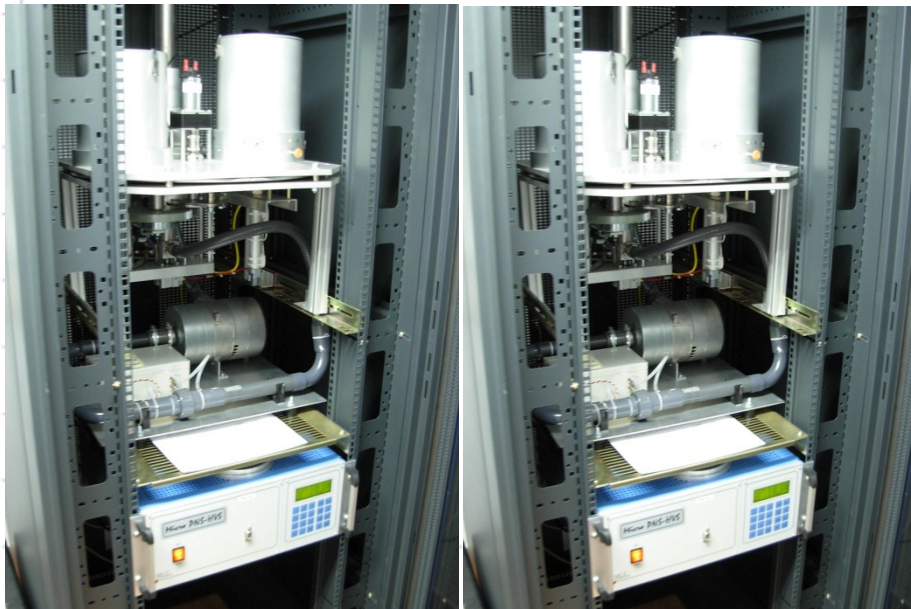
Filter

Complete PM gravimetric method set-up at Izana Atmospheric Research Center

Weight filters conditioned room

PM_{10}

$PM_{2.5}$



Common Gravimetric Ambient Aerosol Sampling Techniques

- Advantages: Recognized reference method, low capital cost
- Disadvantages: Limited time resolution (typically 24-hr), long turnaround times, labor intensive, and gravimetric lab maintenance/cost

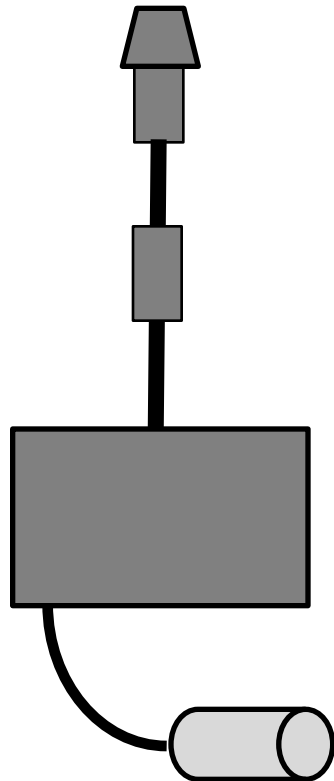
Common Continuous Ambient Aerosol Sampling Techniques

$$(Dm / Dt) / (DV / Dt) = \text{mg/m}^3$$

- Tapered Element Oscillating Microbalance
- Beta (Electron) Attenuation
- Light Scattering, Absorption, and Extinction

PM₁₀ and PM_{2.5} measurements in air quality networks

2. Automated analyzers



1. Impactor PM₁₀ / PM_{2.5}

2. RH reductor / heater

3. Sensor (Beta radiation attenuation or Tapered Oscillating microbalance-TEOM-) → instead of weighting filters

4. Pump / Flow meter

Continuous measurements on PM (PM₁₀, PM_{2.5}, PM₁ or TSP)

PM with Tapered Oscillating microbalance-TEOM- (1)

The system can be considered a simple harmonic oscillator through which the following equation can be derived:

$$\Delta m = K_0 (1/f_f^2 - 1/f_i^2)$$

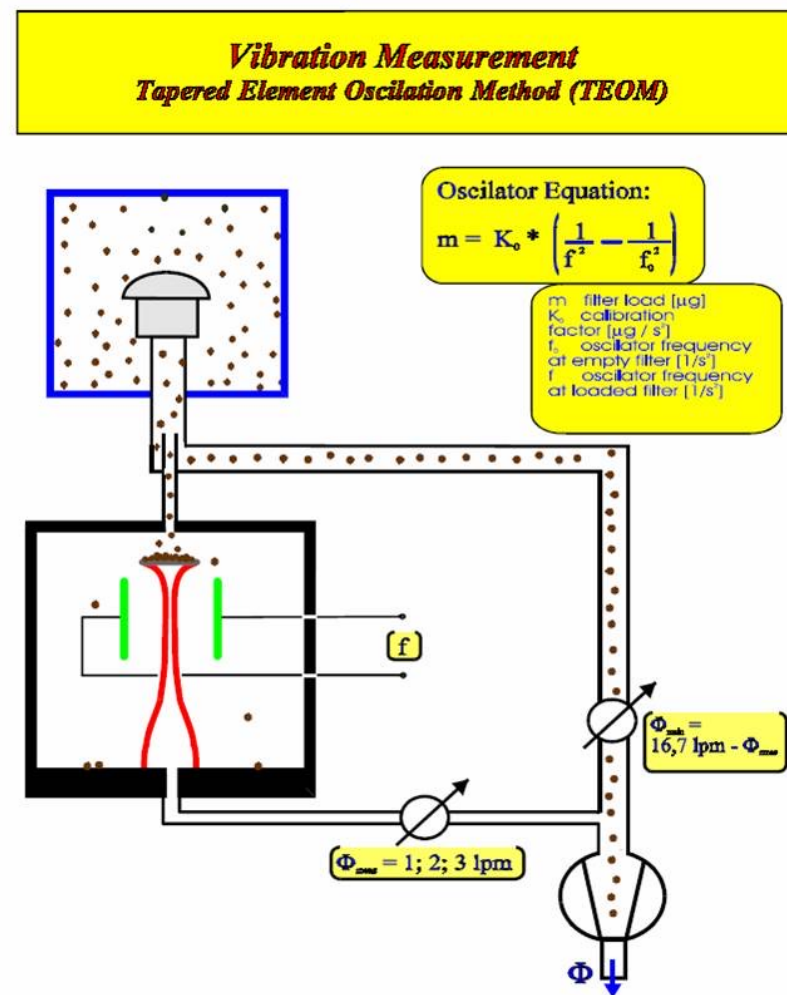
Where:

f_i = initial oscillation frequency of the system;

f_f = oscillation frequency after the addition of mass;

K_0 = calibration (spring) constant of the initial condition.

TEOM mass detectors or microbalances utilize an inertial mass weighting principle.



PM with Tapered Oscillating microbalance-TEOM- (2)

Advantages

- Continuous method
- Highly time resolved
- High resolution
- instantaneous turnaround

Disadvantages

Temperature dependency:

- Volatile losses

- Seasonal and regional dependencies

Affected by vibration

Manual filter changes necessary

Complex systems require some skill

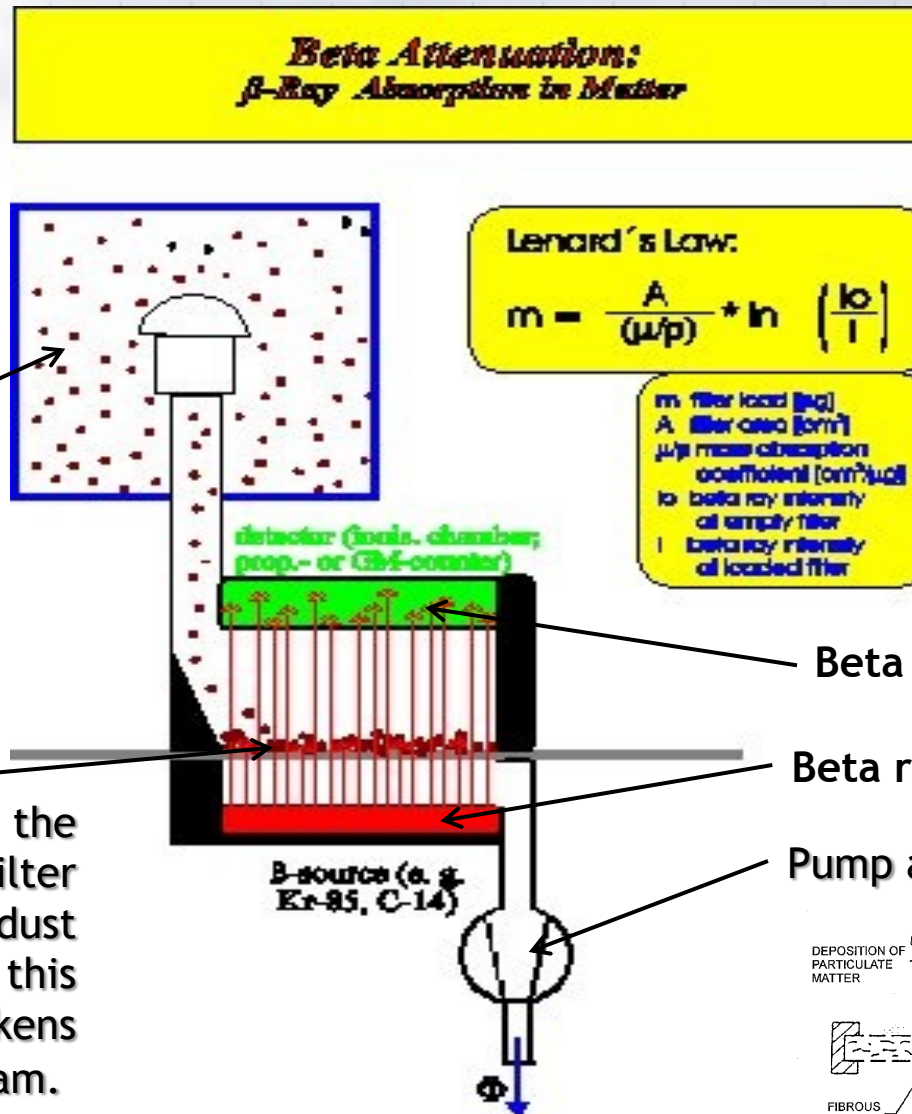
X2 or X3 capita cost

External calibration against reference method

Beta rays are radiation emitted by electrons during the nuclear decay of radioactive elements. Elements such as Krypton-85 or Carbon-14 may be used.

Ambient air is sucked through the sample system

Dust particles contained in the air are deposited on the filter continuously. The layer of dust is building up and this increasing dust mass weakens the intensity of the beta beam.



The absorption of beta rays through the filter and the deposited Particulate matter is described by the Lenard's law

PM with Beta attenuation (2)

$$m = F_{cal} \ln \left(\frac{I_0}{I} \right)$$

- **m**: increasing particle mass [μg]
- **F_{cal}**: calibration factor
- **I₀** beta ray intensity at empty filter
- **I** beta ray intensity at loaded filter

The intensities I_0 and I are measured with the detector system. F_{cal} has to be measured directly during the calibration procedure. This is accomplished by replacing the filter with the element having a known mass (mass calibration kit)

The mass concentration is calculated from:

$$c = \frac{m}{Ft}$$

Where:

c: concentration [$\mu\text{g}/\text{m}^3$]

F: measured air flow [m^3/h]

t: time [h]

Common Continuous Ambient Aerosol Sampling Techniques

$$(Dm / Dt) / (DV / Dt) = \text{mg/m}^3$$

Advantages

- Continuous method
- Highly time resolved
- High resolution
- instantaneous turnaround
- Low operational cost

Disadvantages

Temperature dependency:

- Volatile losses

- Seasonal and regional dependencies

- Affected by vibration

- Manual filter changes necessary

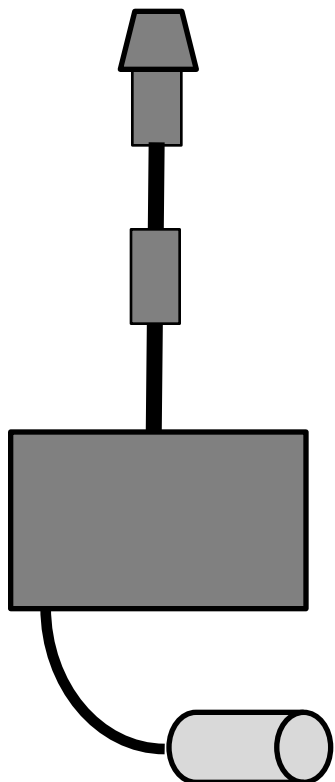
- Complex systems require some skill

- X2 or X3 capita cost

- External calibration against reference method

PM₁₀ and PM_{2.5} measurements in air quality networks

2. Automated analyzers



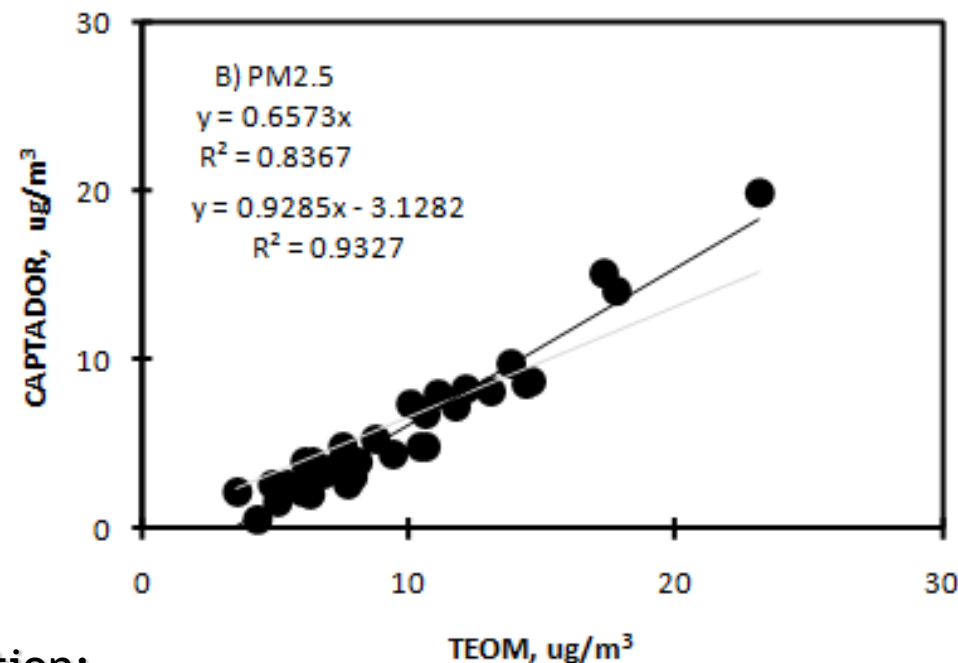
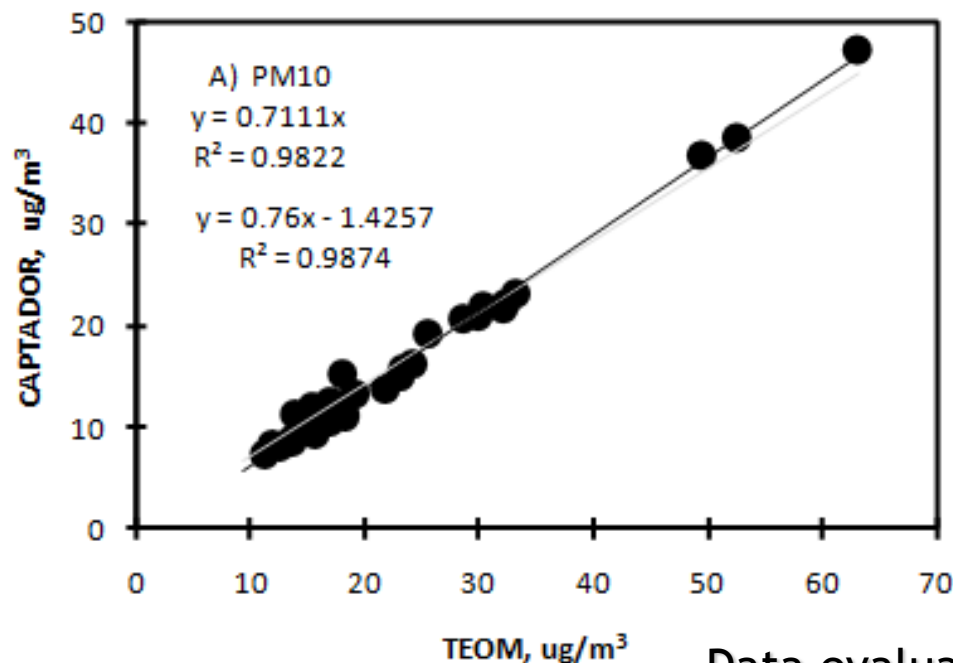
beta



TEOM

PM₁₀ and PM_{2.5} measurements in air quality networks

Intercomparisons



Data evaluation:

Data from continuous analyzer are valid if they fit A or B:

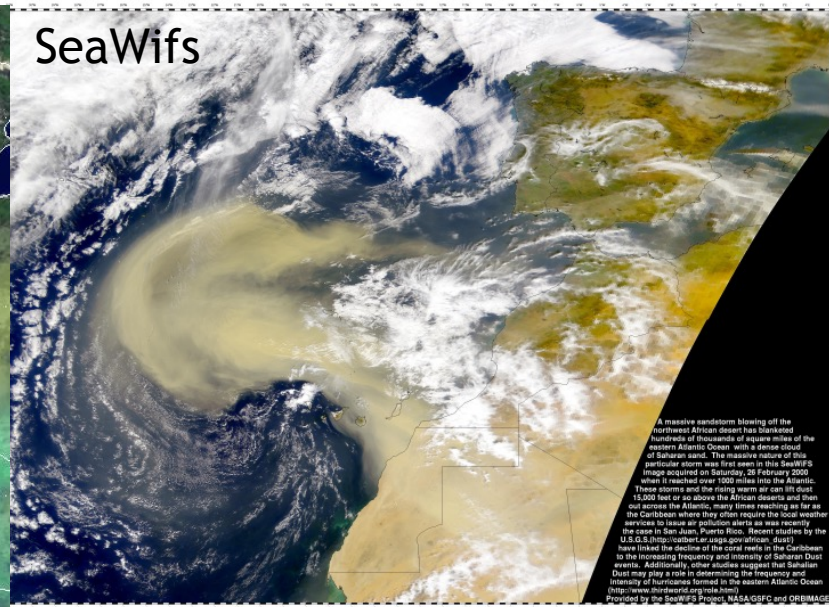
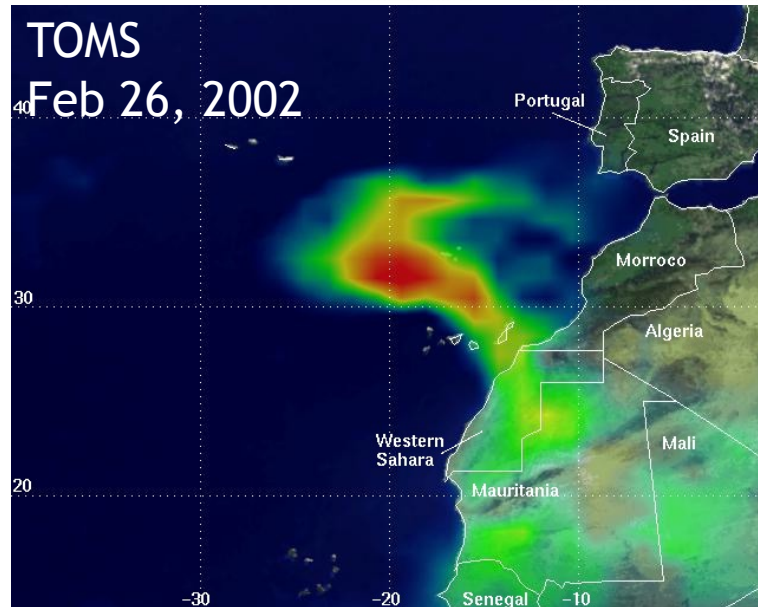
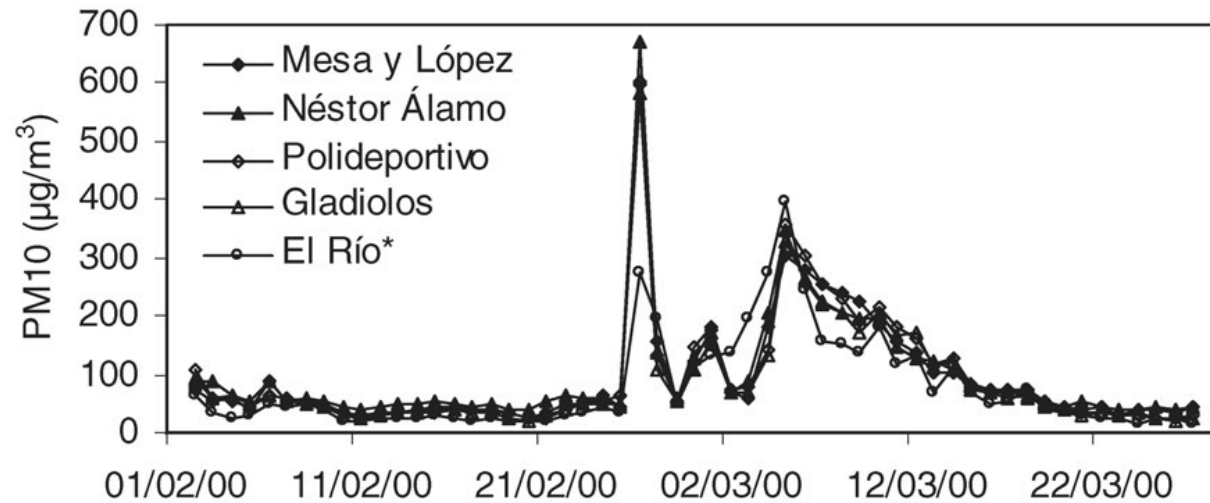
A) $Y = a \cdot X$; $r^2 \geq 0.8$

B) $Y = a \cdot X + b$; $r^2 \geq 0.8$; $\text{abs}(b) < 5$

Y= Reference Method (gravimetric method),

X= Automatic analyzer

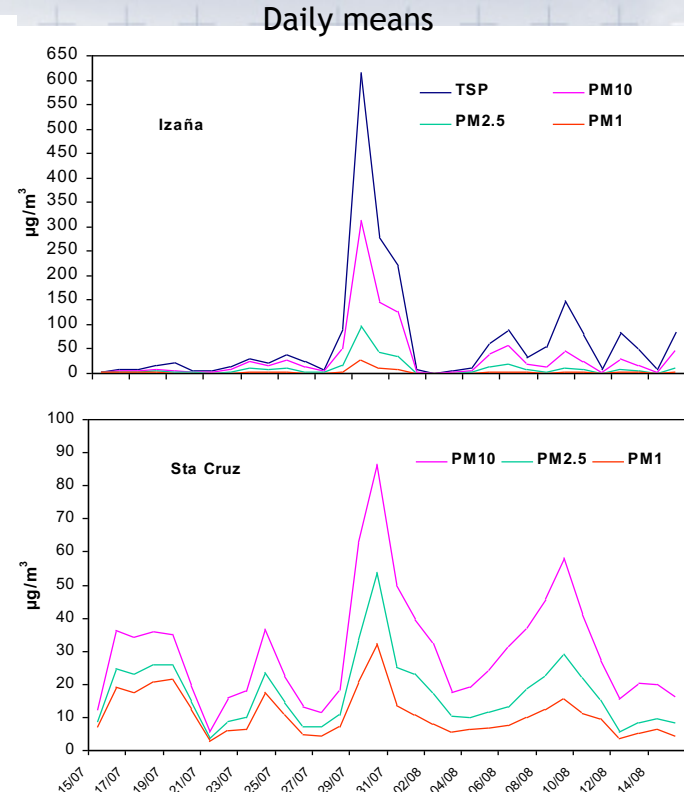
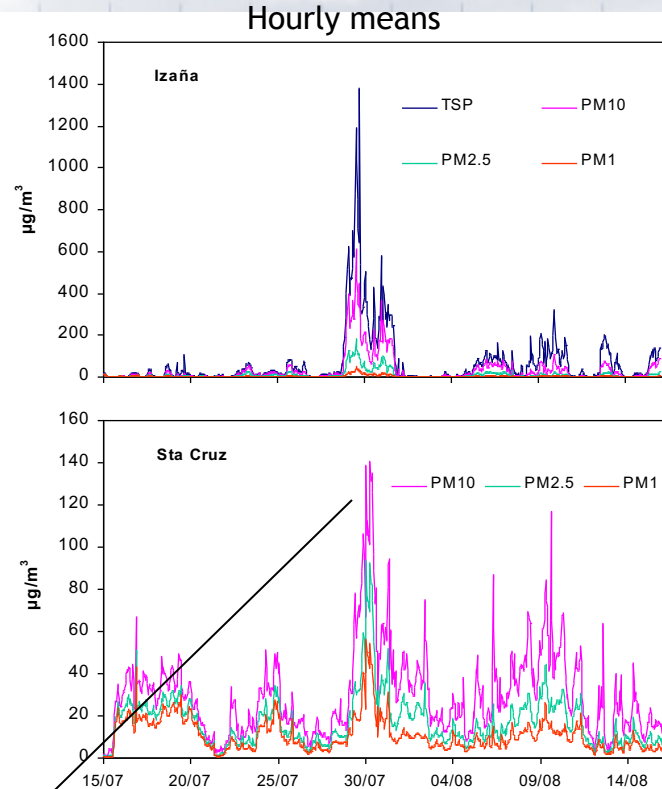
Air quality stations at Tenerife Island



Viana et al., Atmospheric Environment, 2002

Chemical characterisation of TSP and PM_{2.5} at Izaña and Sta. Cruz during the MINATROC field campaign (July, 2002)

Alastuey et al., 2003



Chemical Analysis: methodology (CSIC; Alastauy and Querol)
60 parameters/filter

ACIDIC DIGESTION

Inductively coupled plasma
Atomic Emission Spectroscopy

Major elements

ICP-AES

| | | | |
|--------------------------------|---------|---|----|
| Al ₂ O ₃ | Ca | K | Mg |
| Fe | Ti | P | Na |
| CO ₃ ²⁻ | ind. Ca | | |
| SiO ₂ | ind. Al | | |
| SO ₄ ²⁻ | ind. Na | | |

Inductively coupled plasma
Mass spectroscopy

Trace elements (40)

ICP-MS

As, Ba, Bi, Cd, Ce, Co, Cr, Cs, Cu, Dy, Er,
Ga, Gd, Ge, Hf, La, Li, Mn, Mo, Nd, Ni,
Pb, Pr, Rb, Sb, Sc, Se, Sm, Sn, Sr, Ta,
Th, Ti, Tl, U, V, W, Yb, Zn, Zr

WATER EXTRACTION

Soluble anions and cations

| | |
|-------------------------------|-----------------------|
| NH ₄ ⁺ | FIA C. |
| SO ₄ ²⁻ | IC |
| Cl ⁻ | IC Ion Chromatography |
| NO ₃ ⁻ | IC |

ELEMENTAL ANALYSIS

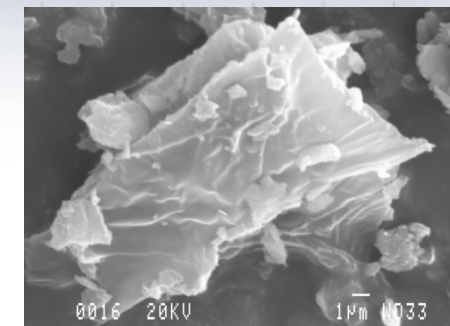
C Elemental A.

DETERMINED

75-85 % PM mass

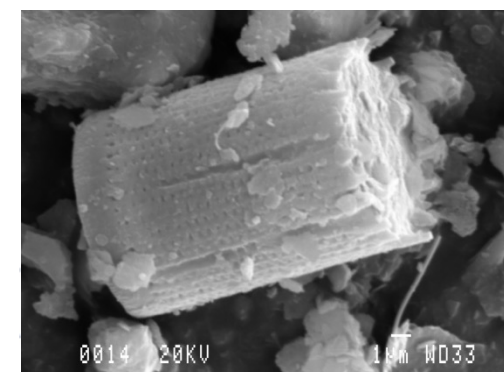
In-situ measurements (PM)

| | A) PM10 | | | | B) PM2.5 | | | |
|-------------------------------|---------------------|----------------------------------|--------------------|------------------------------------|---------------------|----------------------------------|------------------------------------|--------------------|
| | Factor 1 mineral | Factor 2 primario fuel oil | Factor 3 marino | Factor 4 secundario fuel oil | Factor 1 mineral | Factor 2 primario fuel oil | Factor 3 secundario fuel oil | Factor 4 marino |
| OM+EC | | 0.75 | | | | 0.67 | | |
| Al | 0.91 | | | | 0.88 | | | |
| Ca | 0.81 | | | | 0.71 | | | 0.47 |
| K | 0.83 | | 0.45 | | 0.78 | | | |
| Na | | | 0.90 | | | | | 0.71 |
| Mg | 0.59 | | 0.76 | | 0.66 | | | 0.61 |
| Fe | 0.91 | | | | 0.89 | | | |
| SO ₄ ²⁻ | 0.41 | 0.40 | 0.54 | 0.57 | 0.49 | 0.61 | 0.51 | |
| NO ₃ ⁻ | 0.47 | | 0.55 | 0.46 | 0.64 | | | 0.44 |
| Cl ⁻ | | | 0.88 | | | | | 0.88 |
| NH ₄ ⁺ | | | | 0.86 | | 0.53 | 0.63 | |
| P | 0.80 | | | | 0.47 | | | |
| Ti | 0.94 | | | | 0.94 | | | |
| V | | 0.90 | | | | 0.94 | | |
| Mn | 0.92 | | | | 0.78 | | | |
| Co | 0.55 | 0.52 | | | 0.40 | 0.61 | | |
| Ni | | 0.85 | | | | 0.94 | | |
| Cu | 0.53 | | | | | | 0.61 | |
| Zn | | | | 0.51 | | | 0.87 | |
| As | 0.54 | 0.43 | | | 0.53 | 0.56 | | |
| Cd | 0.75 | | | | 0.64 | | 0.50 | |
| Pb | 0.51 | | 0.42 | | 0.67 | | | |
| Exp Var | 8.09 | 3.40 | 3.72 | 2.25 | 7.30 | 4.68 | 2.90 | 2.50 |
| Pm Totl | 0.37 | 0.15 | 0.17 | 0.10 | 0.33 | 0.21 | 0.13 | 0.11 |



• Saharan air masses mainly consist of clays, quartz, feldspars, and calcite.

• Fresh water diatoms and pollen particles from North Africa have also been identified at Izaña.

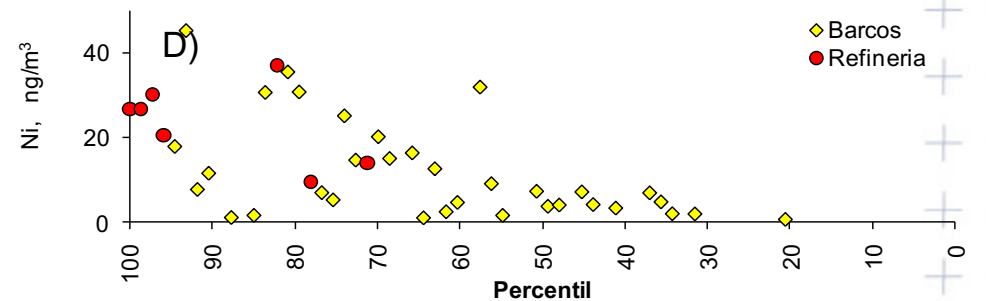
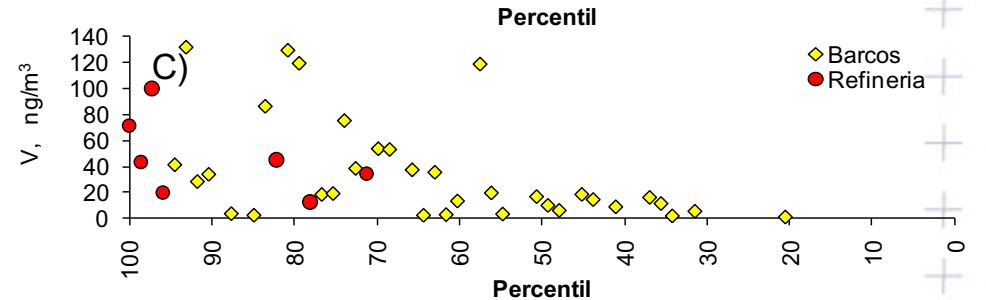
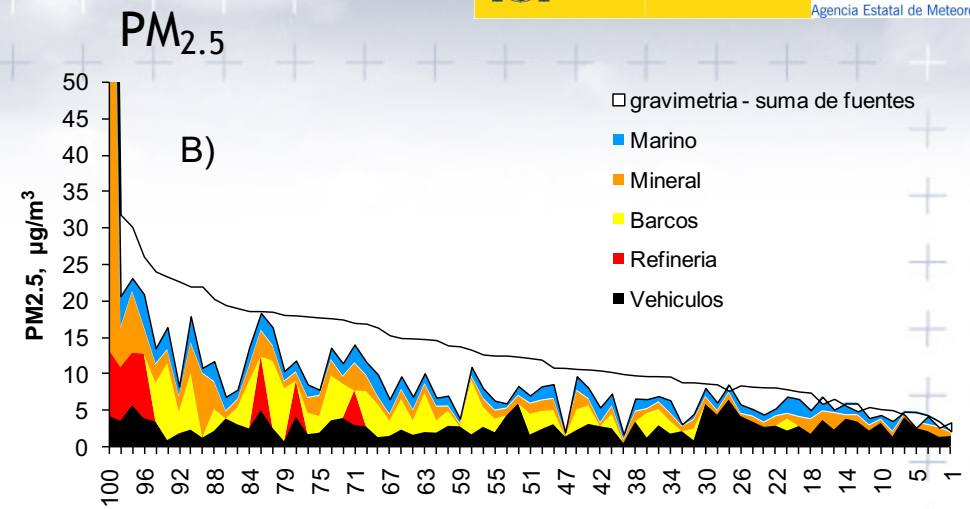
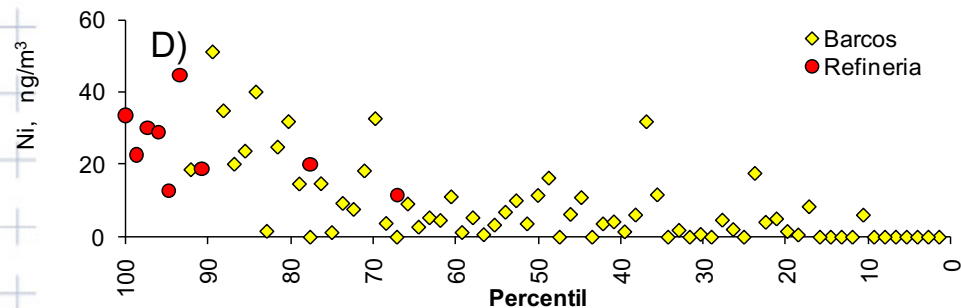
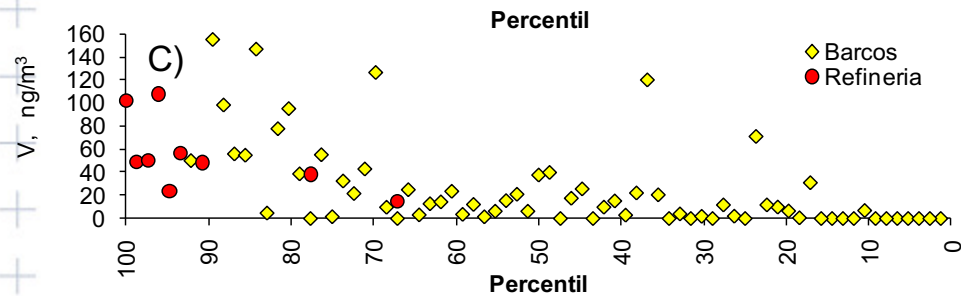
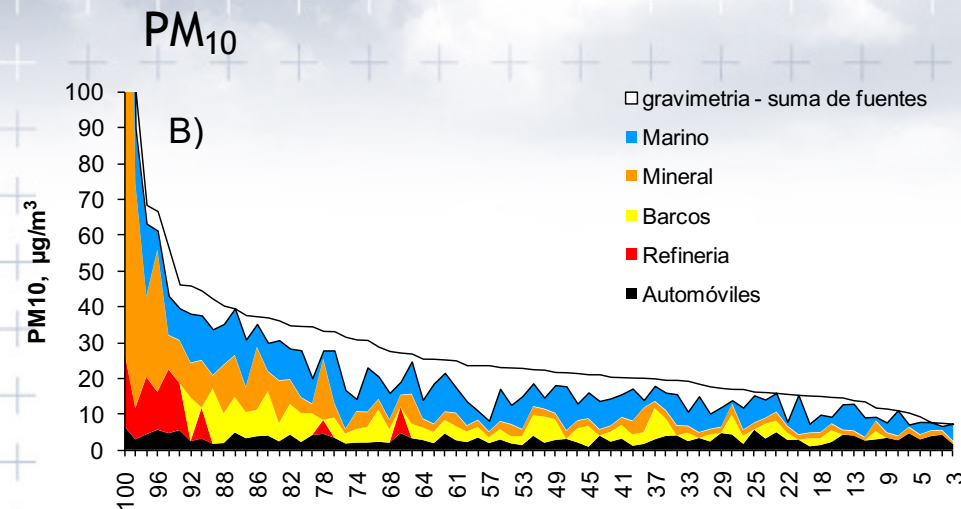


$$PM_{10} > 50 \mu\text{g}/\text{m}^3$$

Tabla 11. Episodios en los que se supera el valor límite diario para partículas PM10 ($50 \mu\text{g}/\text{m}^3$). A) Contribución de fuentes. B) Composición de cada fuente antropogénica de PM10. C) Información complementaria. En A) % indica la contribución a la media de cada fuente a la concentración de PM10; Natural = Mineral + Marino; Antropogénico = Barcos + Refinería + Automóviles. En C) unidades: V ($\mu\text{g}/\text{m}^3$), Ni ($\mu\text{g}/\text{m}^3$), SO_2 ($\mu\text{g}/\text{m}^3$) y $\text{SO}_4^{=}$ ($\mu\text{g}/\text{m}^3$).

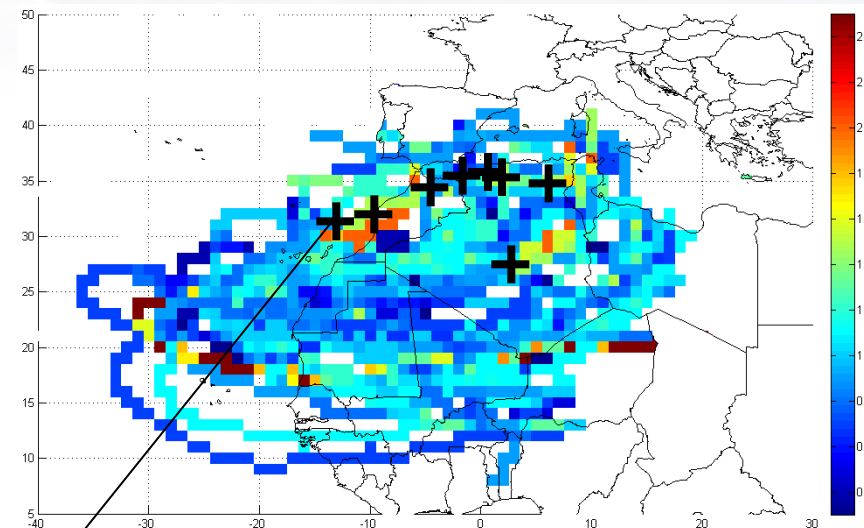
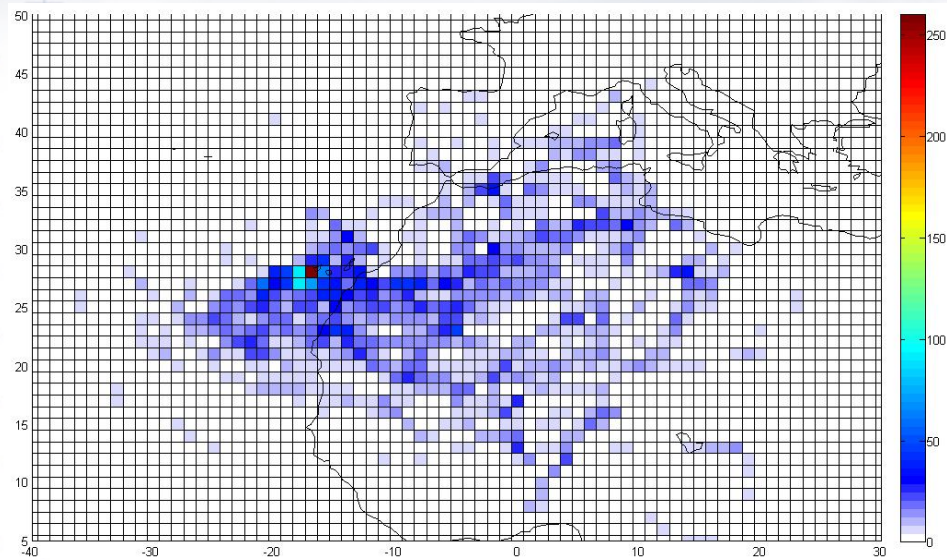
| A) | 26/04/2008 | | 21/01/2008 | | 02/04/2008 | | 11/11/2008 | | 29/06/2008 | |
|------------------|--------------------------|----|--------------------------|----|--------------------------|----|--------------------------|----|--------------------------|----|
| | $\mu\text{g}/\text{m}^3$ | % | $\mu\text{g}/\text{m}^3$ | % | $\mu\text{g}/\text{m}^3$ | % | $\mu\text{g}/\text{m}^3$ | % | $\mu\text{g}/\text{m}^3$ | % |
| PM10 | 376.4 | | 100.6 | | 68.5 | | 66.8 | | 56.6 | |
| Σ contribuciones | 319.1 | 85 | 89.9 | 89 | 63.3 | 92 | 61.3 | 92 | 43.1 | 76 |
| Mineral | 270.4 | 72 | 62.9 | 63 | 22.3 | 33 | 39.6 | 59 | 9.4 | 17 |
| Marino | 21.0 | 6 | 15.3 | 15 | 20.5 | 30 | 5.4 | 8 | 11.1 | 20 |
| Barcos | 0.0 | 0 | 0.0 | 0 | 0.0 | 0 | 0.0 | 0 | 0.0 | 0 |
| Refinería | 20.9 | 6 | 8.8 | 9 | 16.1 | 23 | 10.6 | 16 | 17.9 | 32 |
| Automóviles | 6.8 | 2 | 3.0 | 3 | 4.4 | 6 | 5.7 | 9 | 4.8 | 8 |
| Mineral + marino | 291.4 | | 78.1 | | 42.8 | | 45.0 | | 20.5 | |
| Antropogénico | 27.7 | | 11.7 | | 20.5 | | 16.3 | | 22.6 | |

Rodríguez et al., 2009



Santa Cruz de Tenerife source apportionment study by receptor modeling
Rodríguez et al., 2009

PM₁₀>10 μ g/m³

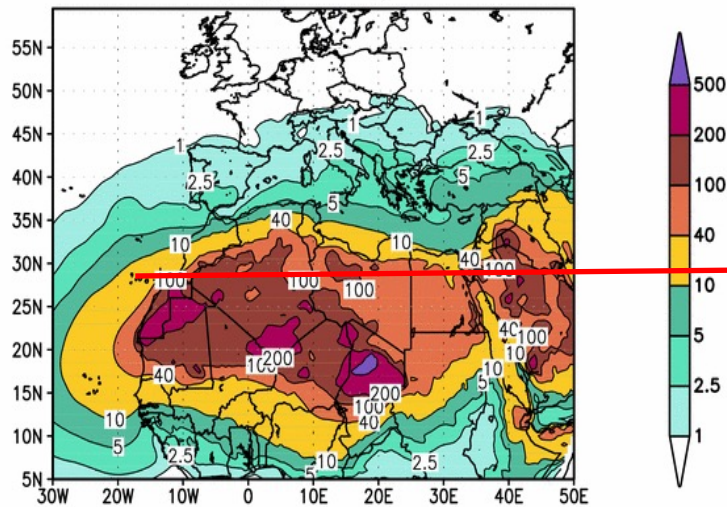


SAMIR (33.68°N, -7.43°E)
Refinery (125,000 bpd)

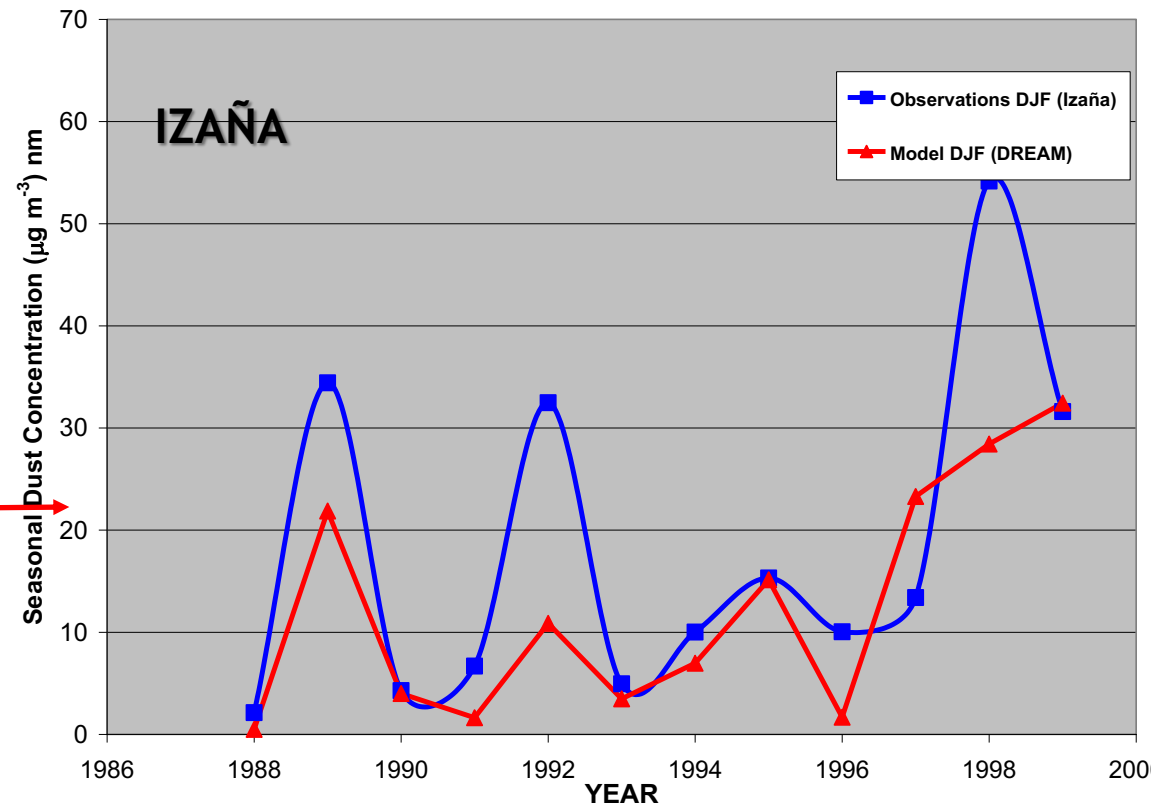


BSC-DREAM validation against TSP at Izaña

Surf. dust conc. [$\mu\text{g}/\text{m}^3$] DJF 1981–2006



Winter



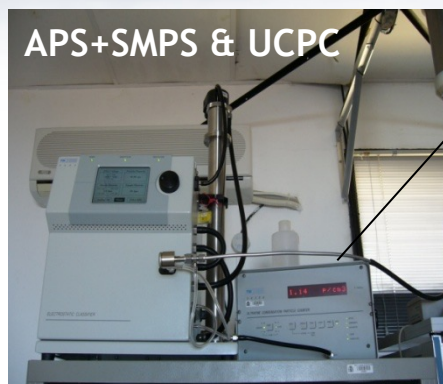
$R=0.79$

In-situ dust concentration validation: Izaña GAW station-DREAM

Pérez et al., 2008

In-situ Aerosol/ African-dust characterization

Number size distribution of coarse, fine and ultra-fine particles (3nm-20 μ m) (APS + SMPS)



Ultrafine Condensation Particle Counter (UCPC): Counts number concentration (up to 10^5 cm^{-3}) of particles of the size 3 nm and higher

Scanning Mobility Particle Sizer (SMPS): Measures particle size distribution in 15-770 nm mobility diameter range. (resolution 64 chan/decade)

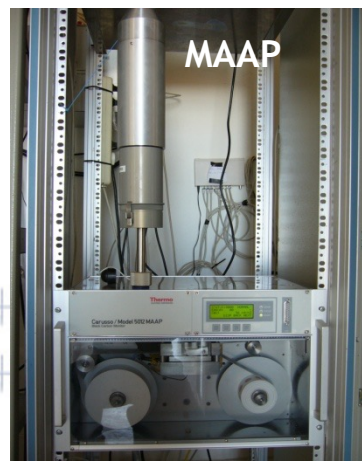
Total and backward scattering coefficients (3 λ) (integrating nephelometer)



TSI 3-wavelength Integrating Nephelometer

Measures total- and back-scattering coefficients at 3 wavelengths: 700 nm (red), 550 nm (green), and 450 nm (blue)

Absorption scattering coefficients (1 λ) (MAAP)



MAAP or Aethalometer

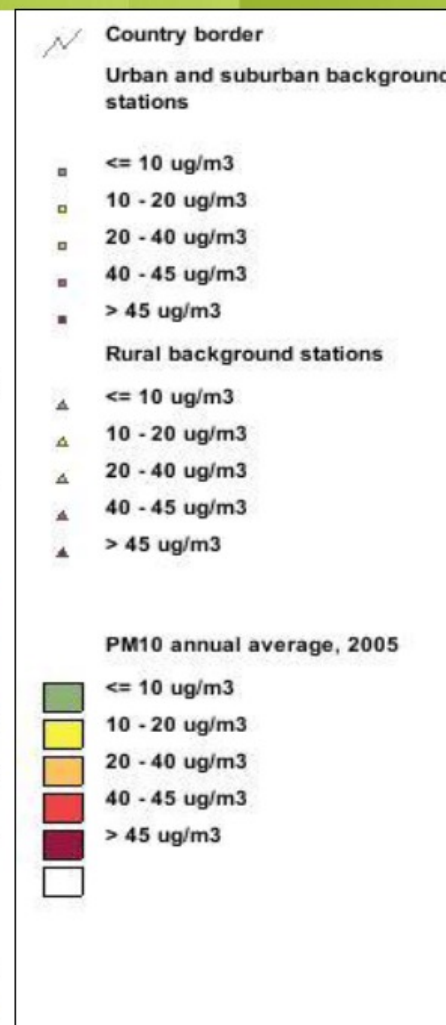
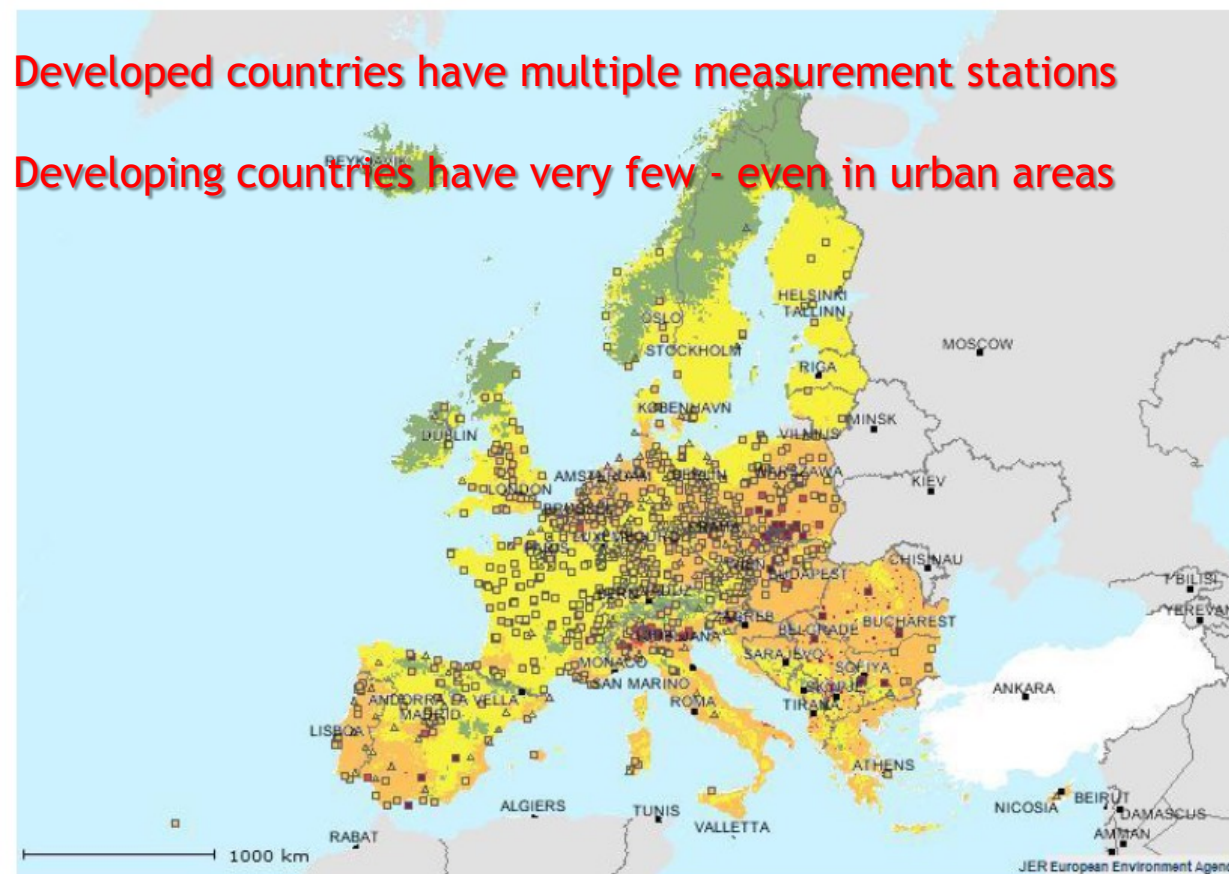
Measures mass loading of black carbon ($\mu\text{g}/\text{m}^3$)

PM10 annual average, 2005

Many different PM instruments and networks

Developed countries have multiple measurement stations

Developing countries have very few - even in urban areas



- ~ 100 PM2.5 & PM10 stations in Europe in 2003 (~ 1600 PM10-only stations). Mostly daily (24h) average data are available.
- Measurement principles: TEOM, beta-attenuation, gravimetry

From ground observations...
to ground estimations...

Visibility

WMO - visibility

The greatest distance that a black object of “suitable dimensions,” situated near the ground, can be seen and recognized when observed against a background of fog

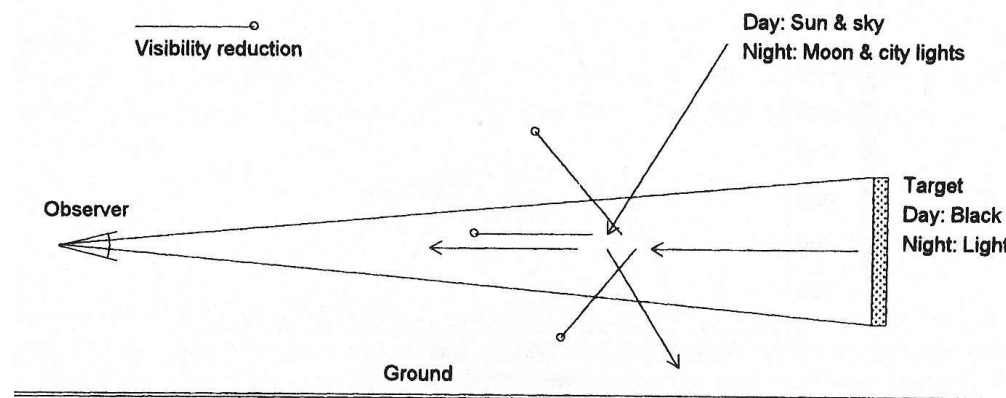


Fig. 11-1 Visibility reduction by scattering.

- Operational surface synoptic weather station reports from Global Telecommunication System (GTS)
- Station reports include past & present weather, visibility (km), temperature (°C), dew point temperature, wind direction (°), and speed (knots)

| | | | | | | | | | | | |
|-------|-------|-------|----------|------------------------------|----|----|-----|----|-----|-----|--------|
| 62733 | 15.32 | 35.60 | 02040818 | Dust, not at time of obs. | 6 | 0 | 18 | 22 | 320 | 2 | 35.5 |
| 62733 | 15.32 | 35.60 | 02041015 | Dust, raised at time of obs. | 7 | 0 | 99. | 30 | 320 | 6 | 34.5 |
| 62733 | 15.32 | 35.60 | 02041121 | | -9 | -9 | -9 | 20 | 23 | 320 | 2 26.0 |
| 62733 | 15.32 | 35.60 | 02041212 | | -9 | -9 | -9 | 20 | 34 | 340 | 3 37.5 |

Measurement of visibility - transmissometer

- A light source with one or two light detectors at fixed distances from the source
- Detectors are designed to receive light only from the source direction
- Often located along and parallel to a runway (runway visual range; RVR)

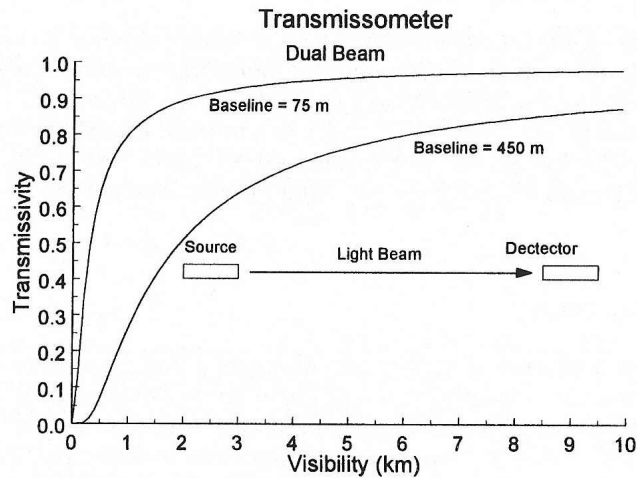


Fig. 11-3 Transfer function for a transmissometer.

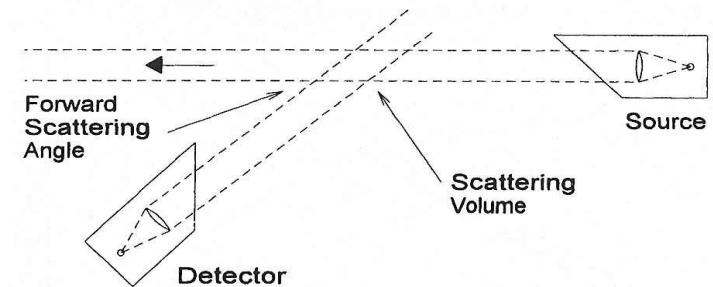
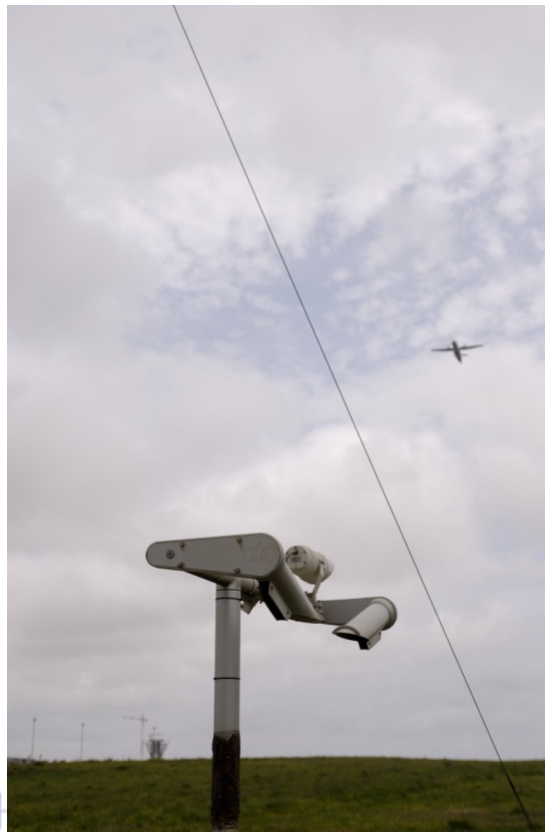
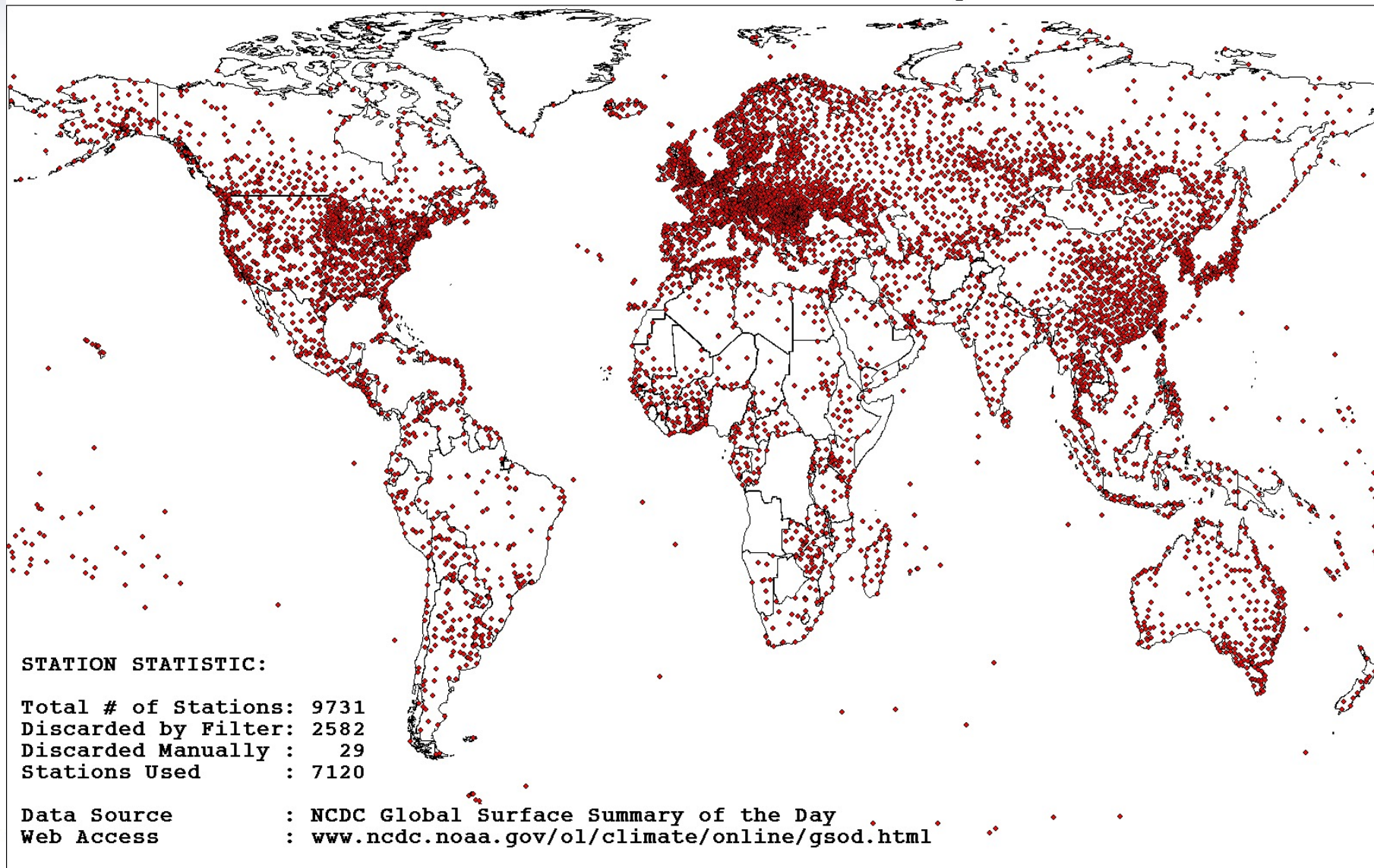


Fig. 11-4 A forward scatter visibility meter.

WMO- World Wide Watch Global Surface Meteorological Network



Relation between horizontal visibility and TSP or PM10

Very few studies on the relation between horizontal visibility and TSP or PM10 levels of mineral dust mass concentration have been carried out in Africa.

$$C_{TSP} = 1339.84 VV^{-0.67}$$

Ben Mohamed et al. (1992)

$$C_{PM10} = 914.06 VV^{-0.73} + 19.03$$

D'Almeida's (1986)

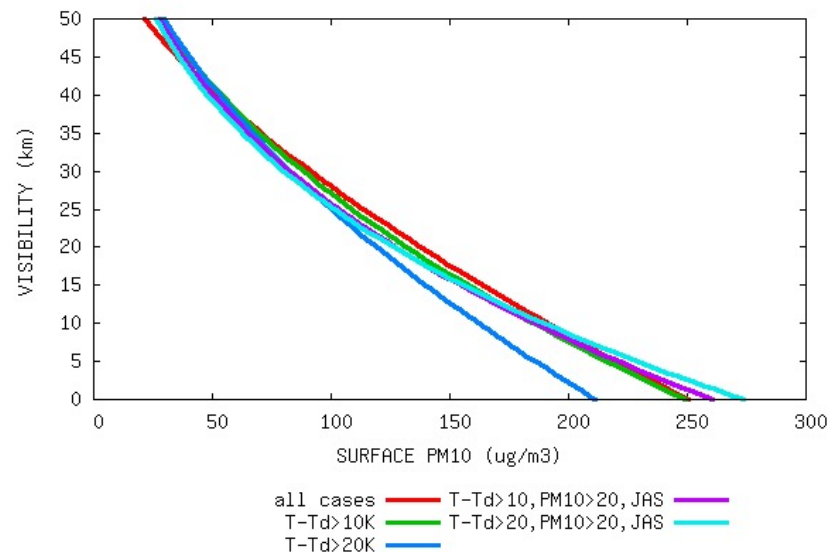
where C is the TSP concentration in μgm^{-3} and VV is the horizontal visibility in km

E. Terradellas

$$\text{Vis} = 63023 - 1838(\text{PM}_{10}^{0.64})$$

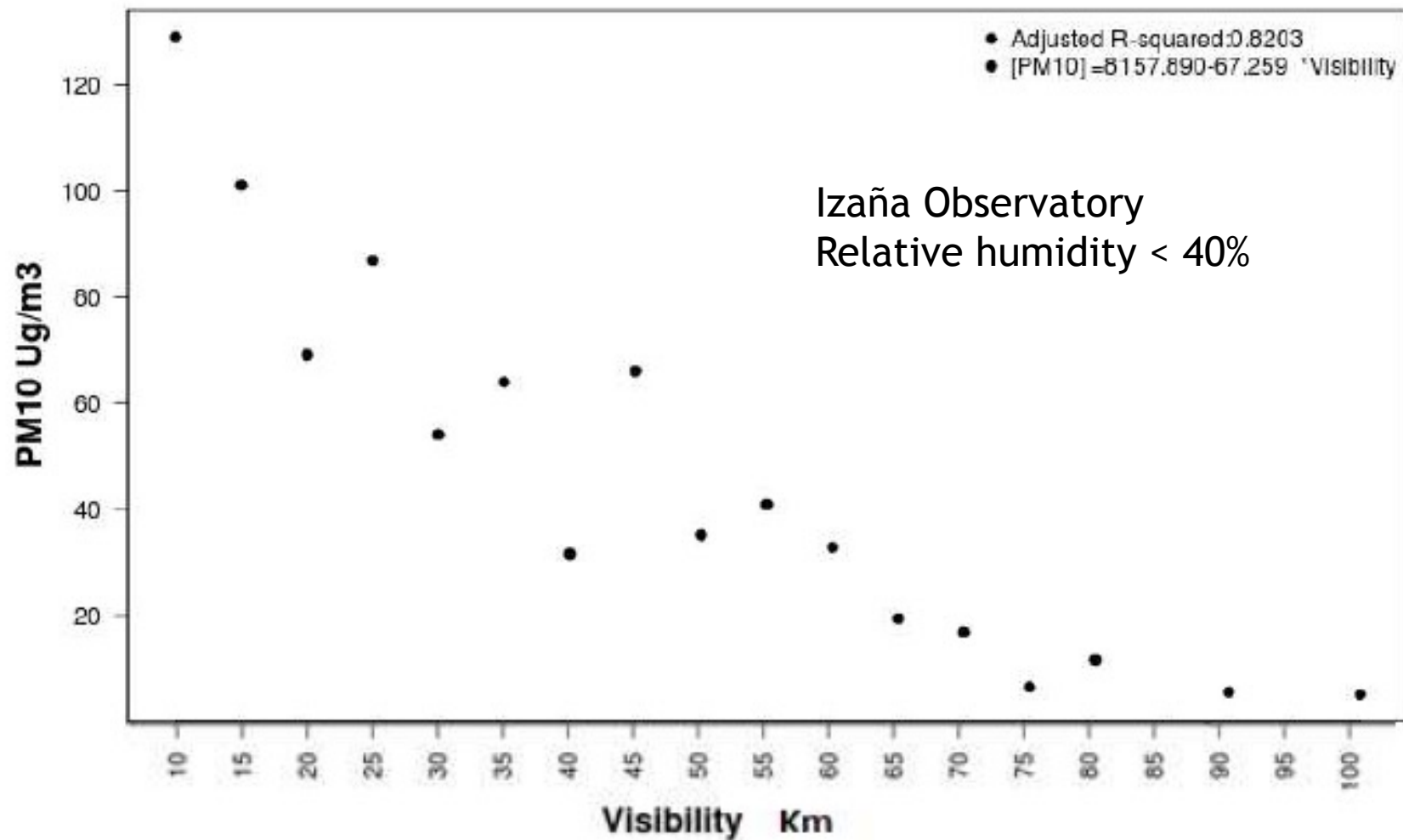
PM_{10} at Izaña

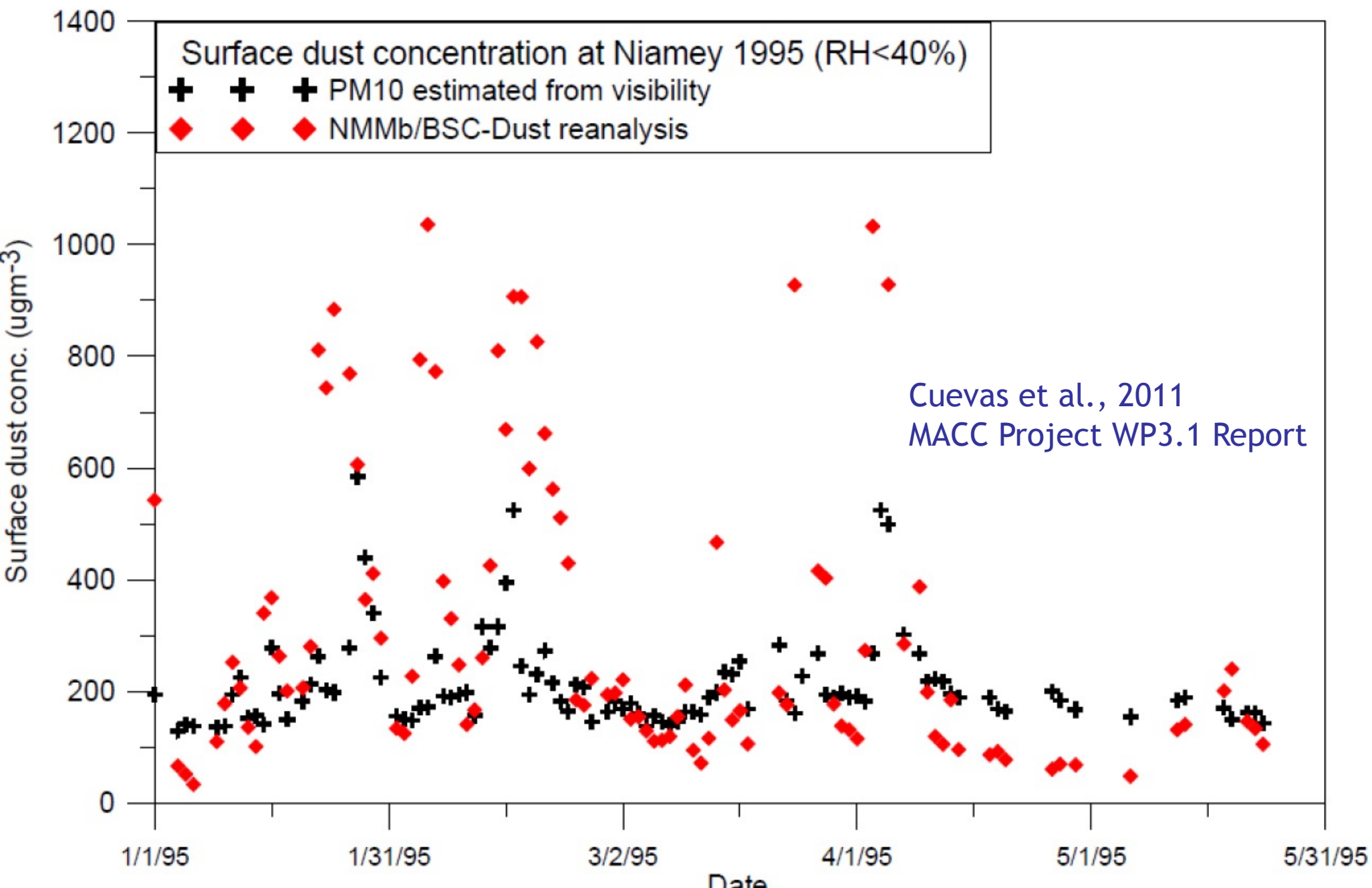
Visibility from SYNOP-Izaña



Identify surface station visibility reports that may be used in simple regression model for estimate ground PM10 or TSP

Scatterplot at Izaña

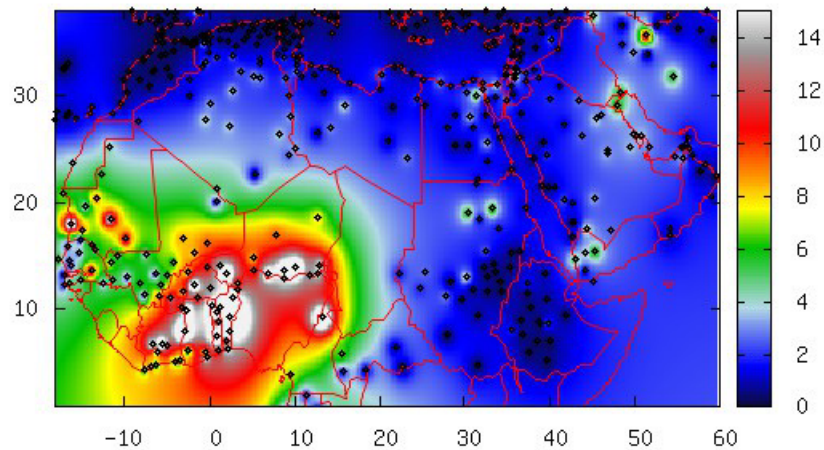




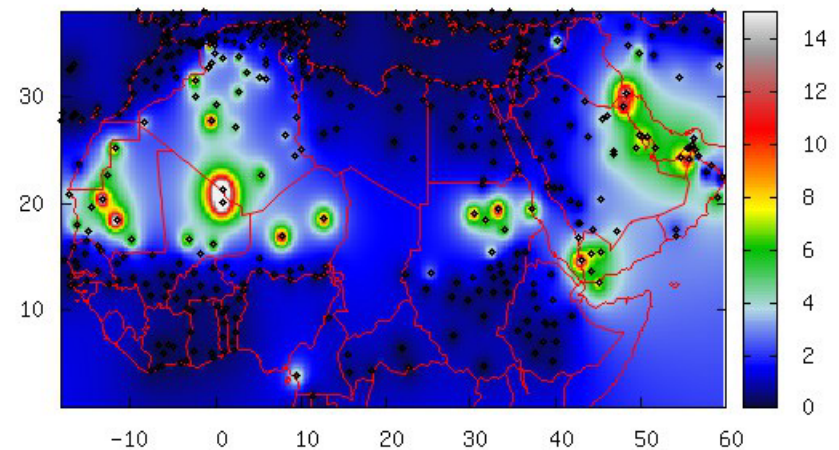
Synop and Metar reports from Met stations

Provided by Enric Terradellas

WINTER



SUMMER



Problems with station visibility estimates

1. Human observations are inherently subjective.
2. No all reductions of visibility are due to dust (fog, biomass burning...)
3. Coarse reporting bins
4. Judgment in distinguishing visibility beyond 10 km

Main advantages

1. Reports are abundant and widespread over land. There is information in remote areas (deserts)
2. There are *some* standards
3. Human detected visibility has been correlated well with surface extinction analyses (Husar et al., 2000)
4. Estimations of PM are possible

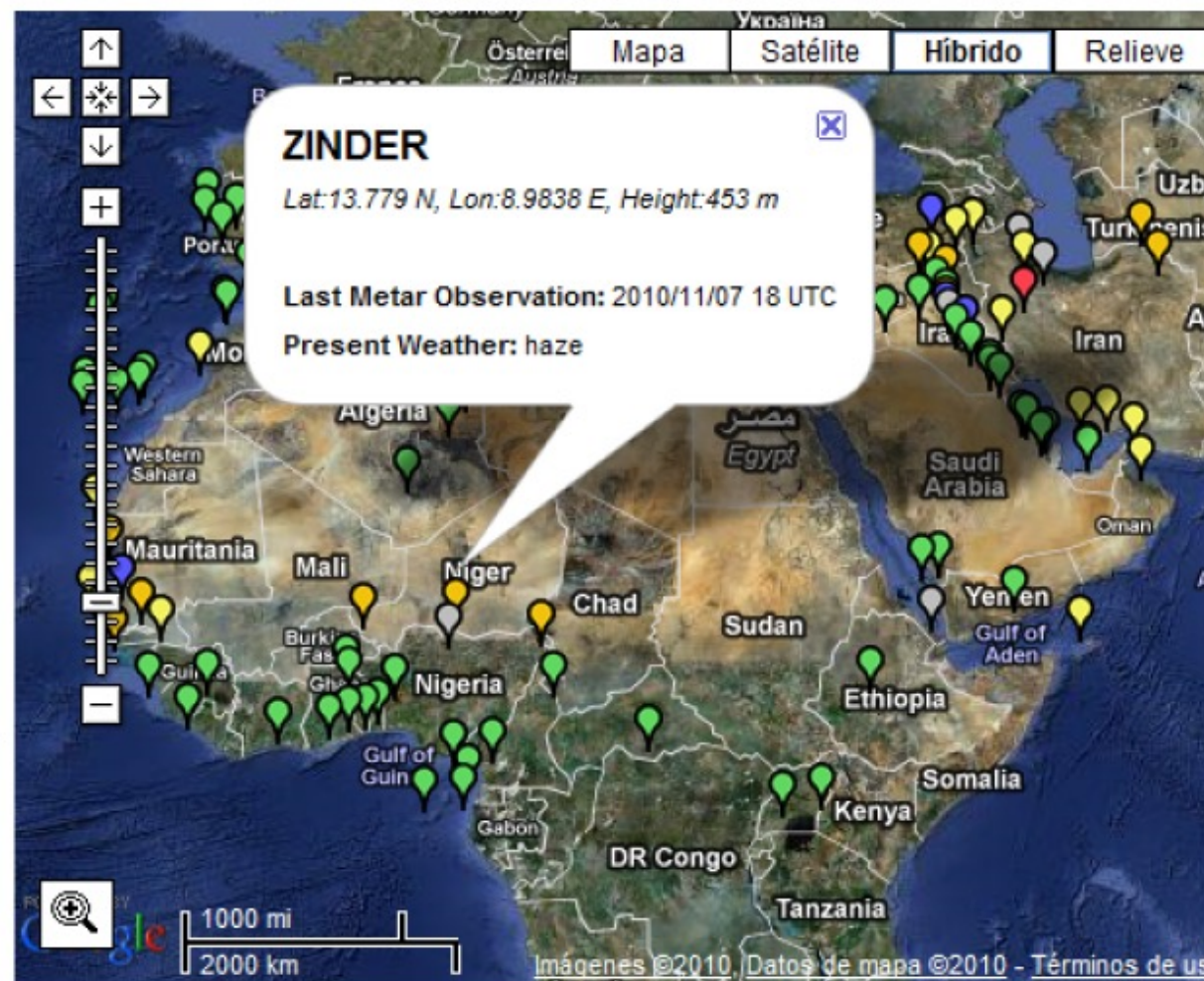
| AERONET Verification | MSG Verification | MODIS Real-Time | Visibility Real-Time |
|---|------------------|-----------------|----------------------|
| http://www.bsc.es/sds-was/node/46 | | | |

-  < 50 m
-  > 50 m And < 2 km
-  > 2 km And < 4 km
-  > 4 km And < 6 km
-  > 6 km And < 8 km
-  > 8 km And < 10 km
-  > 10 km

Last Update:
 2010-11-07 21:37:07

CLICK ON A STATION FOR TIME OF OBSERVATION

Develop a quality-screened station dataset for use in validating dust model forecasts



From ground observations...
to total atmospheric column observations

Sunphotometers

Aerosol Extinction Profiles: A measure of attenuation of the light passing through the atmosphere due to scattering and absorption by aerosol particles.

Extinction coefficient is the fractional depletion of radiance per unit path length (also called attenuation). It has units of km^{-1} .

Aerosol Mass Concentration: The columnar aerosol mass concentration ($\mu\text{g}/\text{m}^3$) is the total aerosol mass in a vertical column of atmosphere.

Aerosol Single Scattering Albedo A measure of the effectiveness of scattering relative to extinction for the light encountering the atmospheric aerosol particles.

The scattering and absorption of radiation by a single aerosol particle is expressed by its **complex refractive index** ($n = n_r + i n_i$), where the real part represents scattering and the imaginary part represents absorption. The refractive index is strongly dependent on the chemical composition of the particle.

Aerosol Asymmetry Factor A measure of the preferred scattering direction (forward or backward) for light encountering aerosol particles.

$$g = \frac{1}{2} \int_{-1}^{+1} \cos \Theta P(\cos \Theta) d \cos \Theta$$
$$P(\cos \Theta) = \frac{1 - g^2}{(1 + g^2 - 2g \cos \Theta)^{3/2}}$$

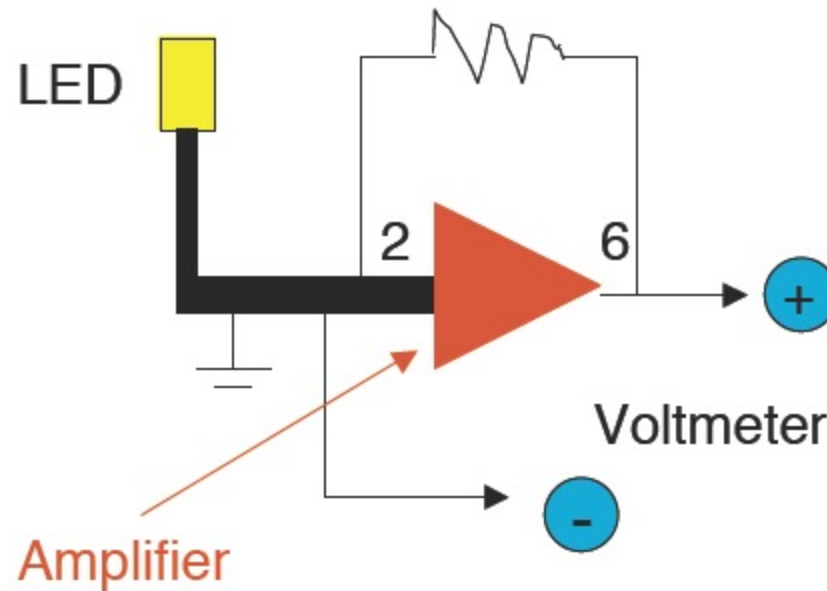
The asymmetry factor approaches +1 for scattering strongly peaked in the forward direction and -1 for scattering strongly peaked in the backward direction.

In general, $g=0$ indicates scattering directions evenly distributed between forward and backward directions, i.e. isotropic scattering (e.g. scattering from small particles)

$g < 0$ scattering in the backward direction (i.e scattering angle > 90 deg.), often referred to as backscattering, is scattering at 180 deg.

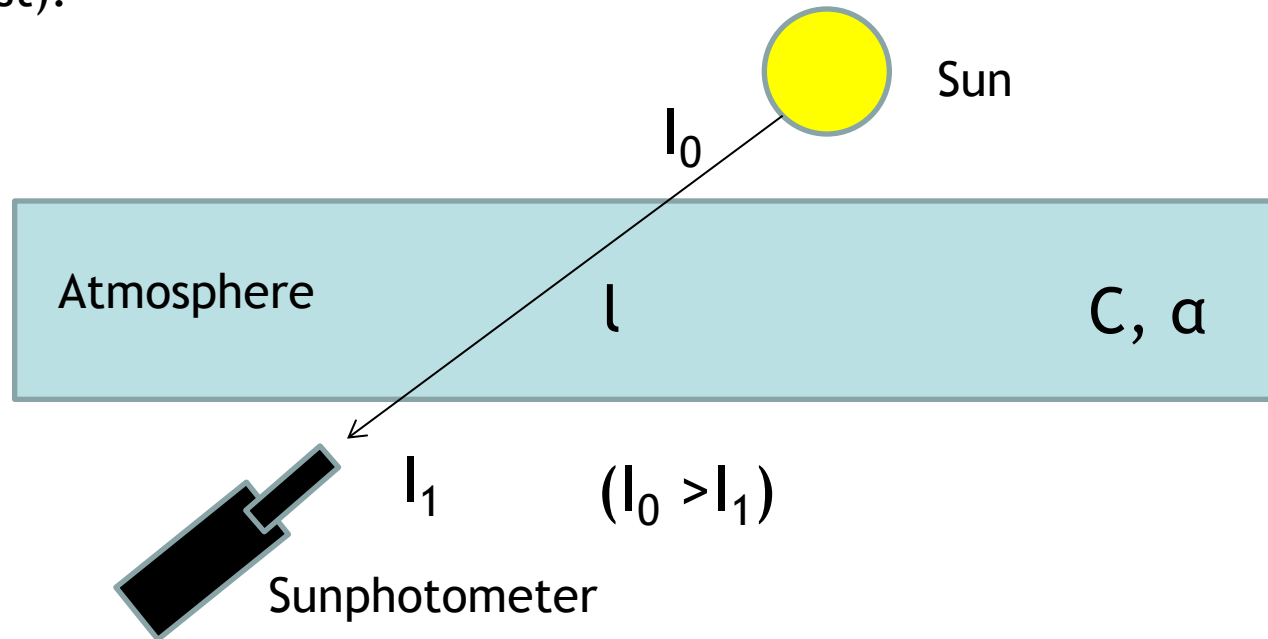
$g > 0$ scattering in the forward direction (i.e scattering angle < 90 deg.), often referred to as forward-scattering, is scattering at 0 deg. **For larger size or Mie particles, g is close to +1.**

Sun Photometers absorb *direct* sunlight energy with a LED light and convert the intensity into a quantified voltage to measure aerosols in the atmosphere.



The intensity of sunlight at the top of the earth's atmosphere is constant. While the sunlight travels through the atmosphere, aerosols can dissipate the energy by scattering (Rayleigh and Mie) and absorbing the light. More aerosols in the atmosphere cause more scattering and less energy transmitted to the surface.

Knowing the sunlight's energy at the top of the atmosphere, the thickness of the atmosphere, and the amount of sunlight transmitted to the earth's surface and can allows us to determine the amount of scattering, and thus, the amount of aerosols (dust).



Beer's Law

$$T = I_1/I_0 = 10^{-\alpha l} = 10^{-\epsilon l c}$$

Transmissivity (T)

Absorption coefficient (α)

path length (l)

molar absorptivity of the absorber (ϵ)

concentration of absorbing species in the material (c)

absorption cross section (σ)

density of absorbers (N)

Langley plot calibration (100 determination for each wavelength):

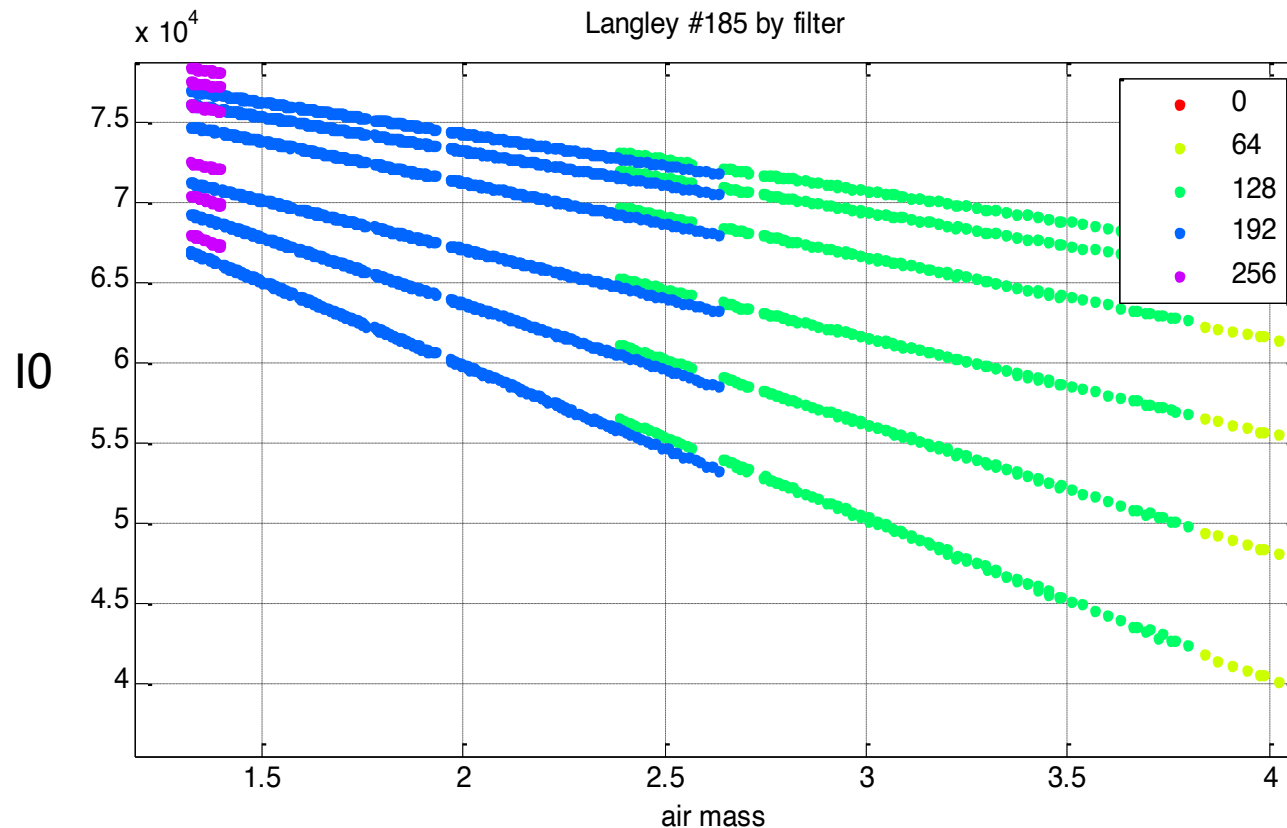
Pristine conditions

No clouds

Stable total ozone and column water vapor

Beer's Law

$$T = I_1/I_0 = 10^{-\alpha l} = 10^{-\epsilon l C}$$



Aerosol Optical Depth (or Thickness)

"Aerosol Optical Depth" (AOD) is the degree to which aerosols prevent the transmission of light. The aerosol optical depth or optical thickness (τ) is defined as the integrated extinction coefficient over a vertical column of unit cross section.

$$\tau_{ext} = \int_{z=0}^{z=toa} K_{ext}(z) dz$$

Angstrom Exponent

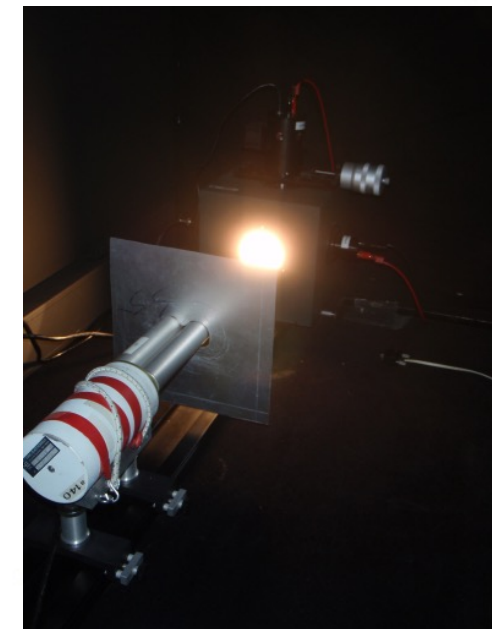
An exponent that expresses the spectral dependence of Aerosol Optical Depth (τ) with the wavelength of incident light (λ). The spectral dependence of aerosol optical thickness can be approximated (depending on size distribution) by:

$$\tau_a = \beta \lambda^\alpha$$

where α is the Angstrom exponent (β = aerosol optical depth at 1 μm)

i.e. If AOD >~ 0.2 and $\alpha < 0.7$ then we are observing dust (aprox.)

- The Cimel Electronique 318 spectral radiometer is a solar-powered, weather-hardy, robotically-pointed sun and sky spectral sun photometer.
- A sensor head points the sensor head at the sun according to a preprogrammed routine.
- The Cimel controller, batteries, and the optional Vitel satellite transmission equipment are usually deployed in a weatherproof plastic case.



Direct irradiance
IFOV 1.2°

Principal plane

Diffuse radiance
IFOV 1.2°

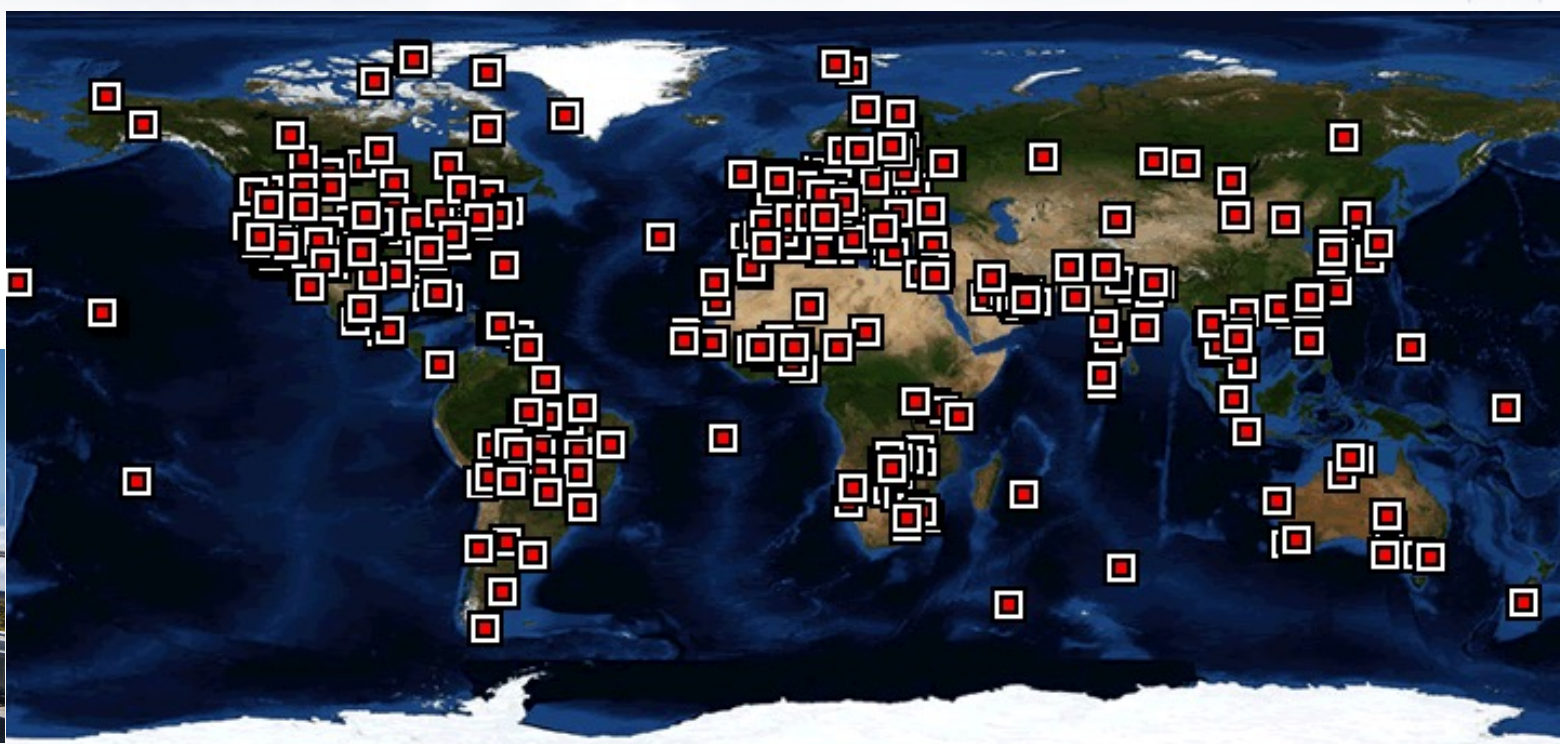
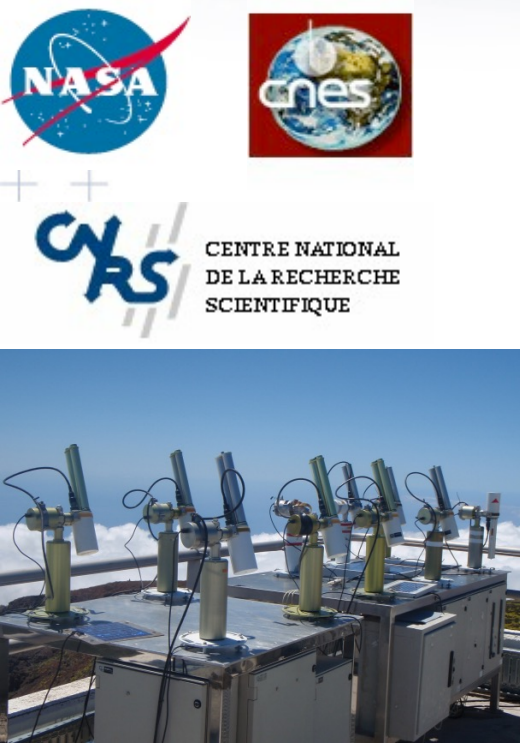
Almucantar

Photometer

Sun measurements
Sky measurements

AERONET (AERosol RObotic NETwork)-

<http://aeronet.gsfc.nasa.gov>



An internationally Federated Network

- Characterization of aerosol optical properties
- Validation of satellite aerosol retrieval
- Near real-time acquisition; long term measurements

AERONET Data Flows

<http://aeronet.gsfc.nasa.gov>

Flux measurements

Direct - $\lambda=340, 380, 440, 500, 670, 870, 940, 1020$ nm

Diffuse - $\lambda=440, 670, 870, 1020$ nm (alm, pp, pol)

Calibration and processing information

Mauna-Loa and Izaña

CNRS-University of Lille and University of Valladolid

Aerosol optical depth and precipitable water computations

Cloud screening and quality control

Inversion products

Volume size distribution ($0.05 < R < 15$ μm),
refractive index, single scattering albedo
($\lambda=440, 670, 870, 1020$ nm)

Holben et al.
RSE, 1998
Holben et al.
JGR, 2001

Eck et al.
JGR, 1999

Smirnov et al.
RSE, 2000

Dubovik and King
JGR, 2000
Dubovik et al.
JGR, 2000
GRL, 2002

AERONET provides:

- global Aerosol Optical Depth of Dust in near real-time
- robust optical properties of Dust: size distribution, ref. Index, etc. (e.g. Asian Dust has stronger and less spectral dependent absorption than Saharan Dust)
- climatological models that reproduce observed optical properties of aerosol (useful for satellite retrievals)

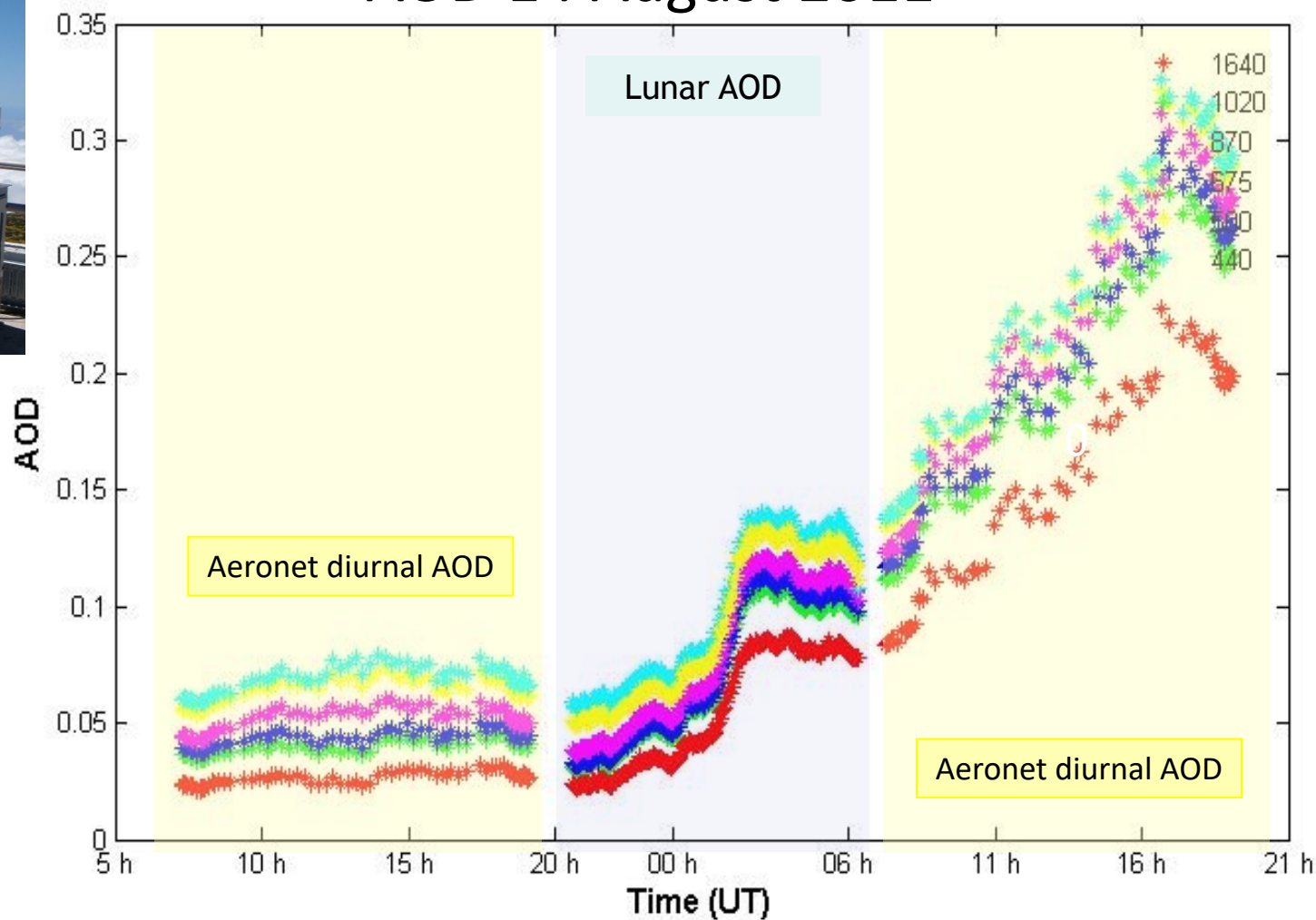


AERONET @ Tamanrasset-Algeria
September 2006
Mohamed Mimouni



AERONET @ Cairo
May 2010
Hamza Mohamed Hamza

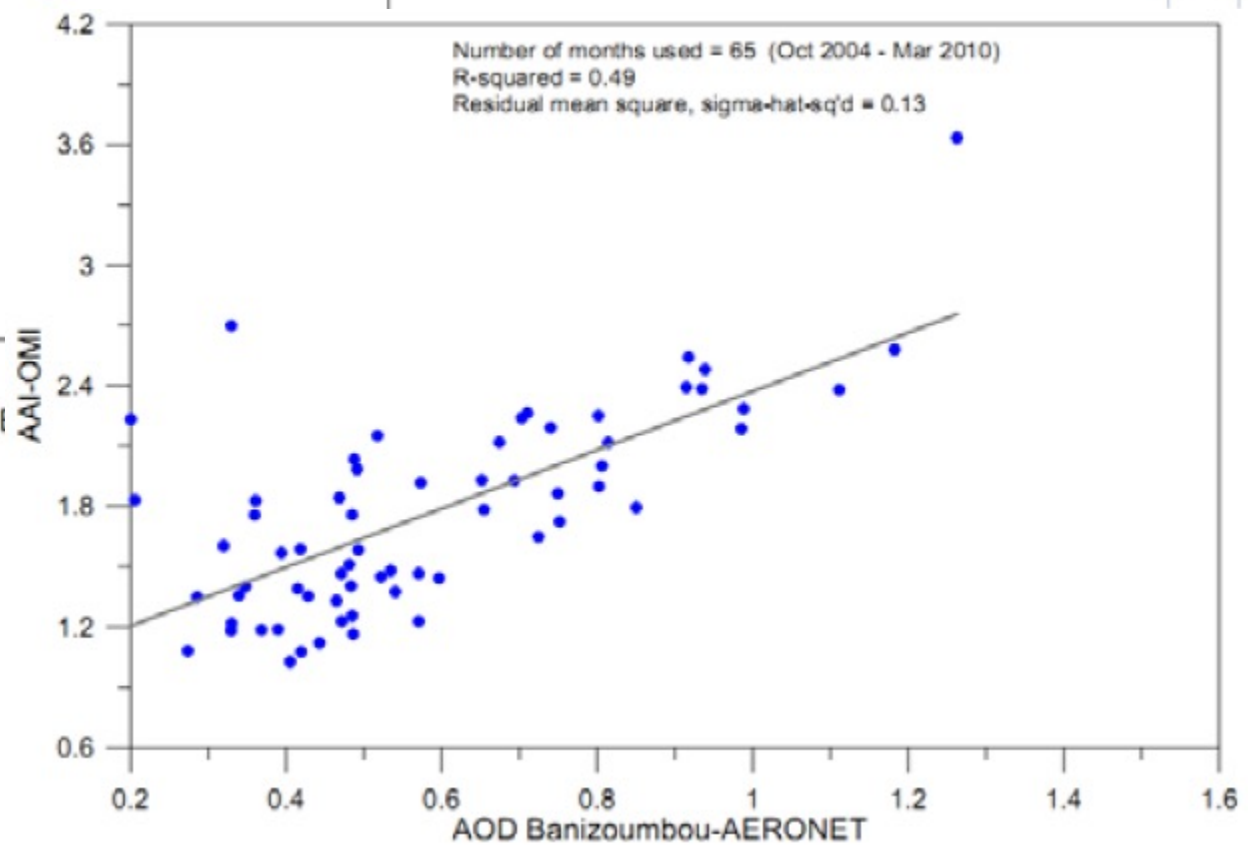
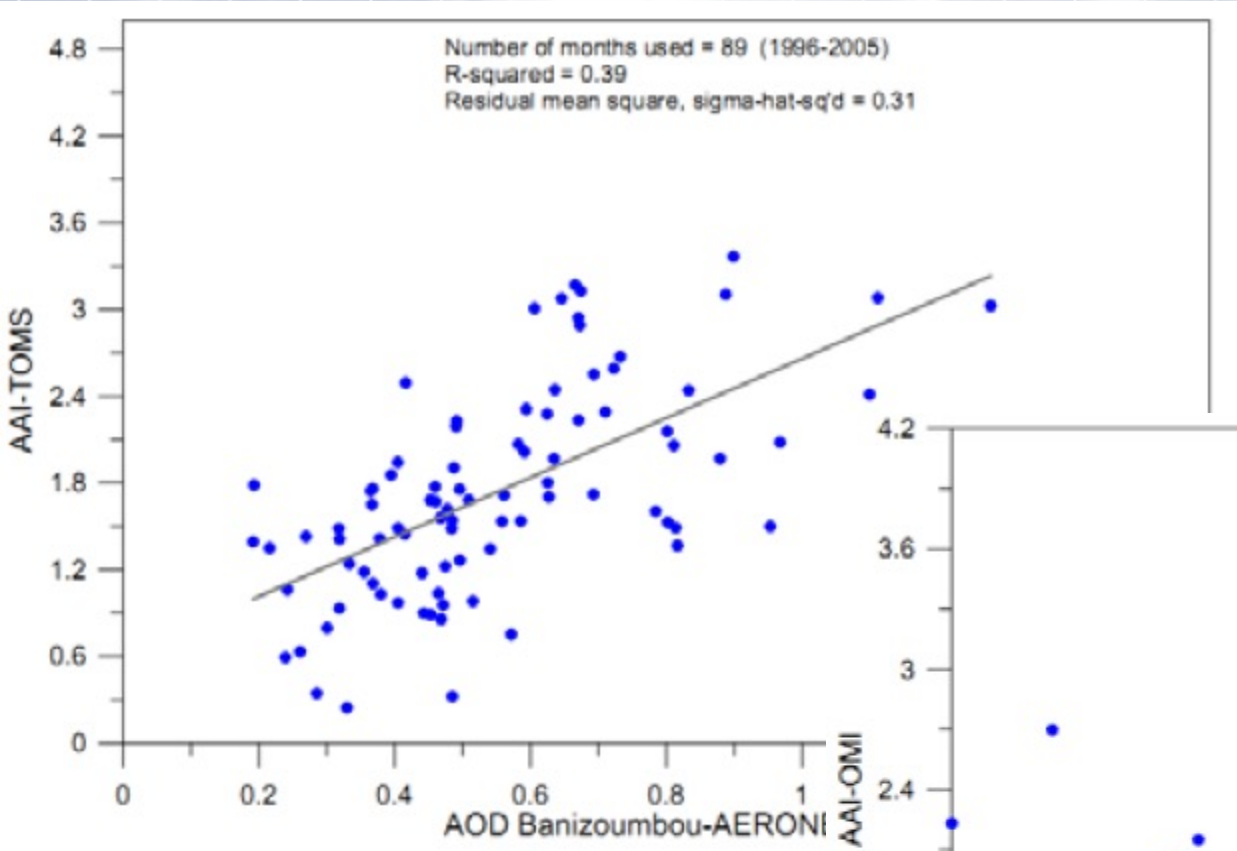
AOD 14 August 2011

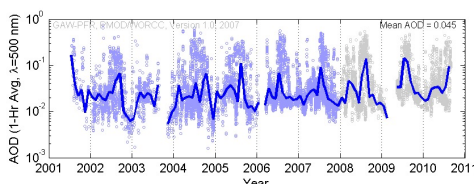
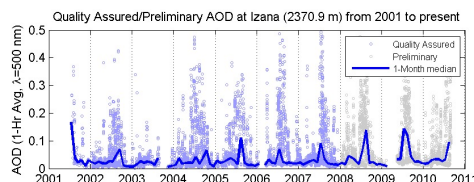
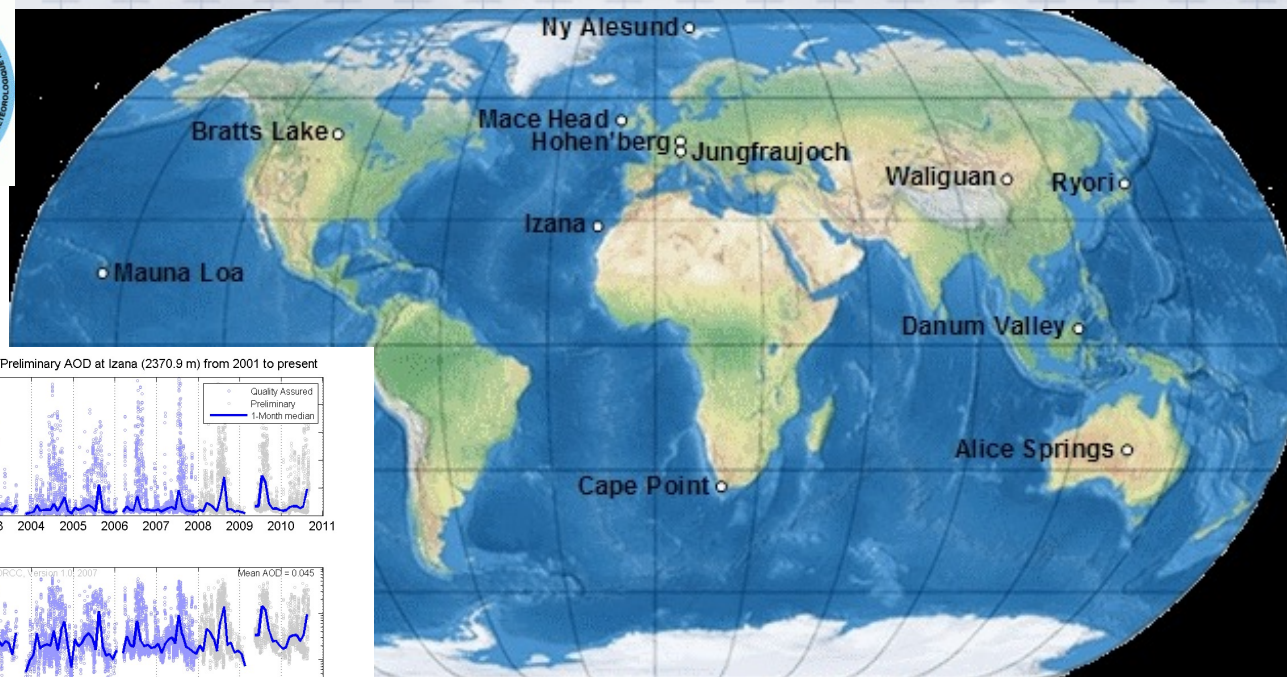


August 13

August 14







- Classic extinction measurements at the recommended 4 WMO wavelengths 368, 415, 500 and 862 nm using Precision Filter Radiometers (PFRs).
- Continuous sampling at a 1- minute frequency by automated systems.
- Data products: AOD and the Angström coefficients alpha and beta (no inversions).
- Hourly mean AOD archived at the [World Data Center for Aerosols \(WDCA\)](https://www.wdc.aero/), Ispra. Data with a 1-minute resolution are available from WORCC upon request.

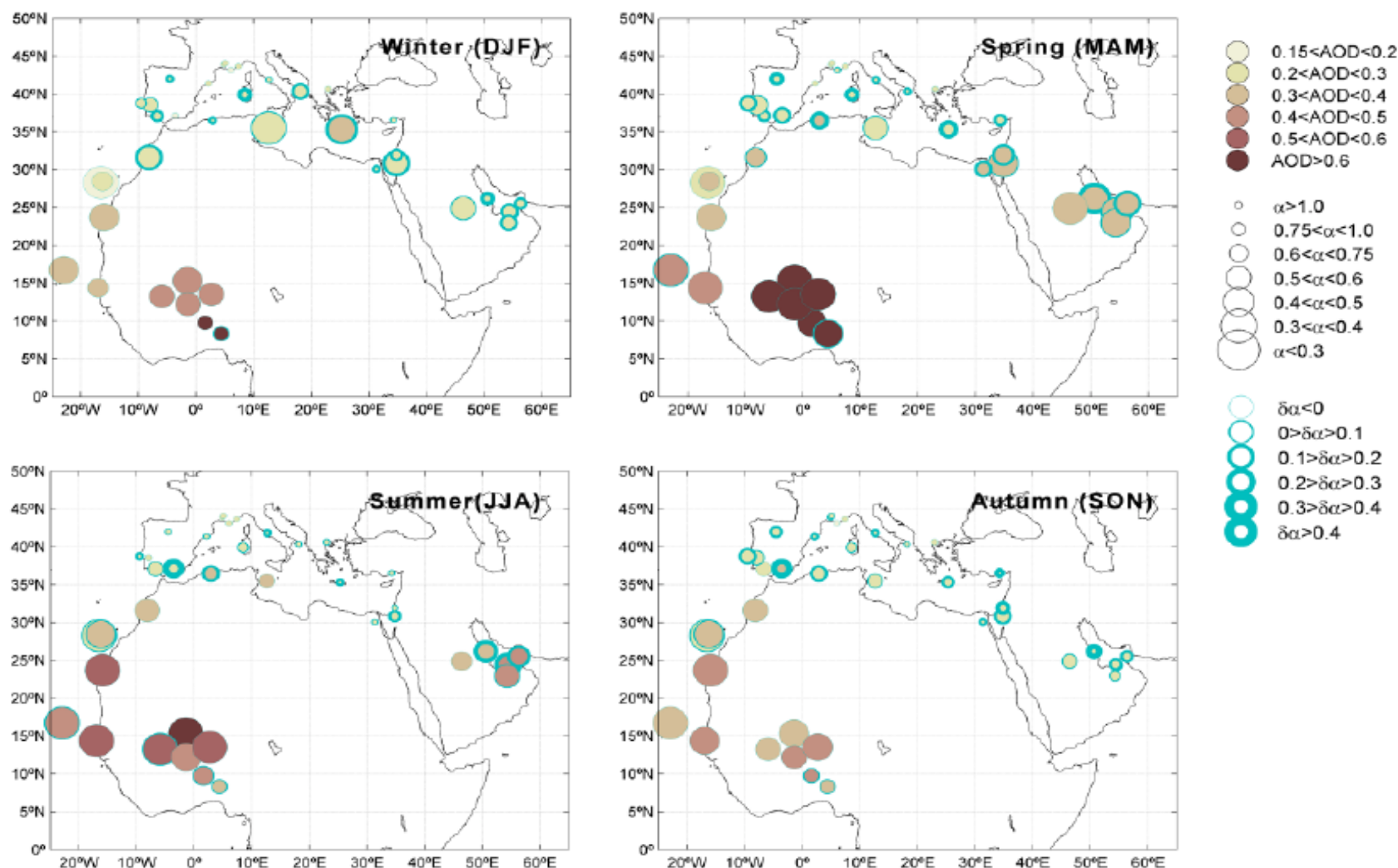
GAW-PFR provides:

- long-term high-accuracy AOD and Angström Coefficients
- GAW-PFR provides AOD Dust in near real-time

Aerosol characterization in Northern Africa, Northeastern Atlantic, Mediterranean Basin and Middle East from direct-sun AERONET observations

S.Basart, C. Pérez, E. Cuevas, J.M. Baldasano, and G.P. Gobbi

(Atmos. Chem. Phys. October-2009)



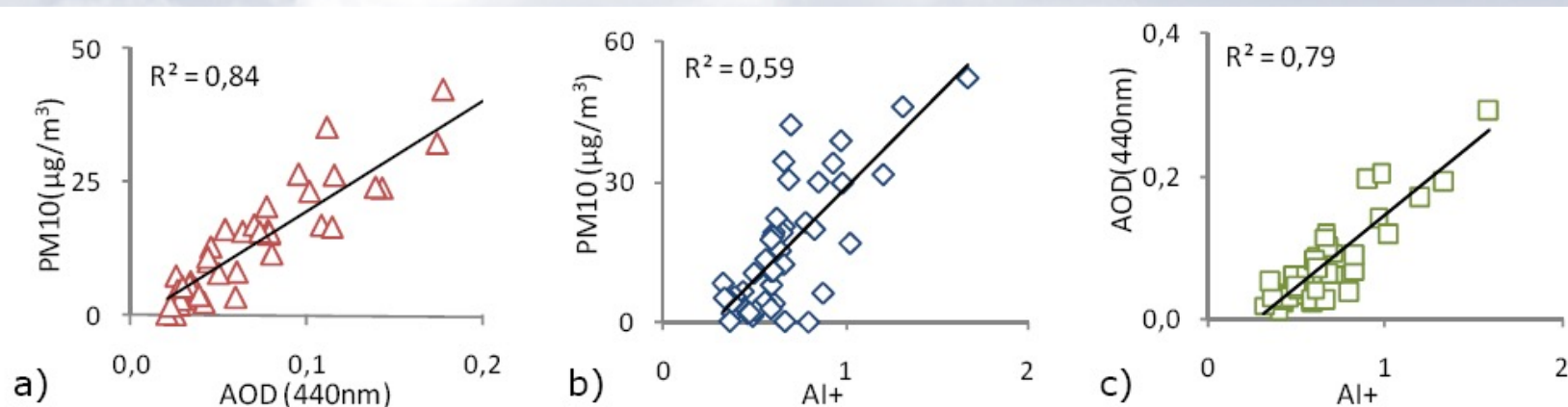


Figure 1. Scatterplot of monthly means of a) AOD vs PM₁₀; b) AI positive values vs PM₁₀; c) AI positive values vs AOD.

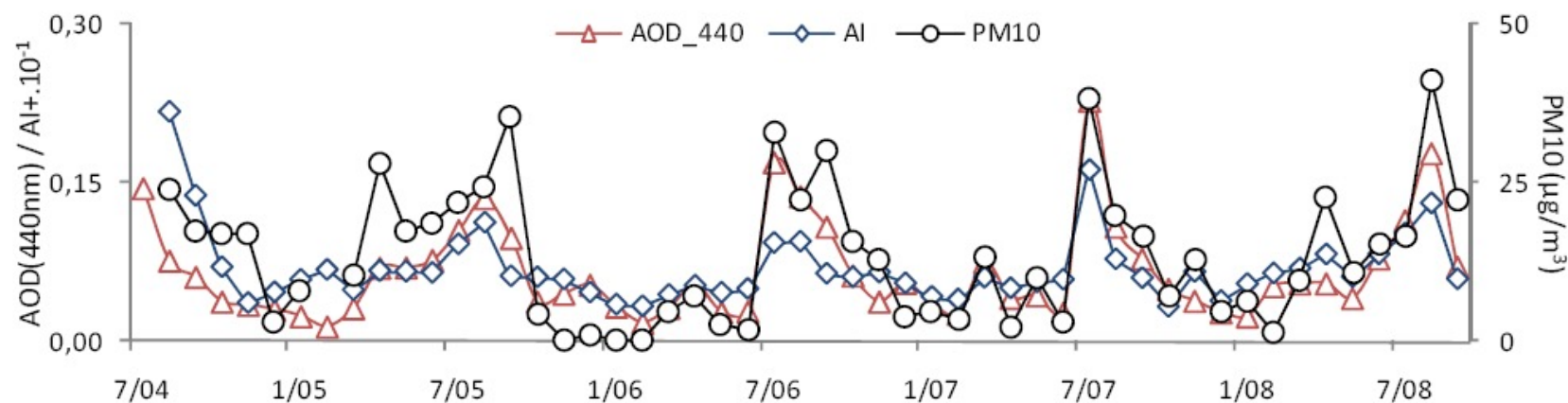


Figure 2. Monthly means of PM₁₀ (µg/m³), AOD and AI positive values.

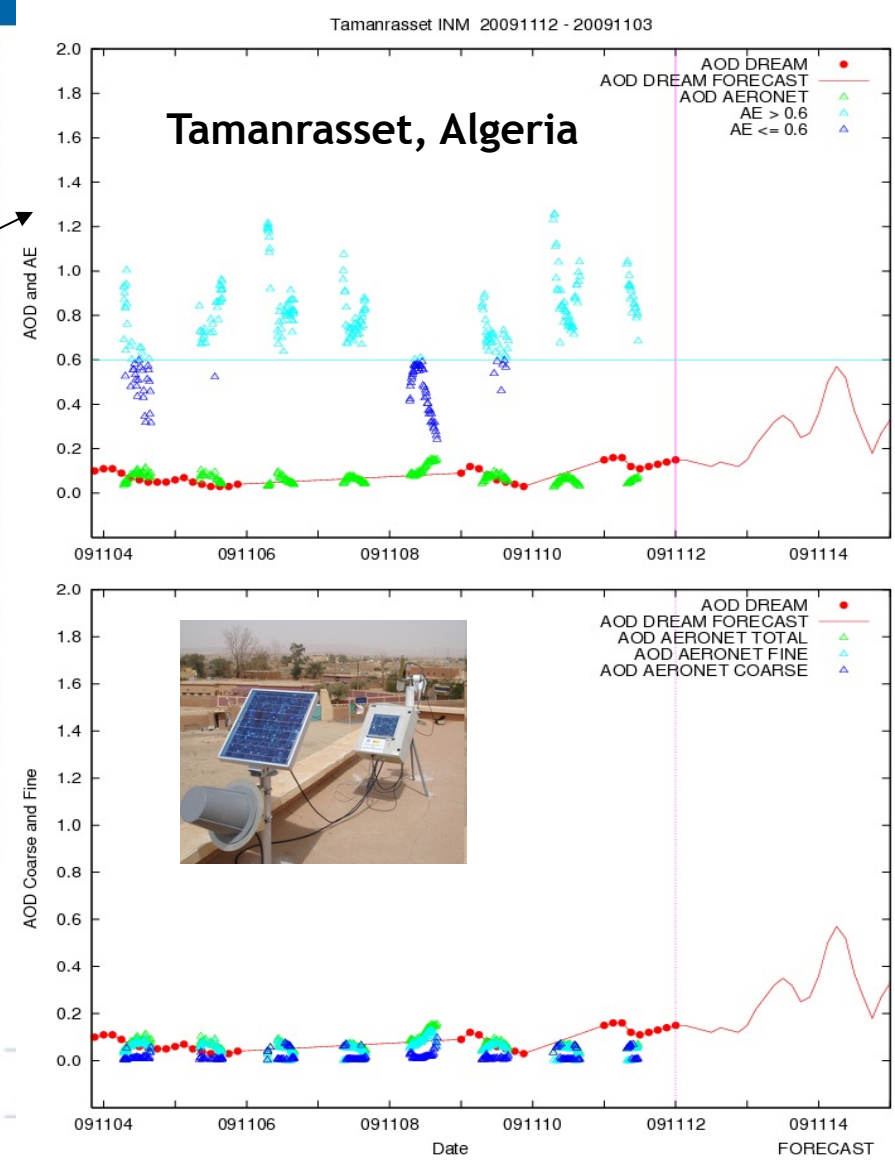
Adam et al., 2010 (ACP-Interlaken): Detection of the Saharan dust air layer in the North Atlantic free troposphere with AERONET, OMI and in-situ data at Izaña Atmospheric Observatory

AERONET
Verification

MSG
Verification

MODIS
Real-Time

Sites Back to Map Archive Plot info



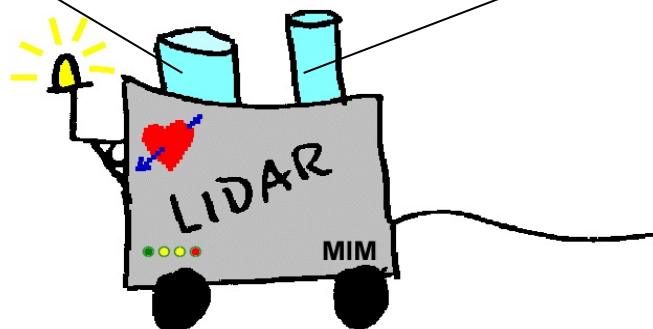
Near real time model verification against AERONET data

From total column observations...
to vertical resolved observations

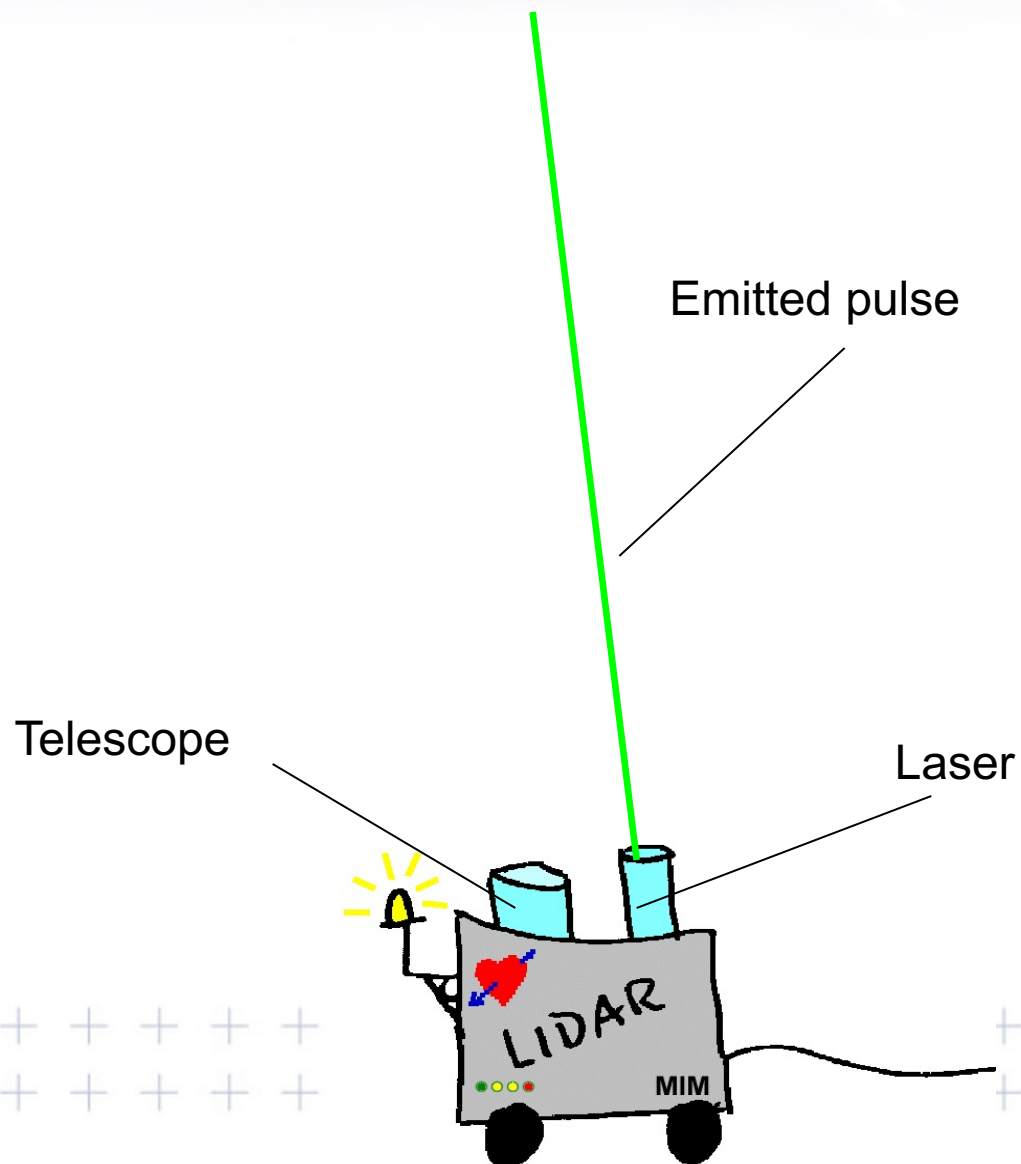
Lidars

Telescope

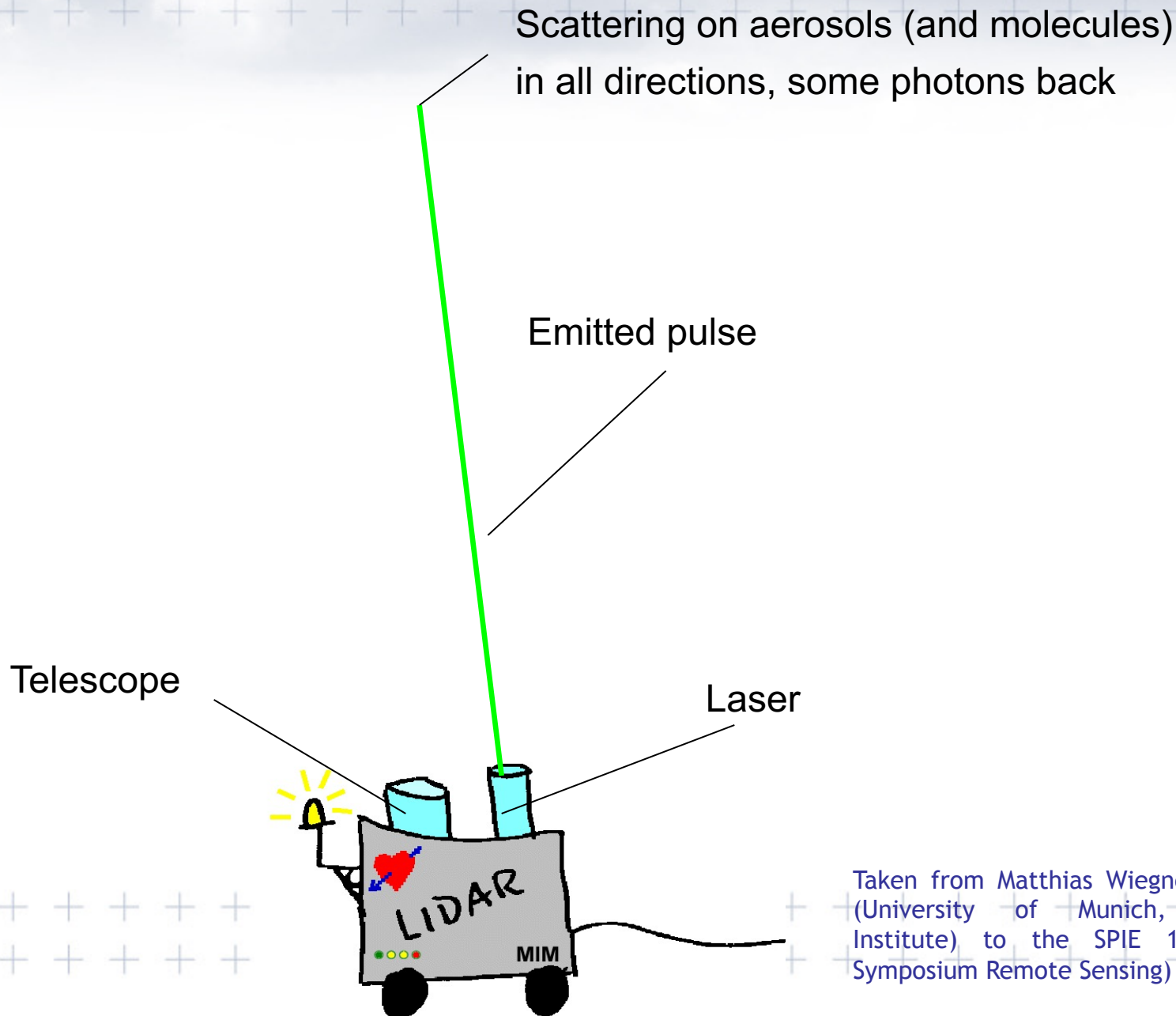
Laser



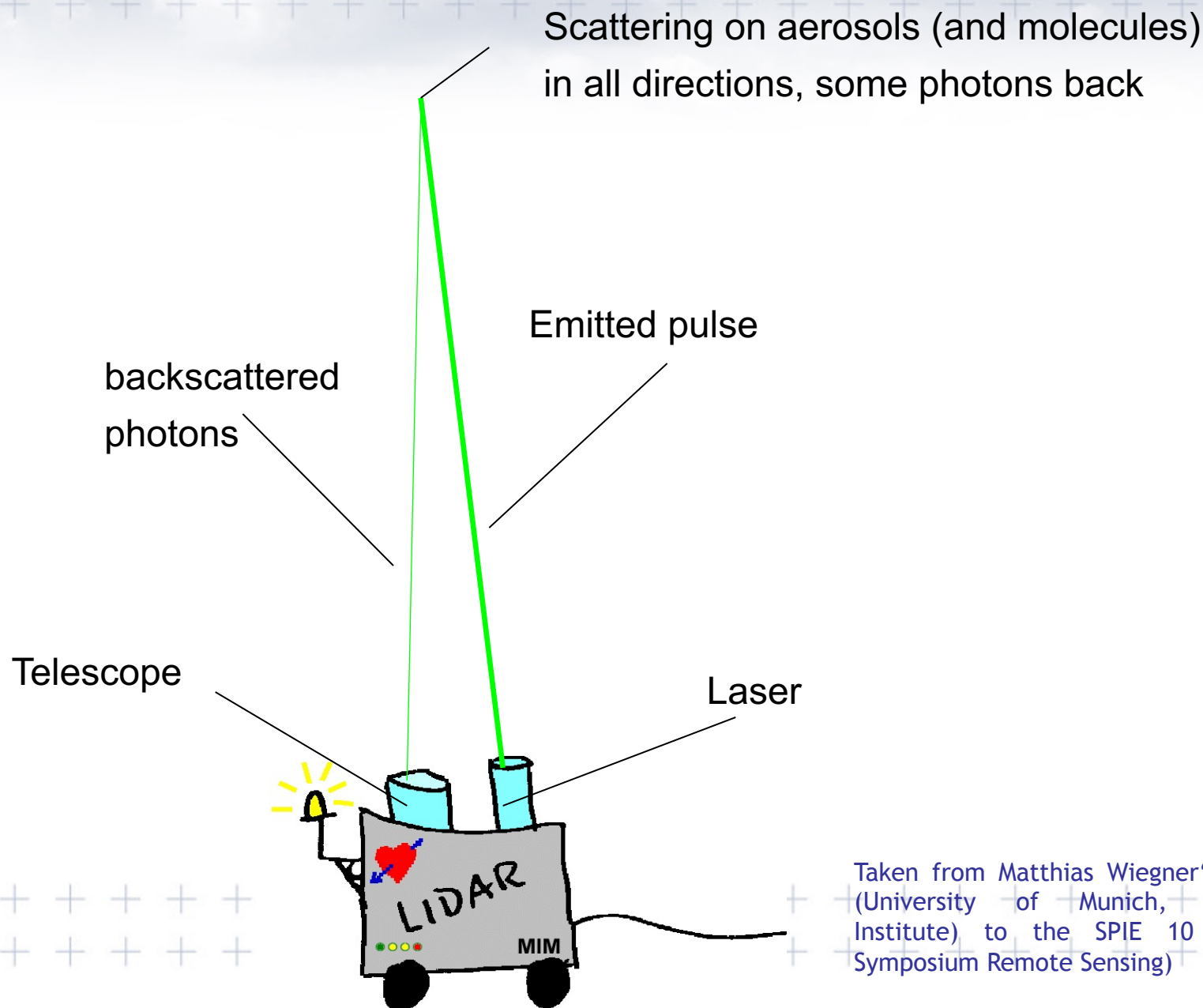
Taken from Matthias Wiegner's presentation
(University of Munich, Meteorological
Institute) to the SPIE 10 (International
Symposium Remote Sensing)



Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)



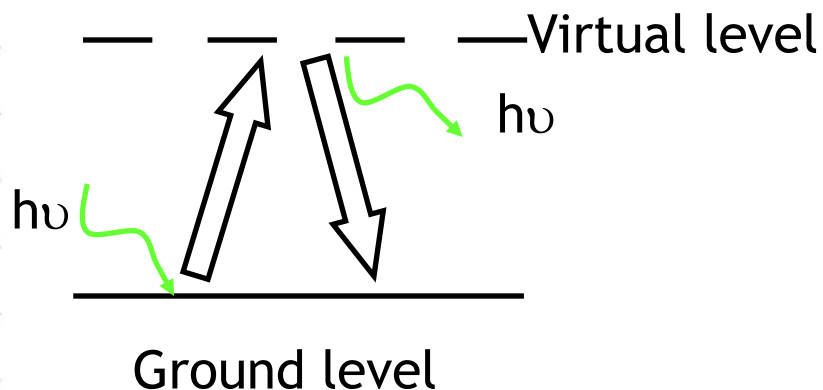
Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)



Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)

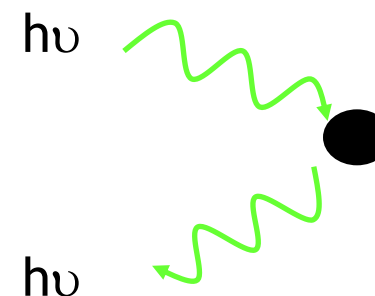
- Rayleigh Scattering

“Laser radiation elastically scattered from atoms or molecules is observed with no change of frequency”



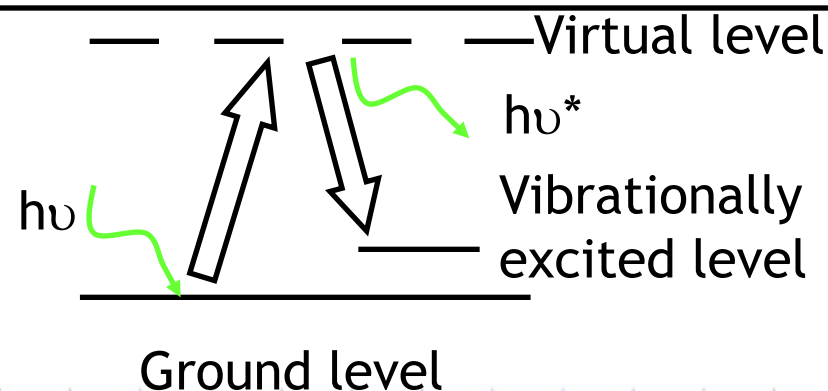
- Mie Scattering

“Laser radiation elastically scattered from small particulates or aerosols (of size comparable to wavelength of radiation) is observed with no change in frequency”



- Raman Scattering

“Laser radiation inelastically scattered from molecules is observed with a frequency shift characteristic of the molecule ($h\nu - h\nu^* = E$)”



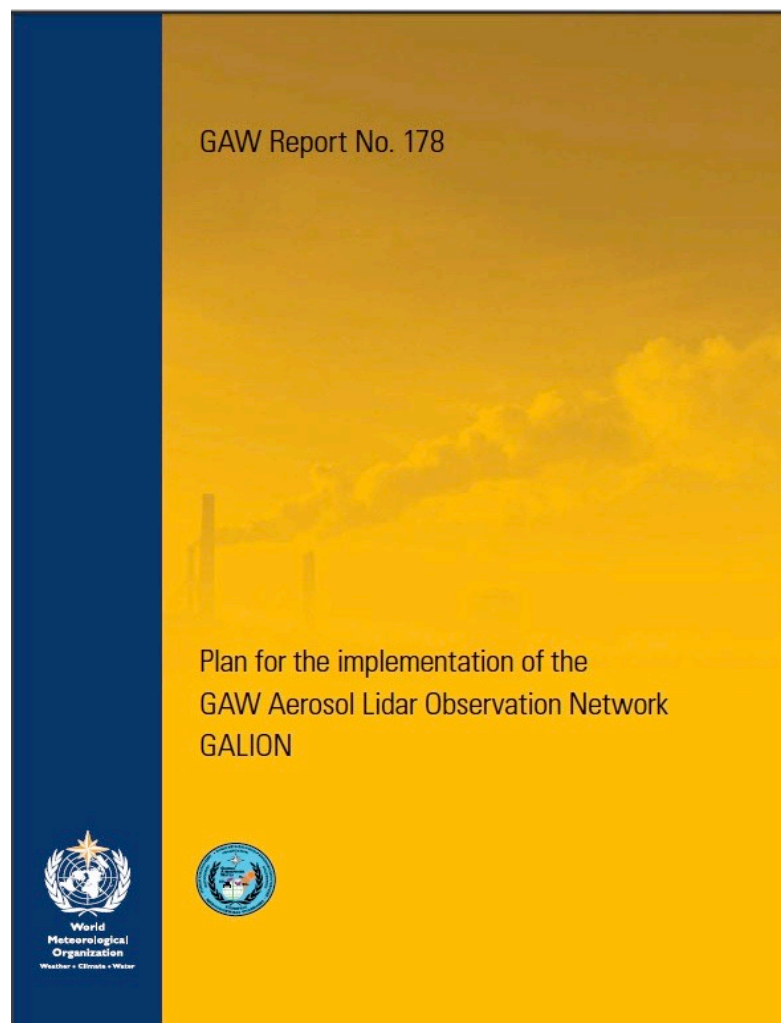


Lidar-Barcelona (UPC)
Raman Lidar
EARLINET-SPALINET



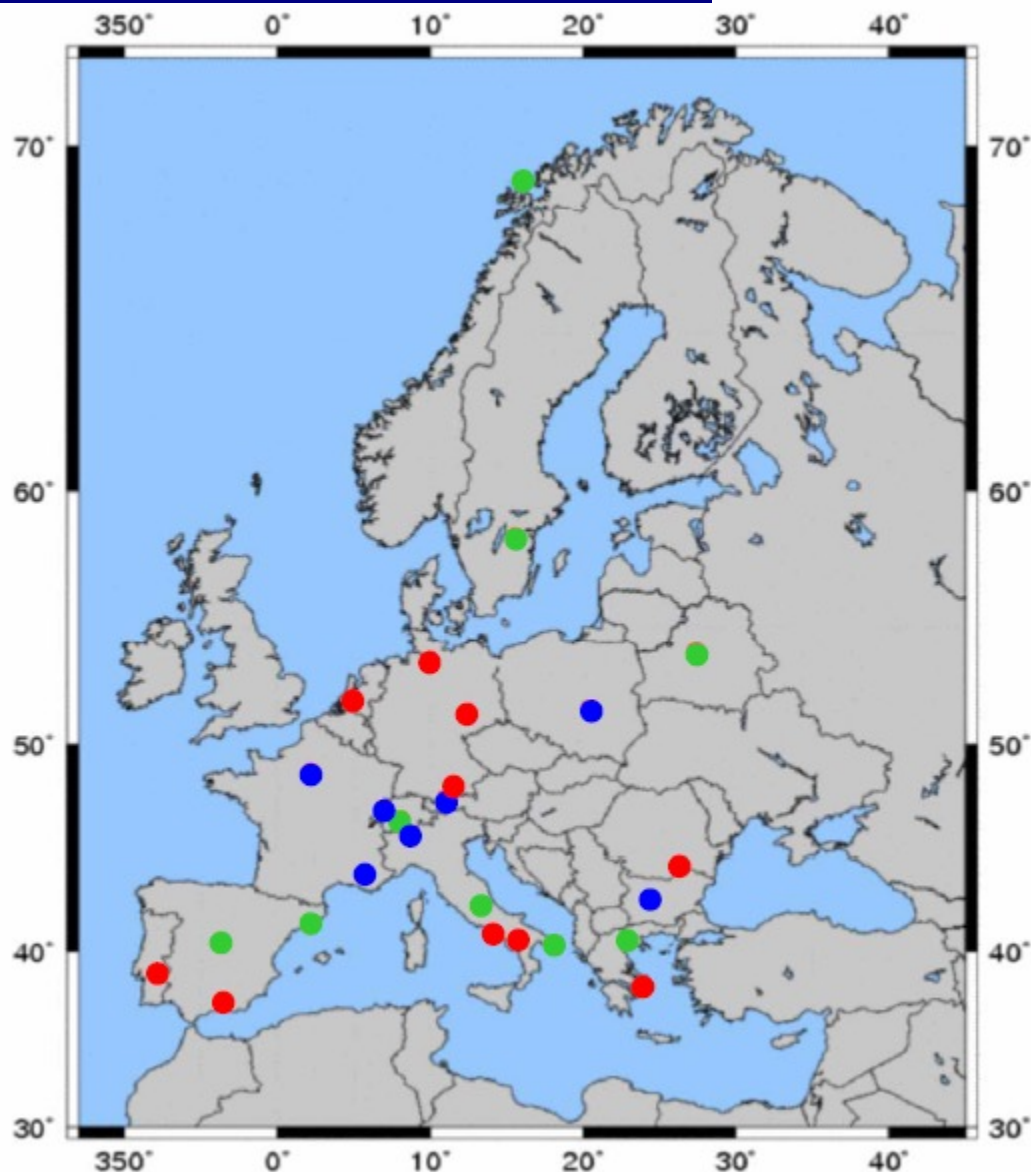
Lidar-Tenerife (INTA-AEMET); Elastic lidar
MPLNET

GAW Atmospheric Lidar Network (GALION)



[ftp://ftp.wmo.int/Documents/
PublicWeb/arep/gaw/gaw178-
galion-27-Oct.pdf](ftp://ftp.wmo.int/Documents/PublicWeb/arep/gaw/gaw178-galion-27-Oct.pdf)





EARLINET

EARLINET (European Aerosol Research Lidar NETwork) is a network of advanced lidar stations distributed over Europe with the main goal to provide a comprehensive, quantitative, and statistically significant data base for the aerosol distribution on a continental scale. EARLINET provides independent measurements of aerosol extinction and backscatter, and retrieval of aerosol microphysical properties.

10 EARLINET stations are equipped also with sunphotometers (they are part of AERONET).

26 lidar stations

- 10 multiwavelength Raman lidar stations

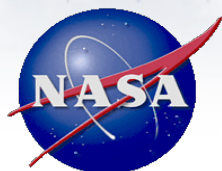
backscatter (355, 532 and 1064 nm)
+ extinction (355 and 532 nm) +
depolar ratio (532 nm)

- 9 Raman lidar stations

- 7 single backscatter lidar stations

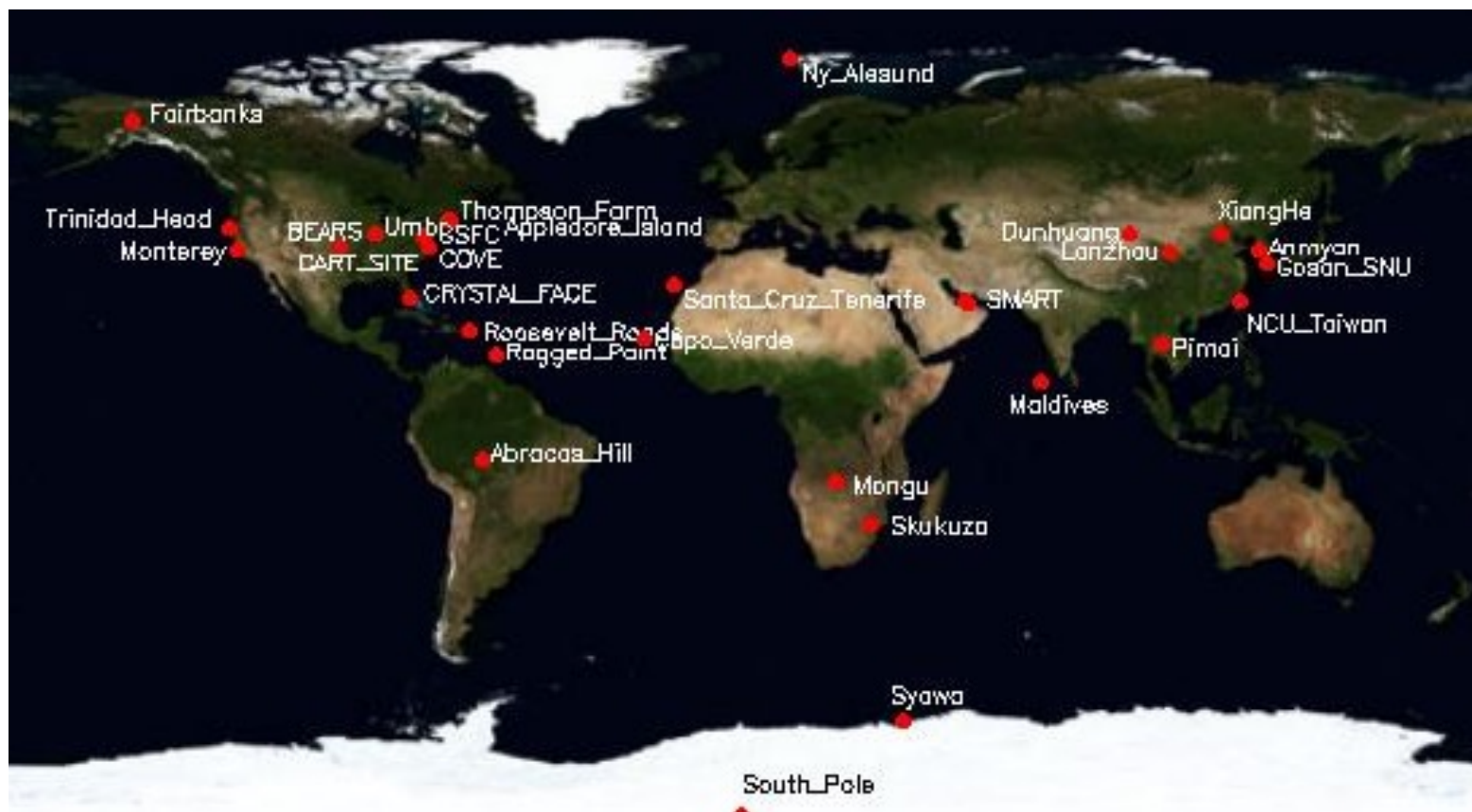
Aerosol lidar (MPLNet)

<http://mplnet.gsfc.nasa.gov/>

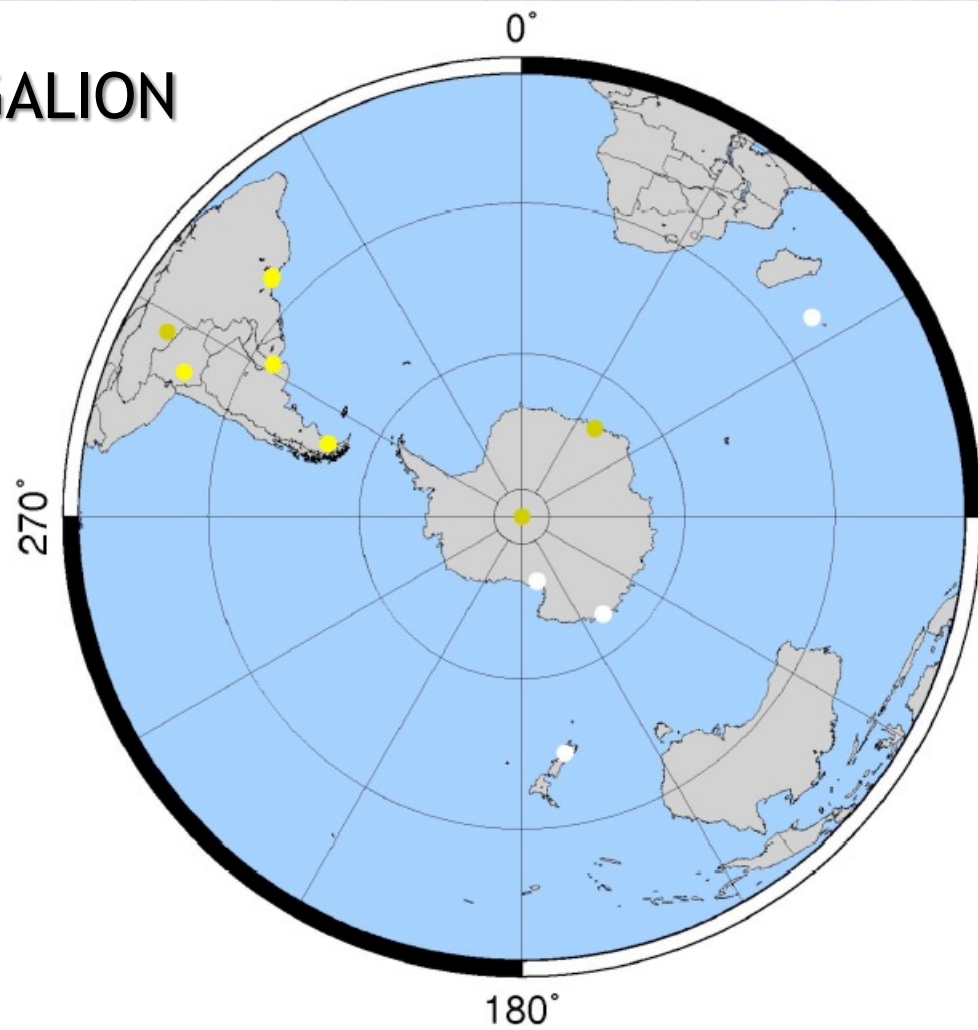
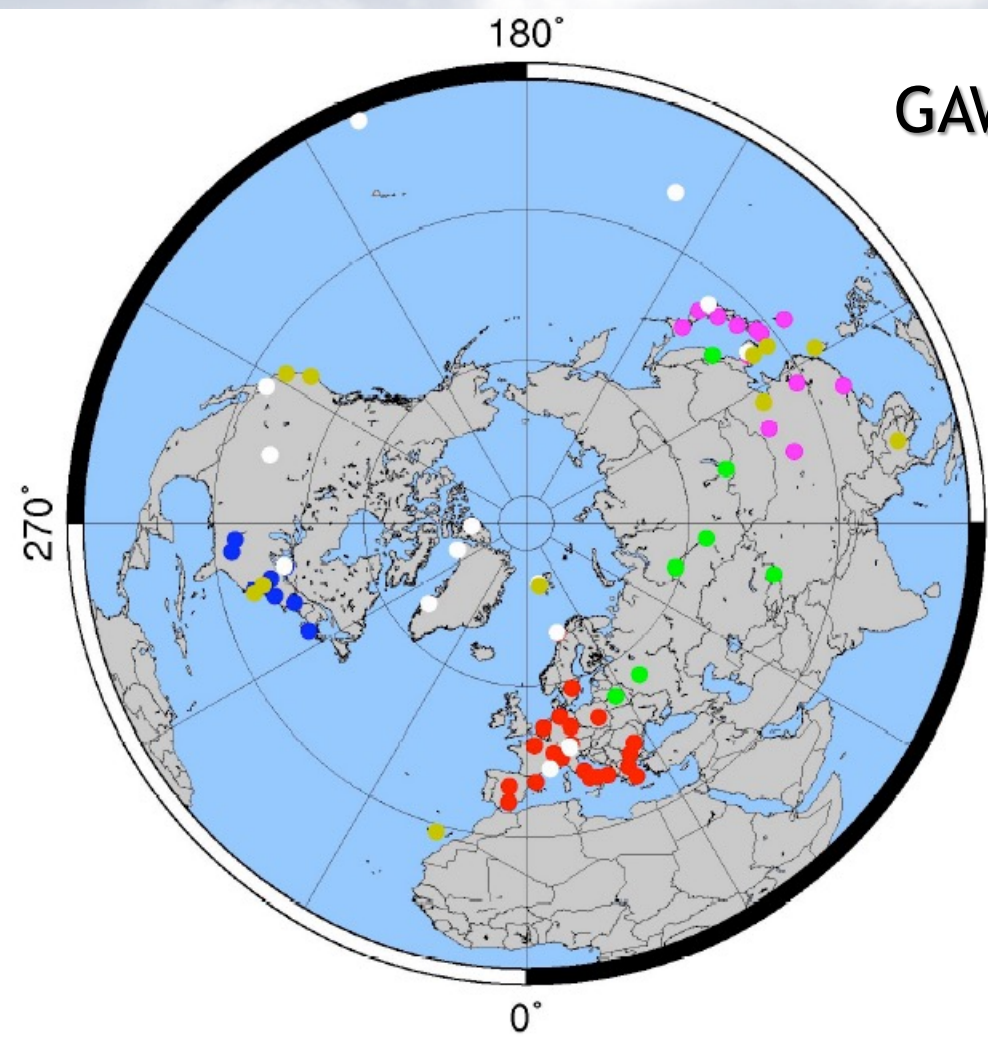


523 nm MPLNET

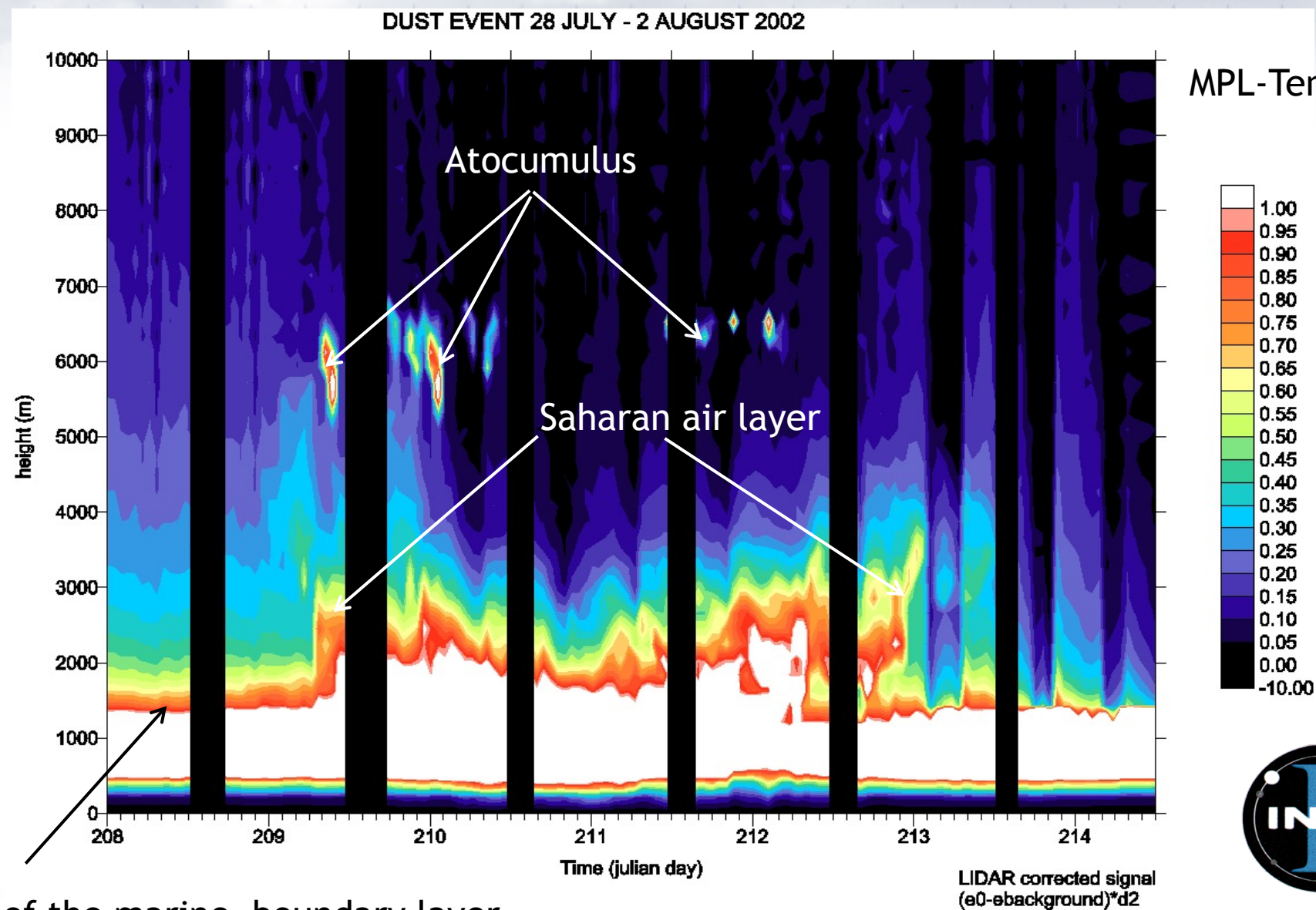
Automatized since July 2005



GAW-GALION

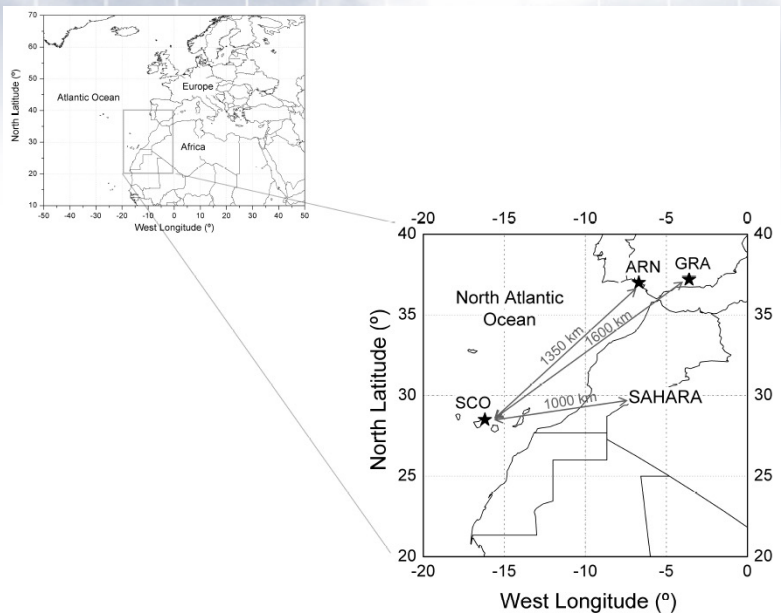


Distribution of stations as available through the cooperation between existing networks: **AD-NET** , **ALINE** , **CISLiNet** , **EARLINET** , **MPLNET** , NDACC , **REALM** .



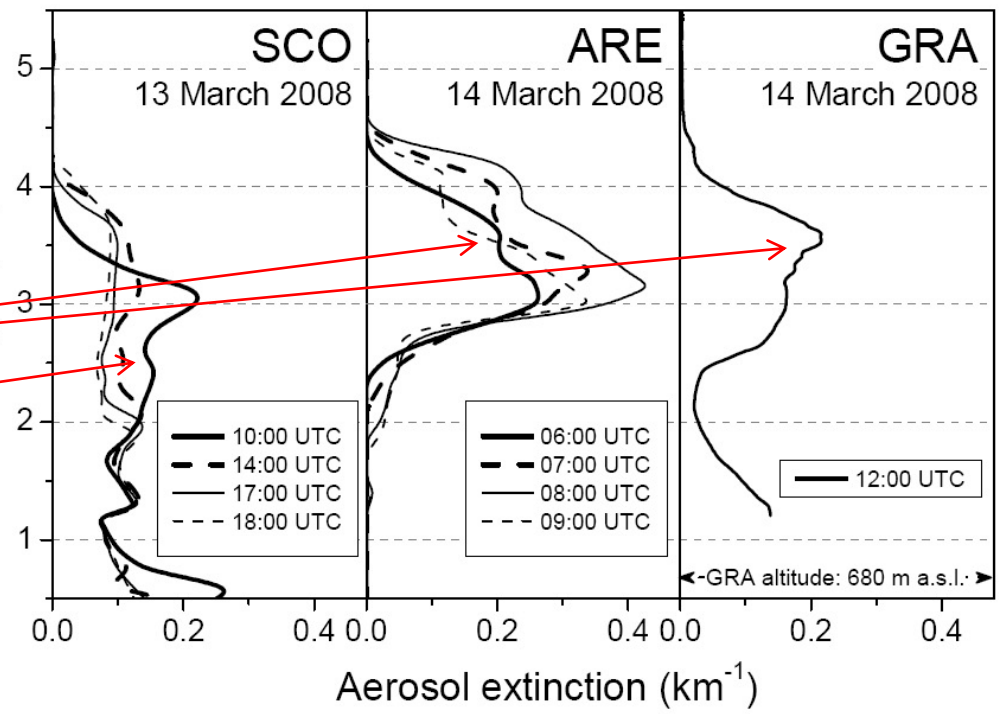
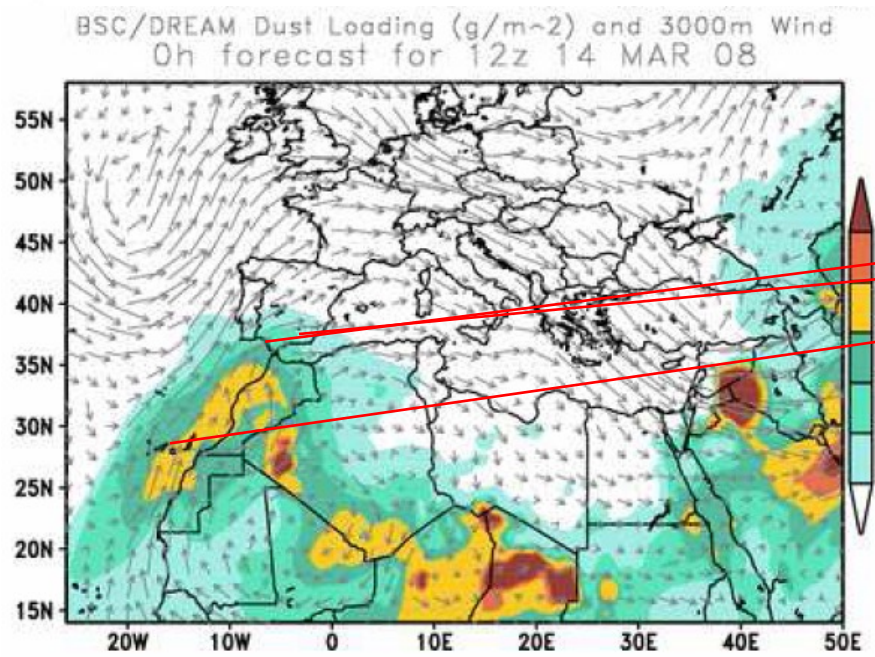
MPL-Tenerife





A case study of dust transport from Canary Islands to Iberian Peninsula

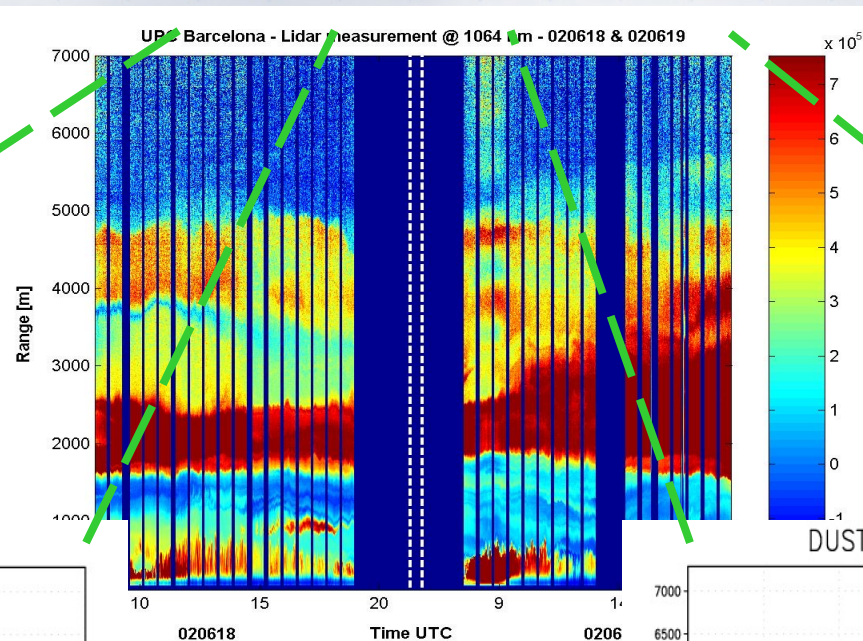
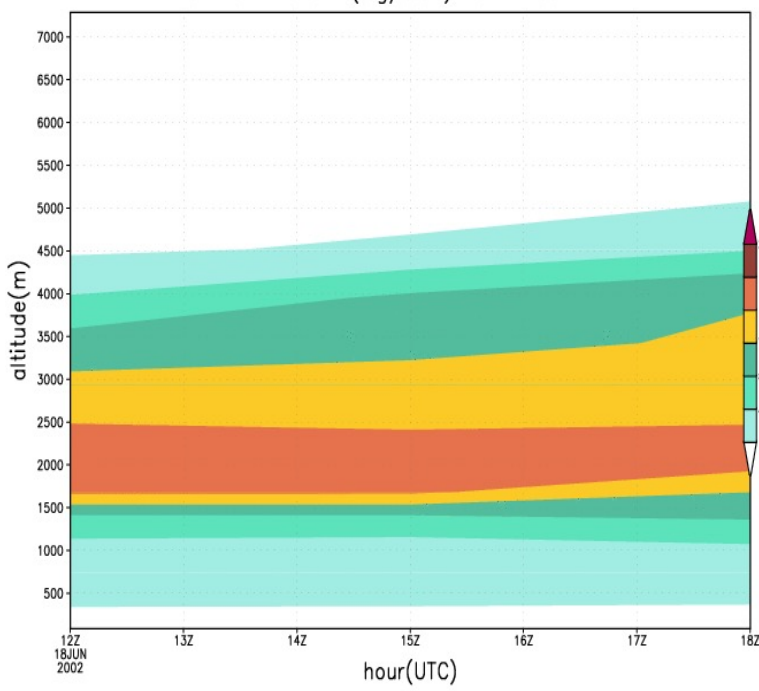
Córdoba-Jabonero et al., ACP Discuss., 2010



Barcelona lidar vs DREAM BSC



DUST CONC. ($\mu\text{g}/\text{m}^3$) 18 JUN 2002

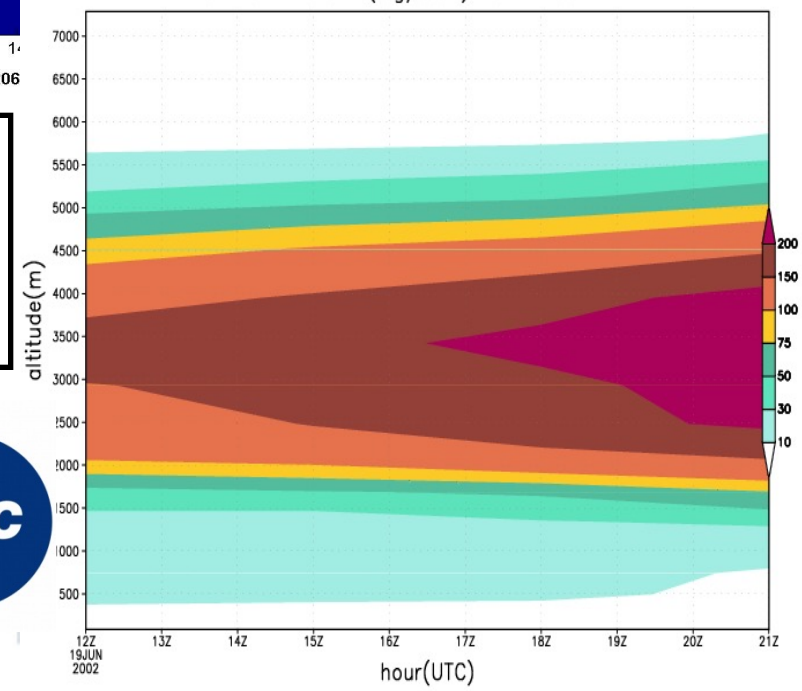


EARLINET: Lidar-UPC,
Barcelona

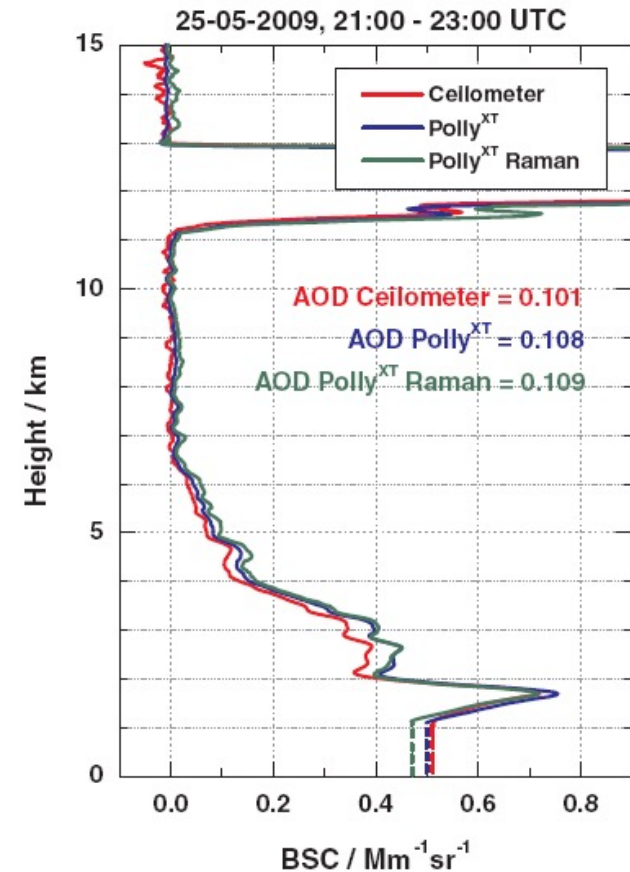
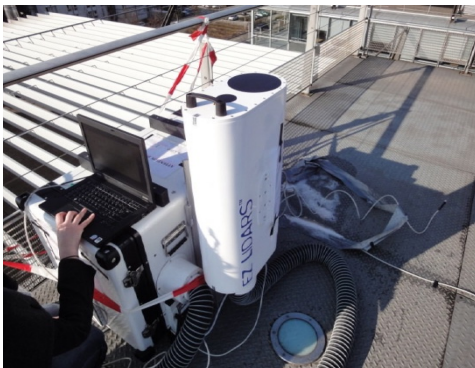
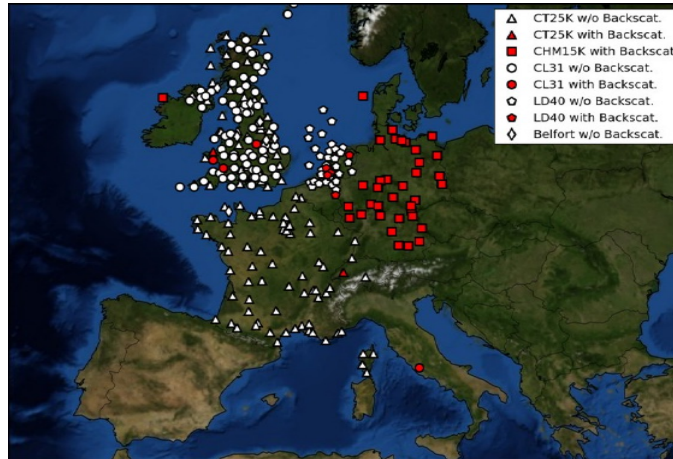
18-19 June 2002

DUST CONC. ($\mu\text{g}/\text{m}^3$) 19 JUN 2002

Vertical dust
distribution
validation:
AIRLINET-DREAM

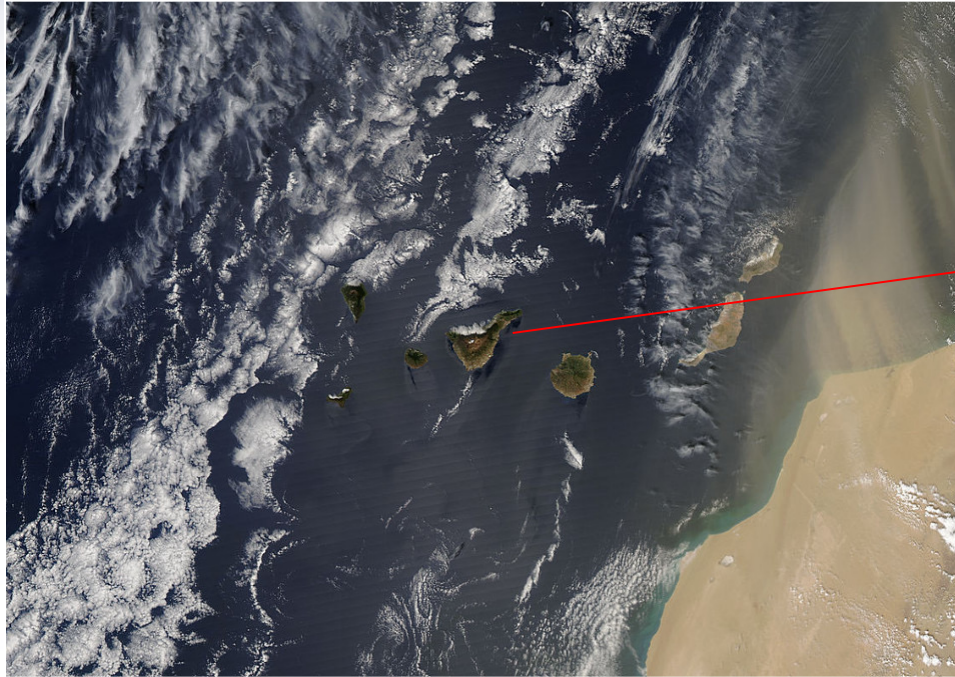


Met Services are replacing cloud-base ceilometer networks by aerosol backscatter profiling ceilometers (**IR wavelength**).
Objective: To monitor MLD based on several hundred profiling ceilometers (100km sampling)



*Heese et al., Atmos. Mes. Tech. 2010,
Ceilometer-lidar inter-comparison:
backscatter coefficient retrieval and
signal-to-noise ratio determination*

Optimal for desertic areas !!



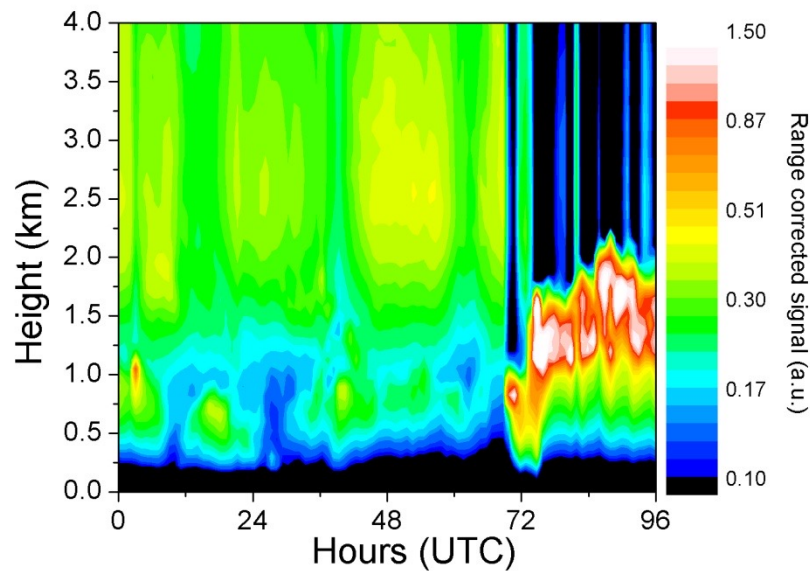
Viasala Ceilometer
CL-51

MicroPulse Lidar and Ceilometer inter-comparison during Saharan dust intrusions over the Canary Islands

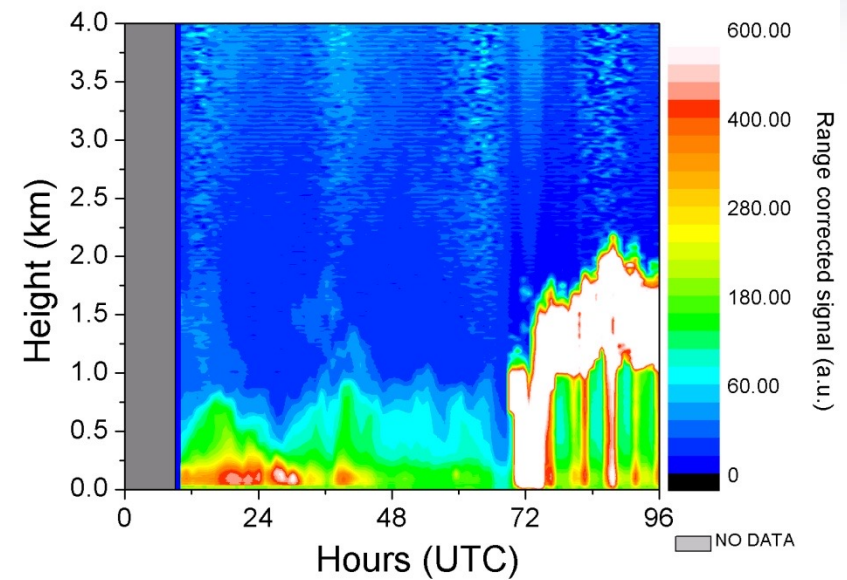
Y. Hernández, S. Alonso-Pérez, E. Cuevas, C. Camino, R. Ramos, J. de Bustos, C. Marrero, C. Córdoba-Jabonero and M. Gil (2011)

Campaign performed from January to March 2011 in Tenerife island

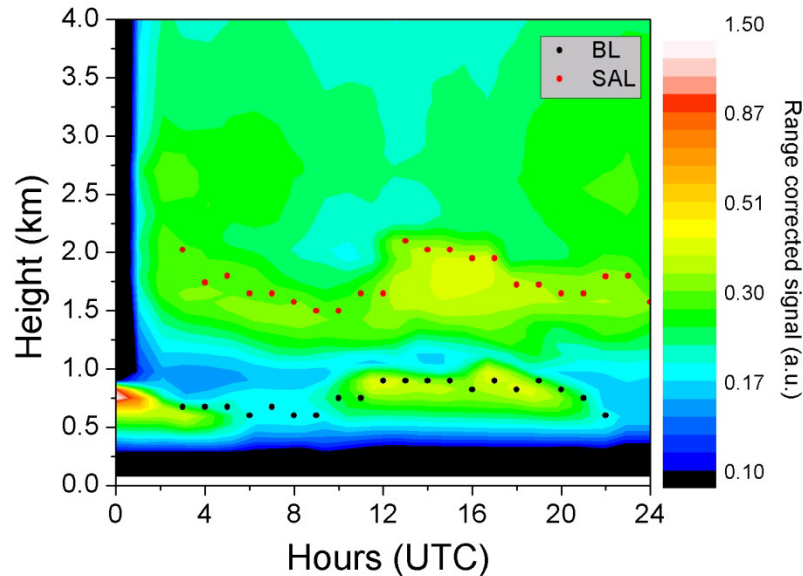
MPL-3 - Sta. Cruz de Tenerife. Mar 31- Apr 3, 2011



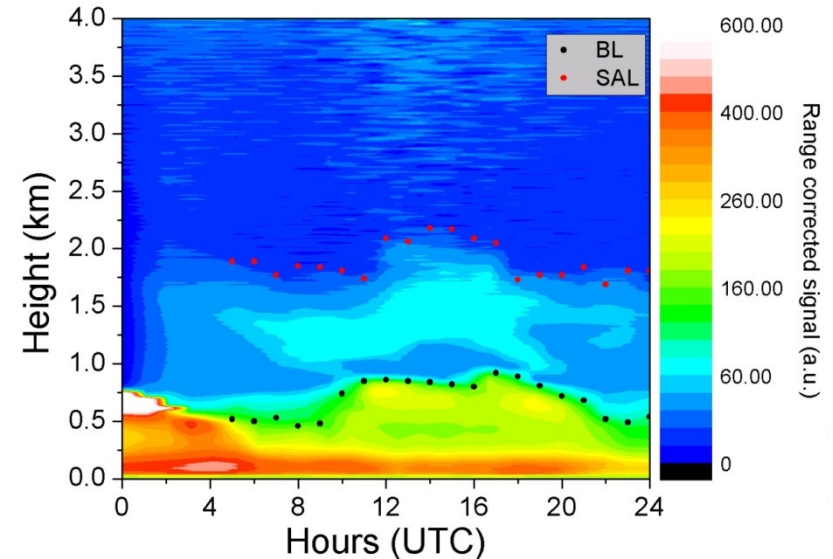
CL51 - Sta. Cruz de Tenerife. Mar 31- Apr 3, 2011



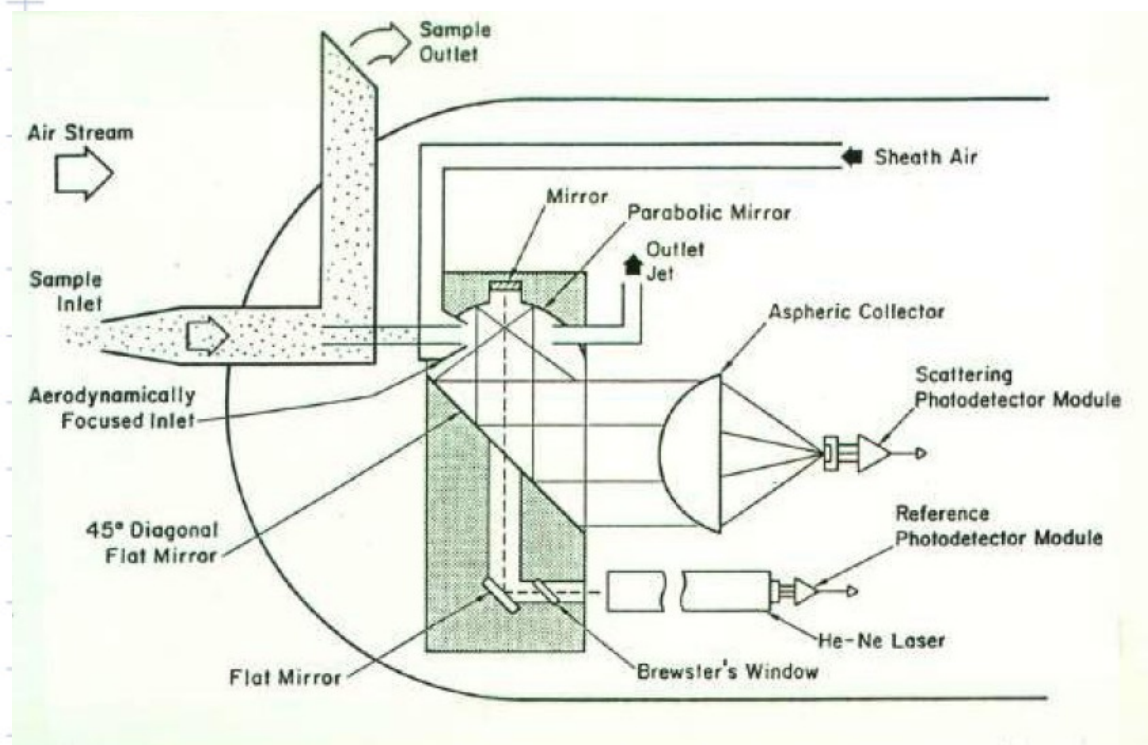
MPL-3 - Sta. Cruz de Tenerife. Feb 24, 2011



CL51 - Sta. Cruz de Tenerife. Feb 24, 2011



INTA C212-200 N/S 301



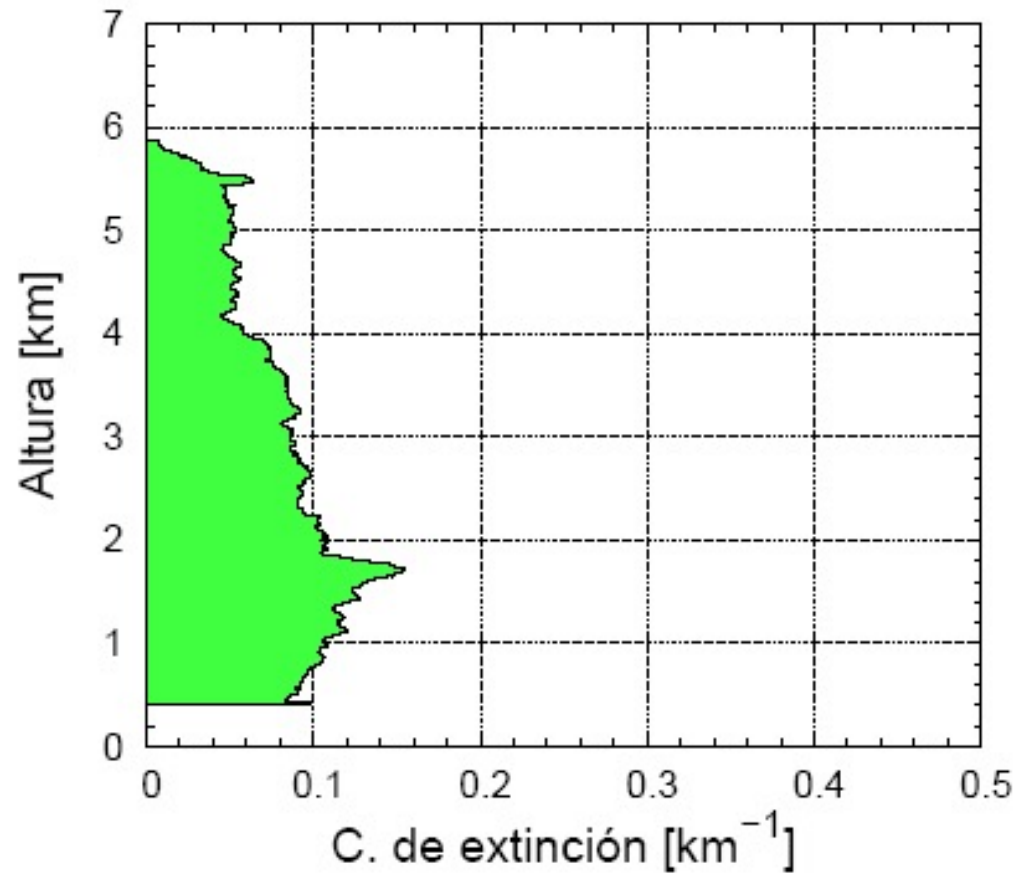
PCASP-100X sonde

Optical counter
0.1 - 3.0 μm en 15 channels
Up to 20,000 particles /s
A size distribution /s

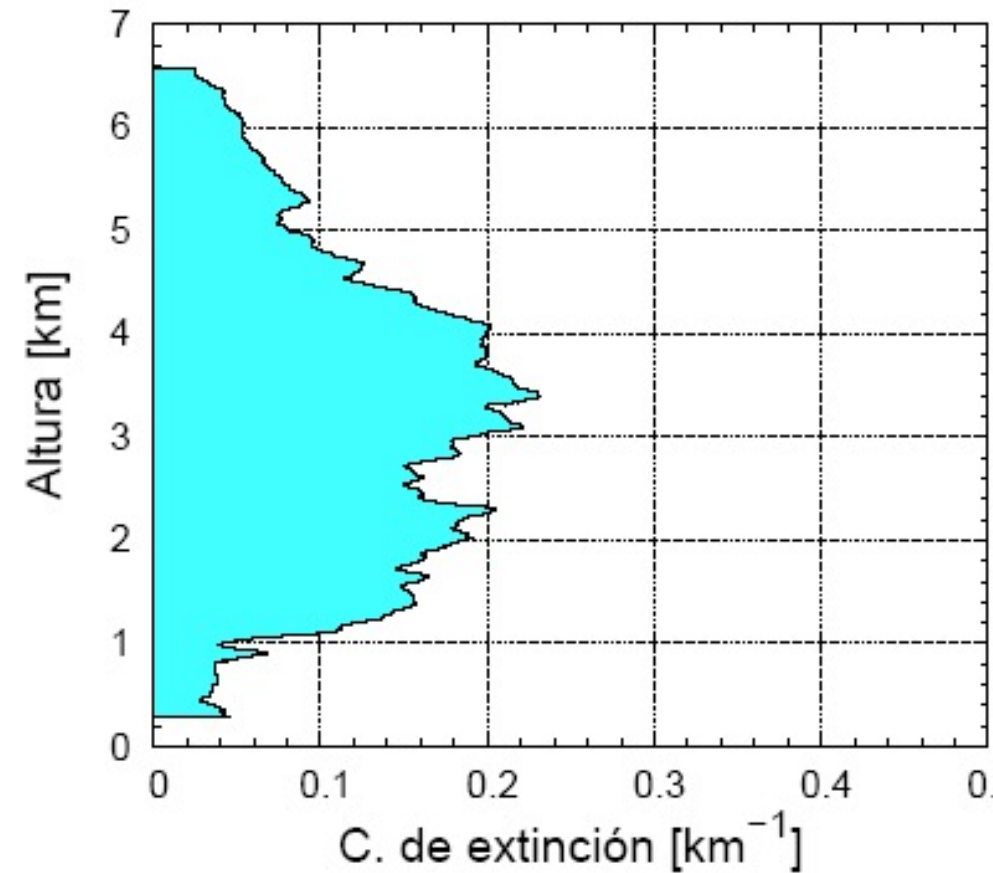




20 - julio - 05



22 - julio - 06



Andrey, 2011 (PhD)

In-situ measurements and surface remote sensing compared to satellite

Advantages

Disadvantages

In-situ
measurements

Ground-based
remote sensing

Satellite remote
sensing

In-situ measurements and surface remote sensing compared to satellite

Advantages

In-situ measurements

- very straightforward;
- unique dust physical and chemical information;
- universal applicability (no sky conditions dependent)
- Time high resolution (minutes)

Ground-based remote sensing

- high information on dust (transmitted light dominates over reflected);
- non-intrusive measurements;
- easy access to equipment;
- column dust information

Satellite remote sensing

- global coverage; (global dust)
- non-intrusive measurements

Disadvantages

- intrusive measurements;
- local coverage in some sites

- local coverage;
- indirect measurements;
- very limited capability in presence of clouds (Photom.)

- limited on information aerosol (aerosol and surface effects to be separated);
- no access to equipment

Ground-based dust observations are essential for:

- High-accuracy optical and chemical characterization of dust
- Dust model verification, validation and assimilation
- Dust satellite-based products validation

Ground based ‘supersites’ equipped with complete high-quality dust observation programs (in-situ and remote sensing) constitute unique platforms for satellite-based dust observations and dust models quality assurance

Thank you for your attention

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